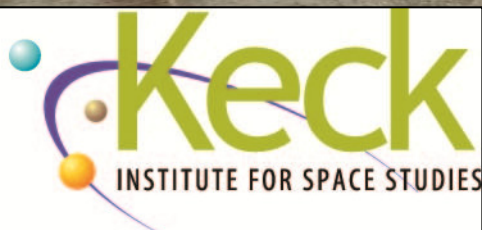


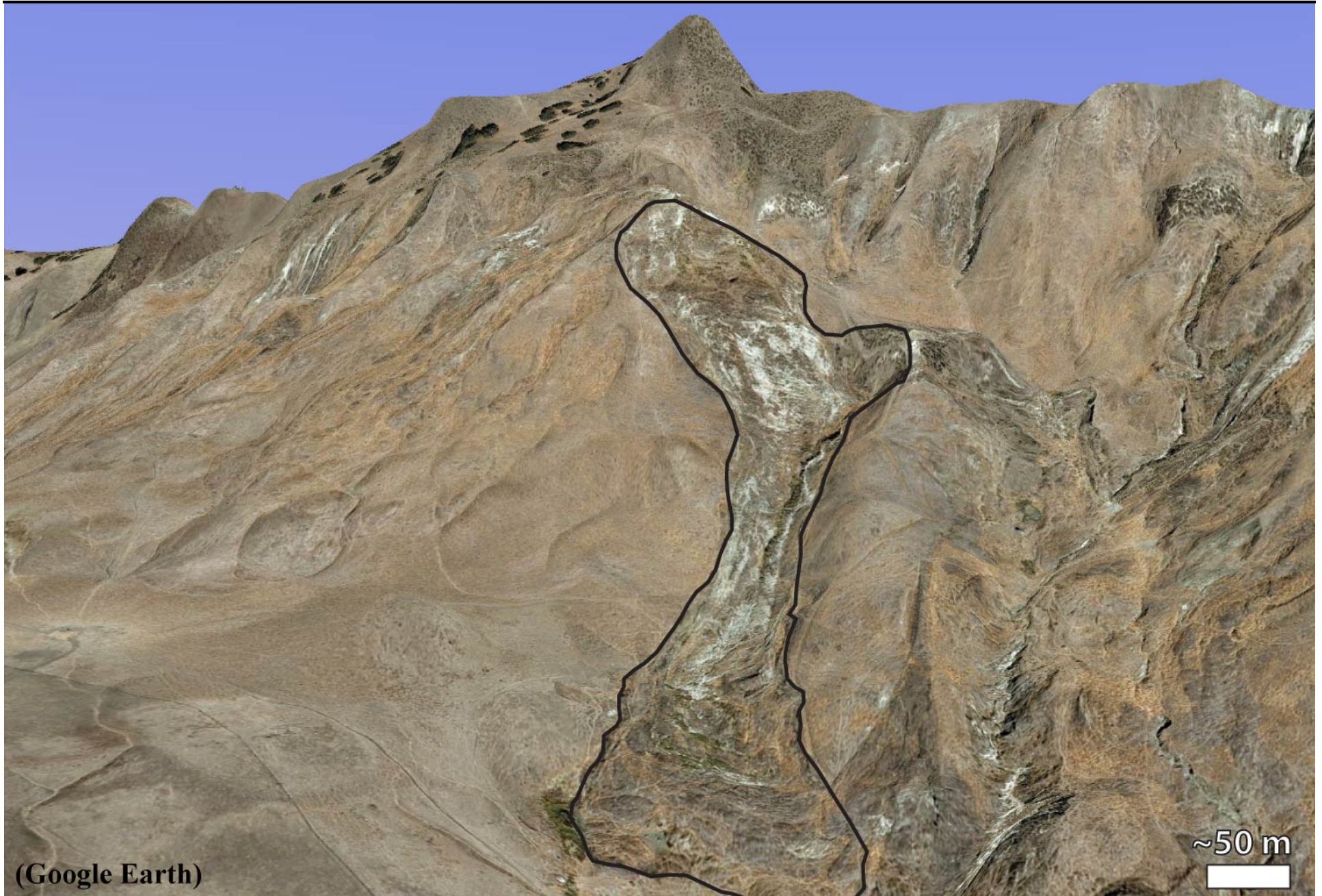
The influence of large-magnitude earthquakes and fault zone damage on the spatial distribution of slow-moving landslides



Joel Scheingross, Brent Minchew, Benjamin Mackey,
Mark Simons, Michael Lamb, Scott Hensley



Slow-moving landslides (“earthflows”)

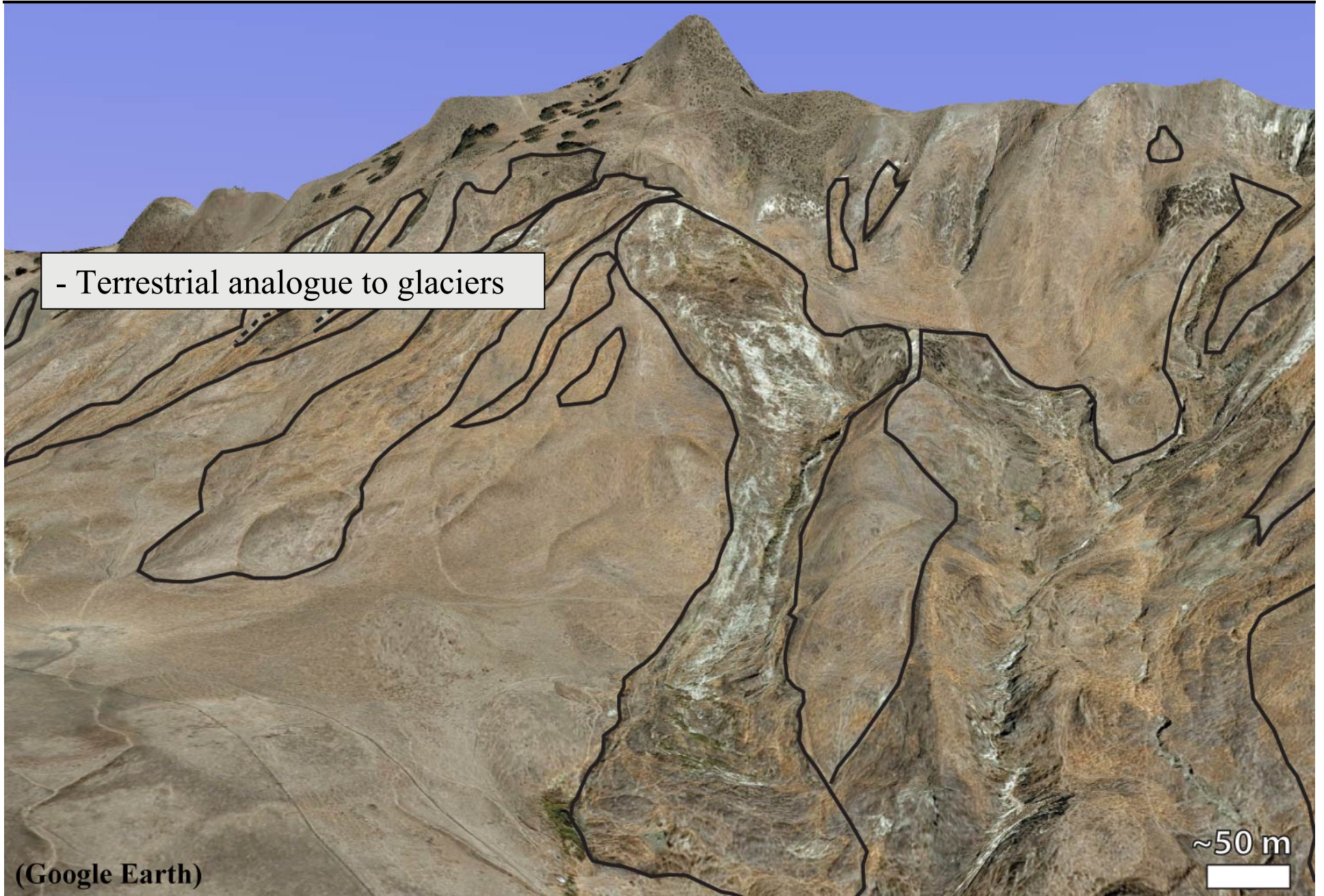


Slow-moving landslides (“earthflows”)



(Google Earth)

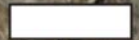
Slow-moving landslides (“earthflows”)



- Terrestrial analogue to glaciers

(Google Earth)

~50 m



Slow-moving landslides (“earthflows”)

- Terrestrial analogue to glaciers
- Move at rates of mm/yr to m/yr



(Google Earth)

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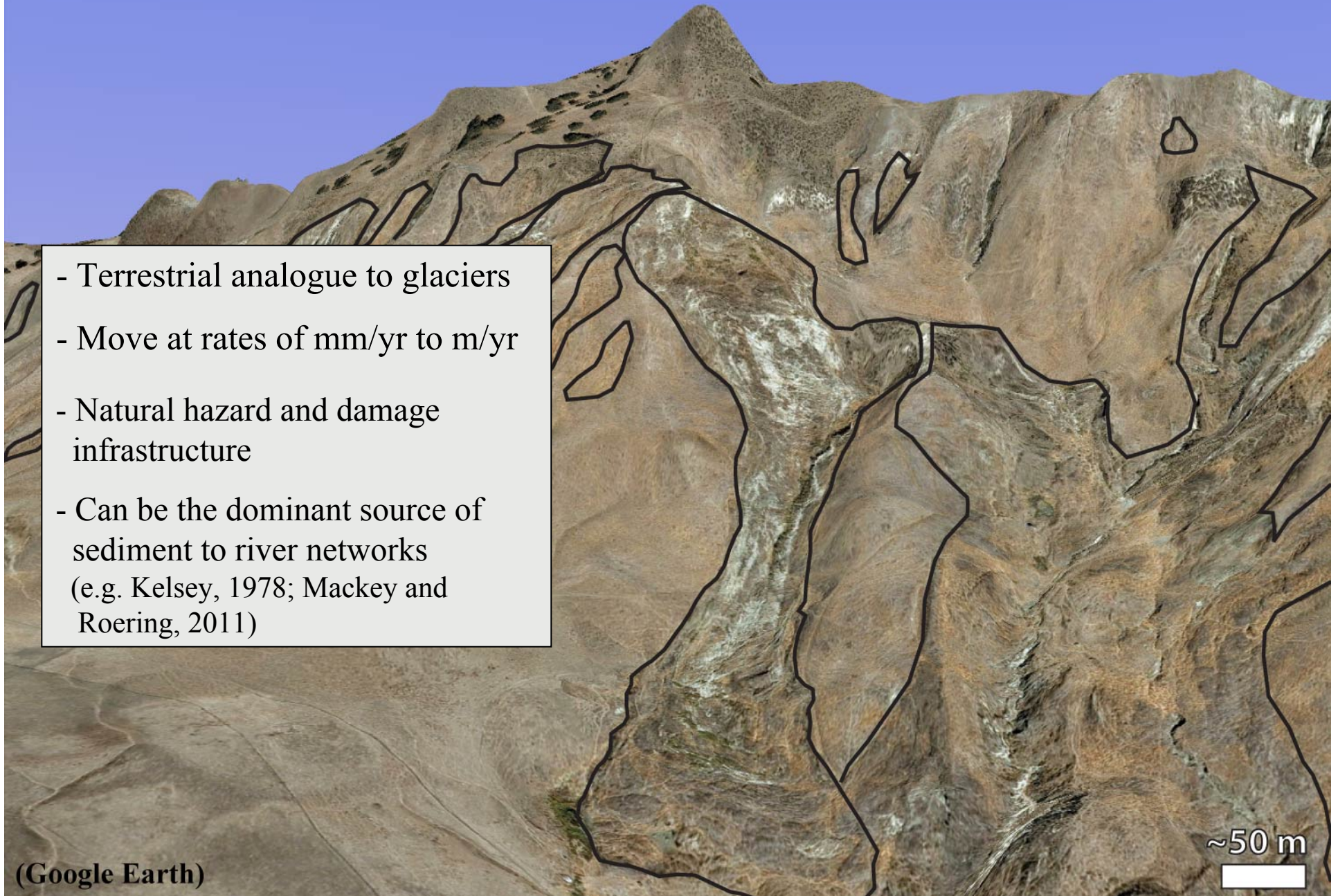
- Terrestrial analogue to glaciers
- Move at rates of mm/yr to m/yr
- Natural hazard and damage infrastructure



(Google Earth)

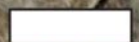
Slow-moving landslides (“earthflows”)

- Terrestrial analogue to glaciers
- Move at rates of mm/yr to m/yr
- Natural hazard and damage infrastructure
- Can be the dominant source of sediment to river networks (e.g. Kelsey, 1978; Mackey and Roering, 2011)

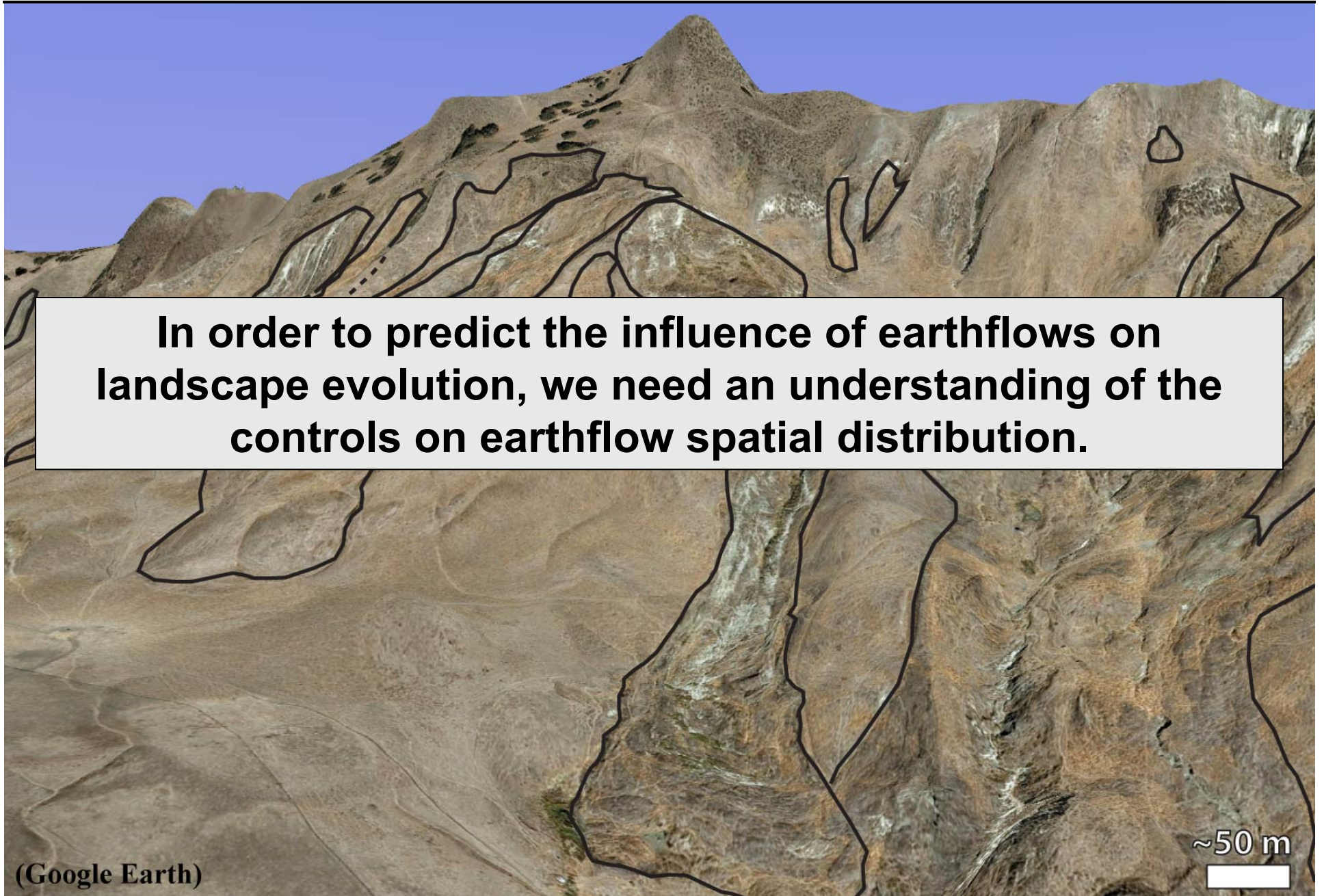


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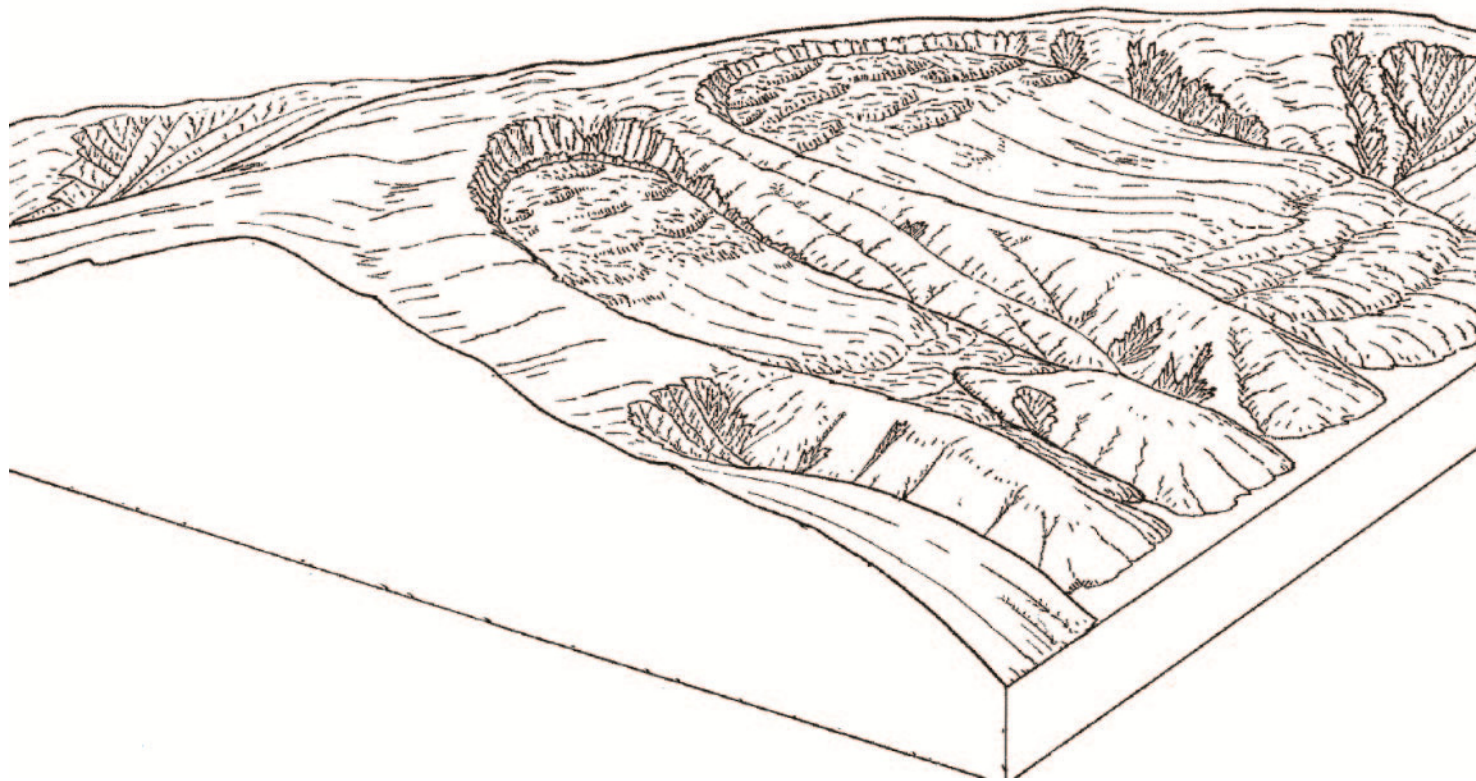


Slow-moving landslides (“earthflows”)



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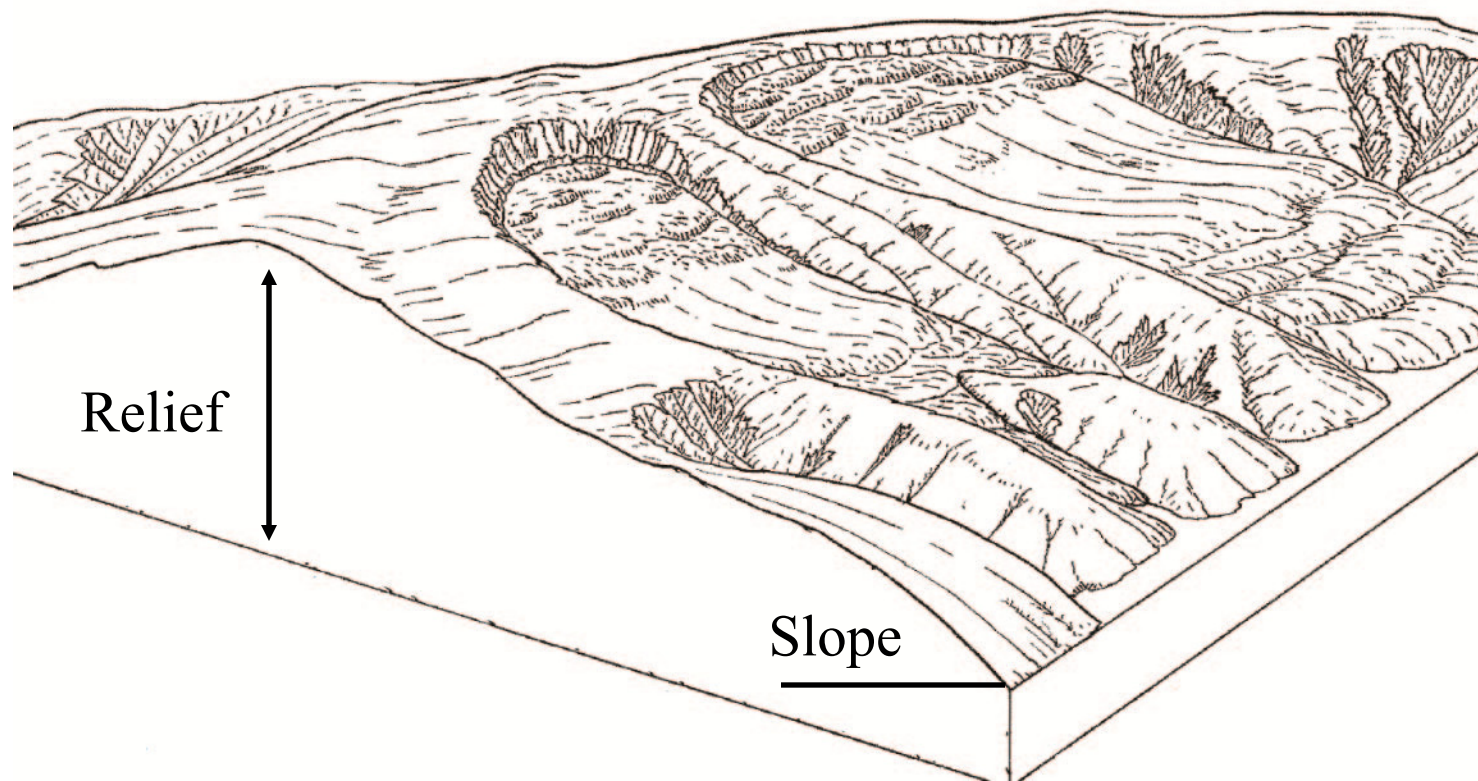
What controls earthflow spatial distribution?



Modified from Putnam and Sharp (1940)

What controls earthflow spatial distribution?

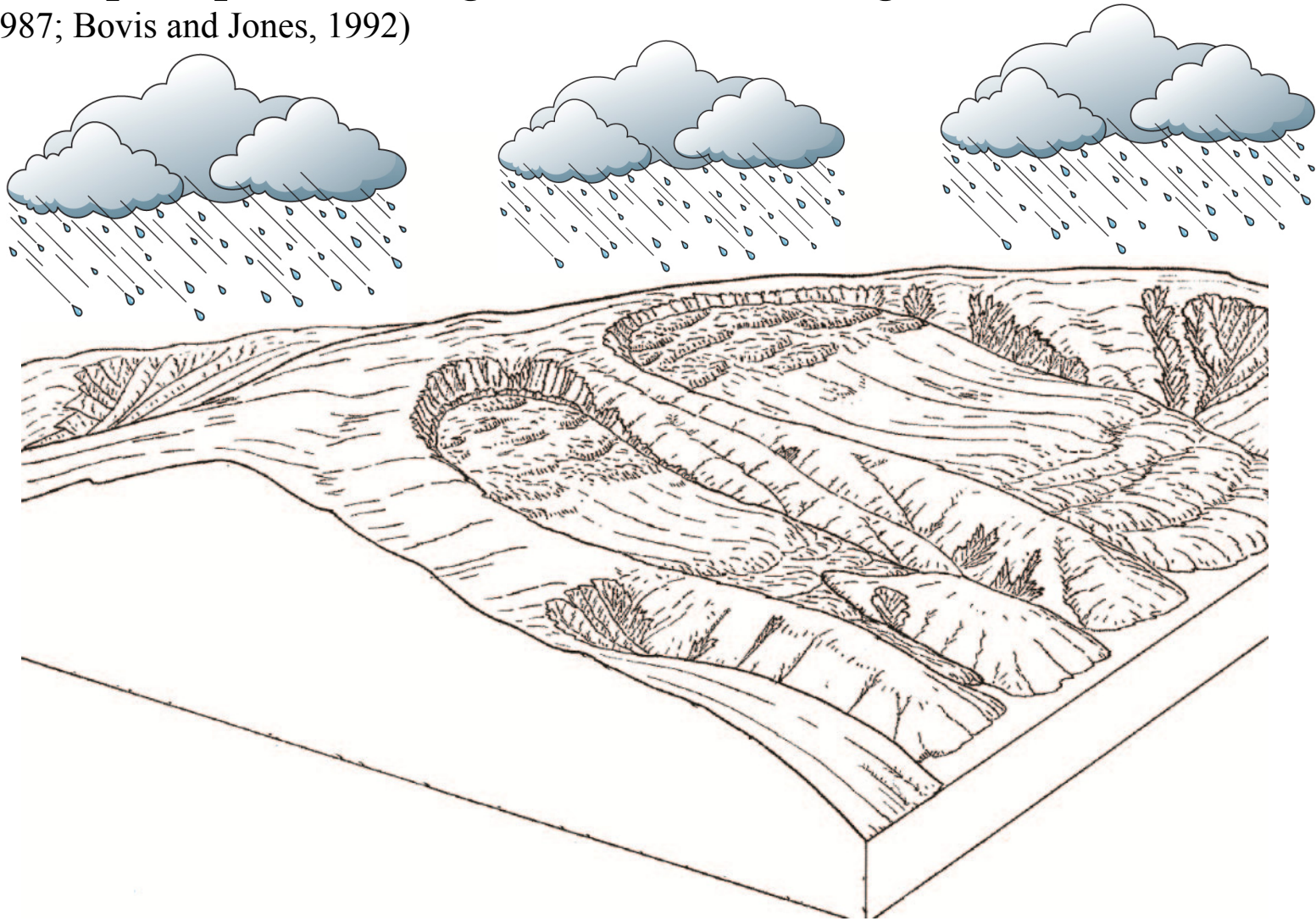
- Topography: hillslope gradient, relief, aspect, uplift (e.g, Keefer and Johnson, 1983; Booth and Roering, 2011; Mackey and Roering, 2011)



Modified from Putnam and Sharp (1940)

What controls earthflow spatial distribution?

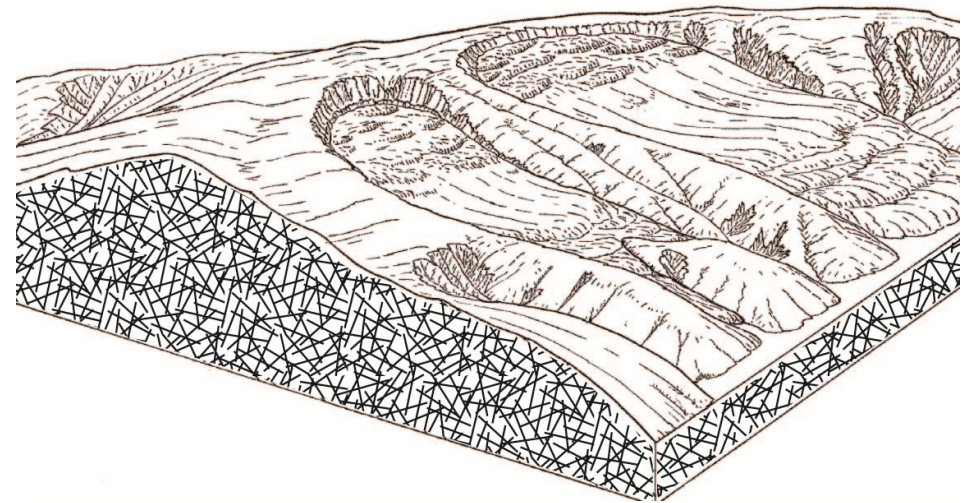
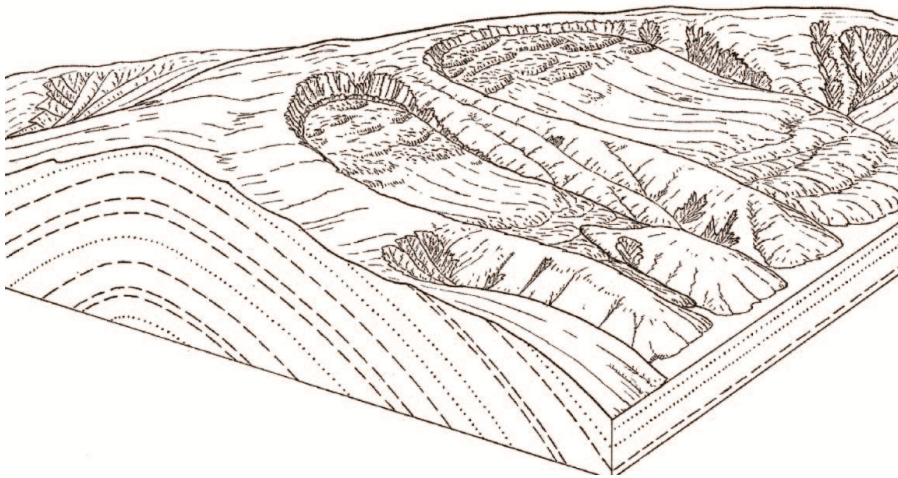
- Climate: precipitation, vegetation, weathering (e.g. Kelsey, 1978; Iverson and Major, 1987; Bovis and Jones, 1992)



Modified from Putnam and Sharp (1940)

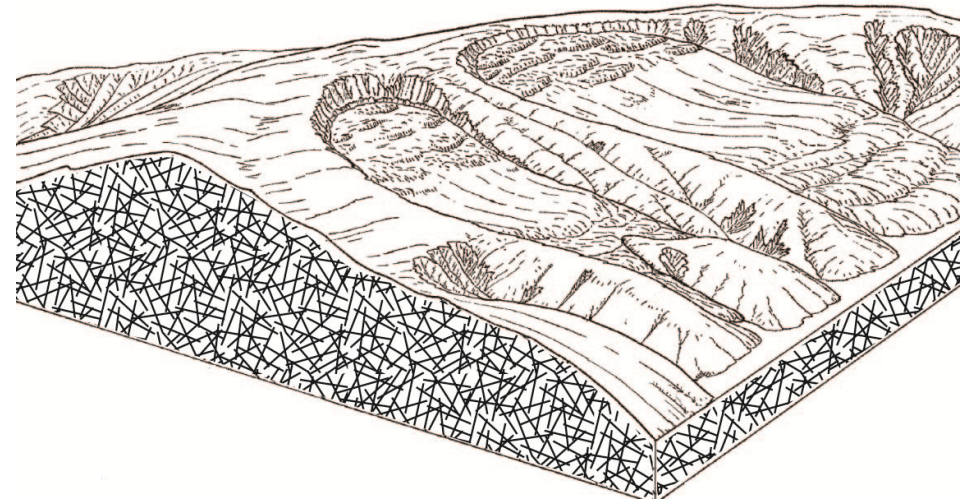
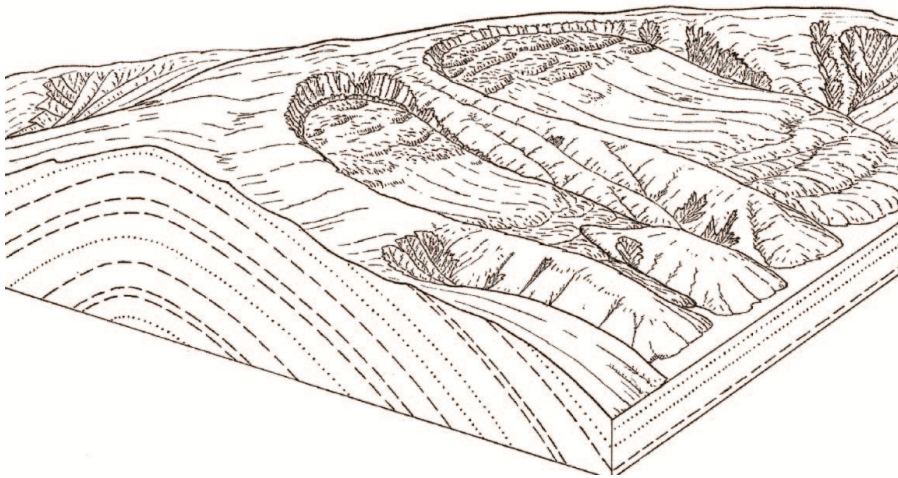
What controls earthflow spatial distribution?

- Lithology: rock strength, structure, clay content, soil thickness (e.g., Putnam and Sharp, 1940; Kelsey, 1978; Keefer and Johnson, 1983)



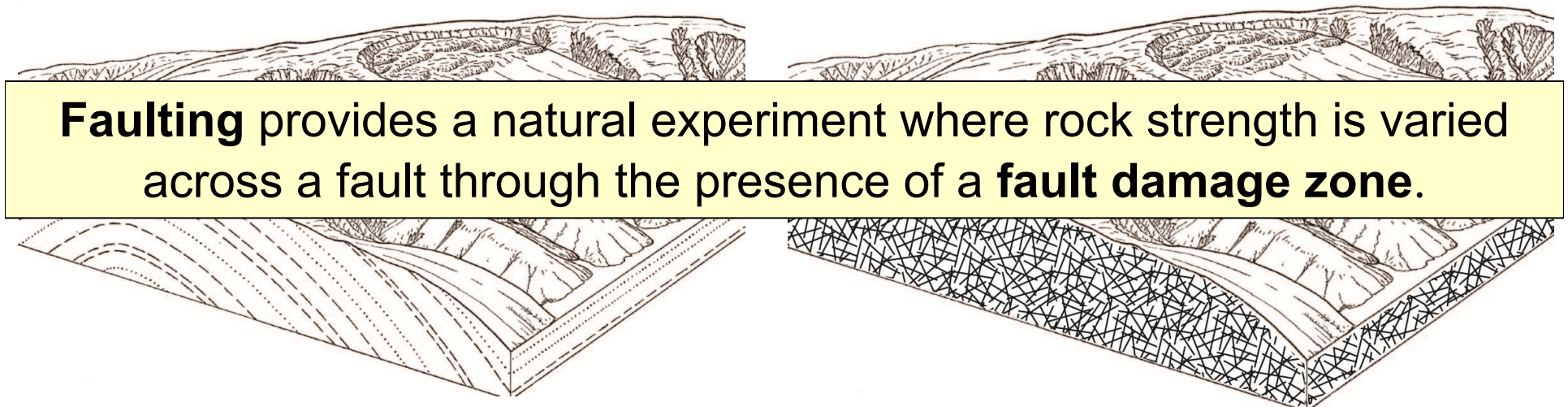
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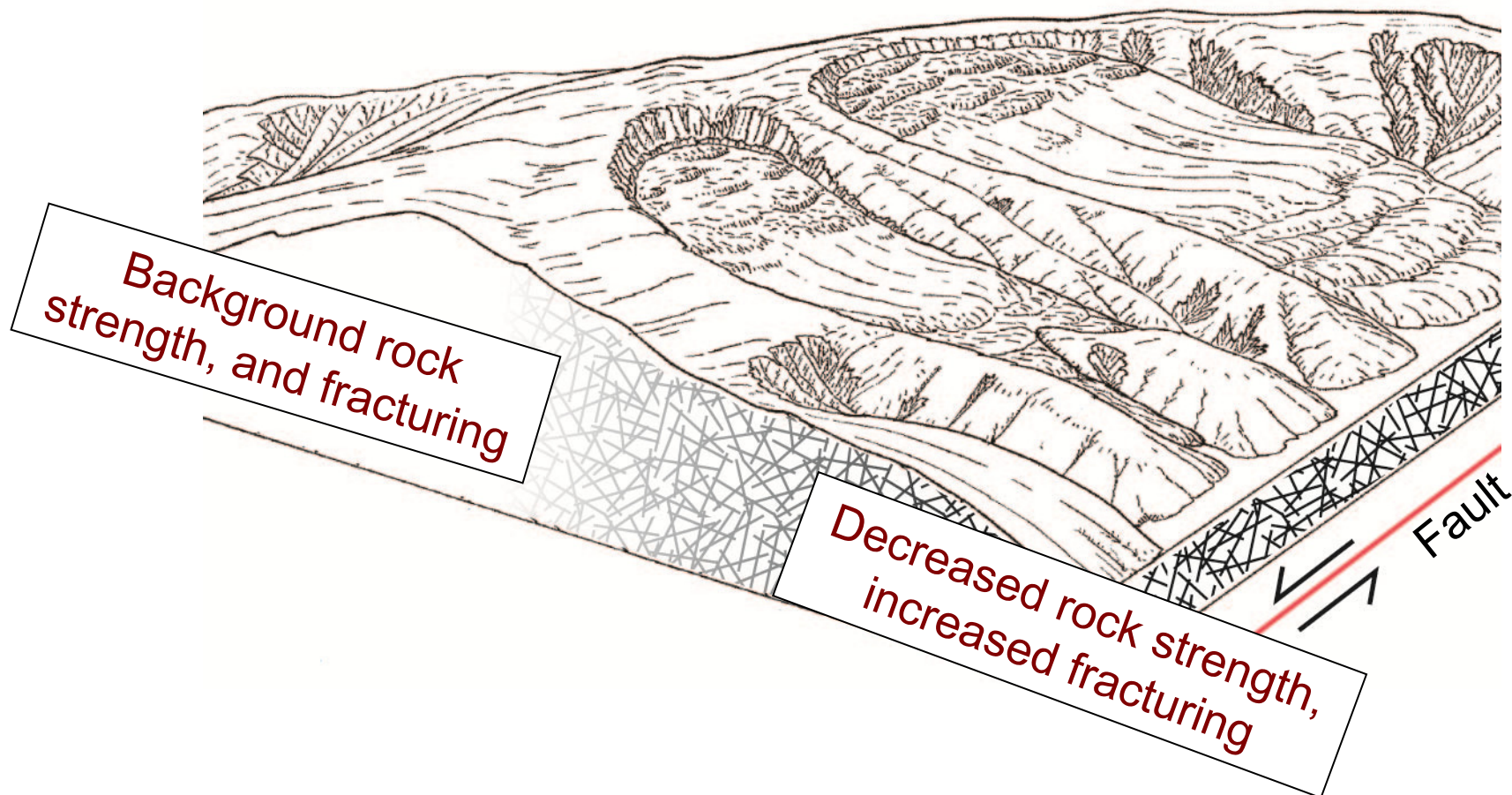


Fault damage zone (proxy for rock strength)

Fault damage zone: An area of increased bedrock fracturing and reduced rock strength extending meters to kilometers from the fault trace (e.g., Chester and Logan, 1986; Fialko et al., 2002, Ben-Zion and Sammis, 2003; Dor et al., 2006; Savage and Brodsky, 2011).

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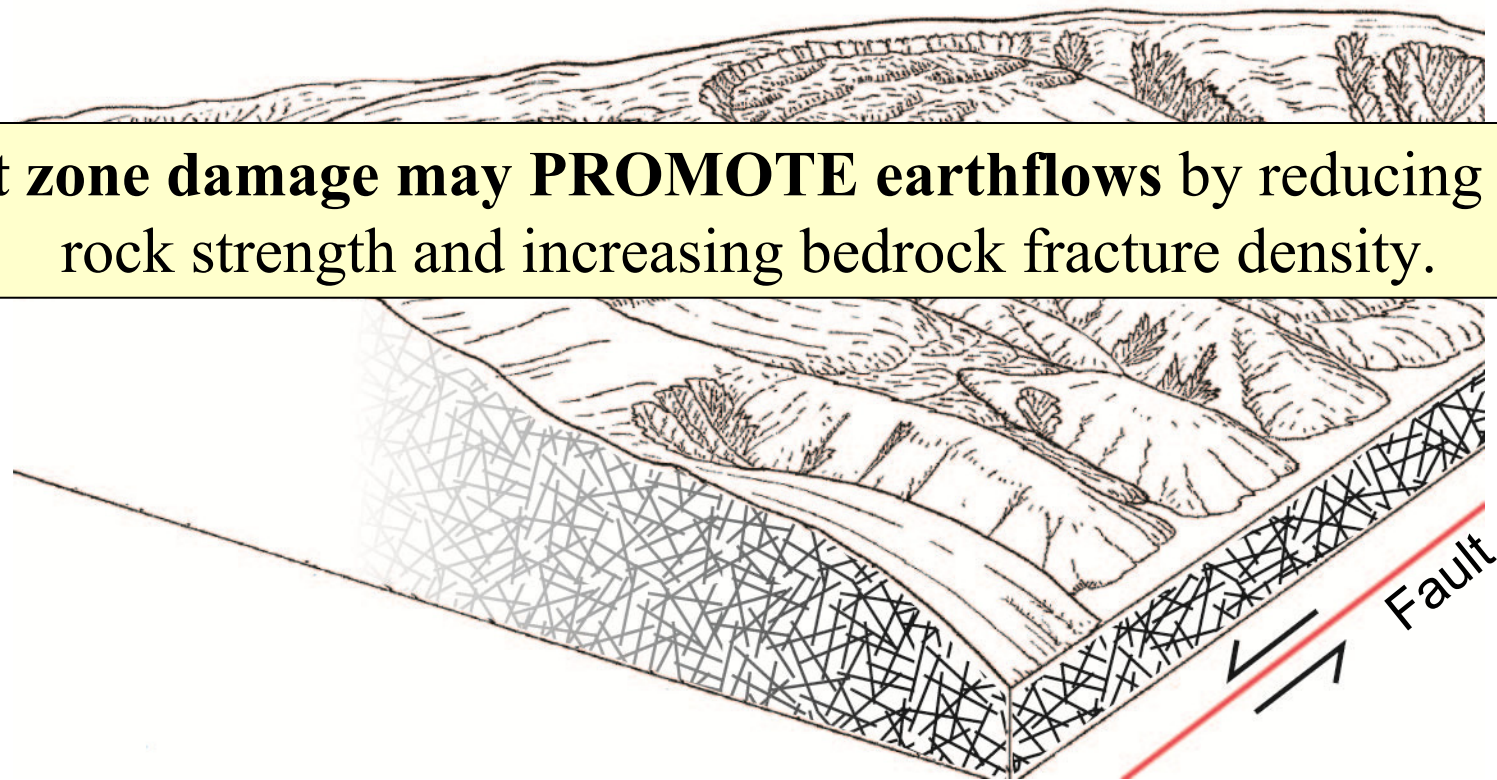


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Fault zone damage may PROMOTE earthflows by reducing bulk-rock strength and increasing bedrock fracture density.



Earthquakes induce co-seismic landslides

Sumatra, 2009, $M_w = 7.9$



AP Photo / Dita Alangkara



Adek Berry / AFP / Getty Image

Earthquakes induce co-seismic landslides

Sumatra, 2009, $M_w = 7.9$



AP Photo / Dita Alangkara



Adek Berry / AFP / Getty Image

Large-magnitude earthquakes may SUPPRESS earthflows by inducing co-seismic landslides (e.g., Keefer 1984) which preferentially remove fractured and weathered rock from the fault damage zone (i.e., the system may become **supply limited).**

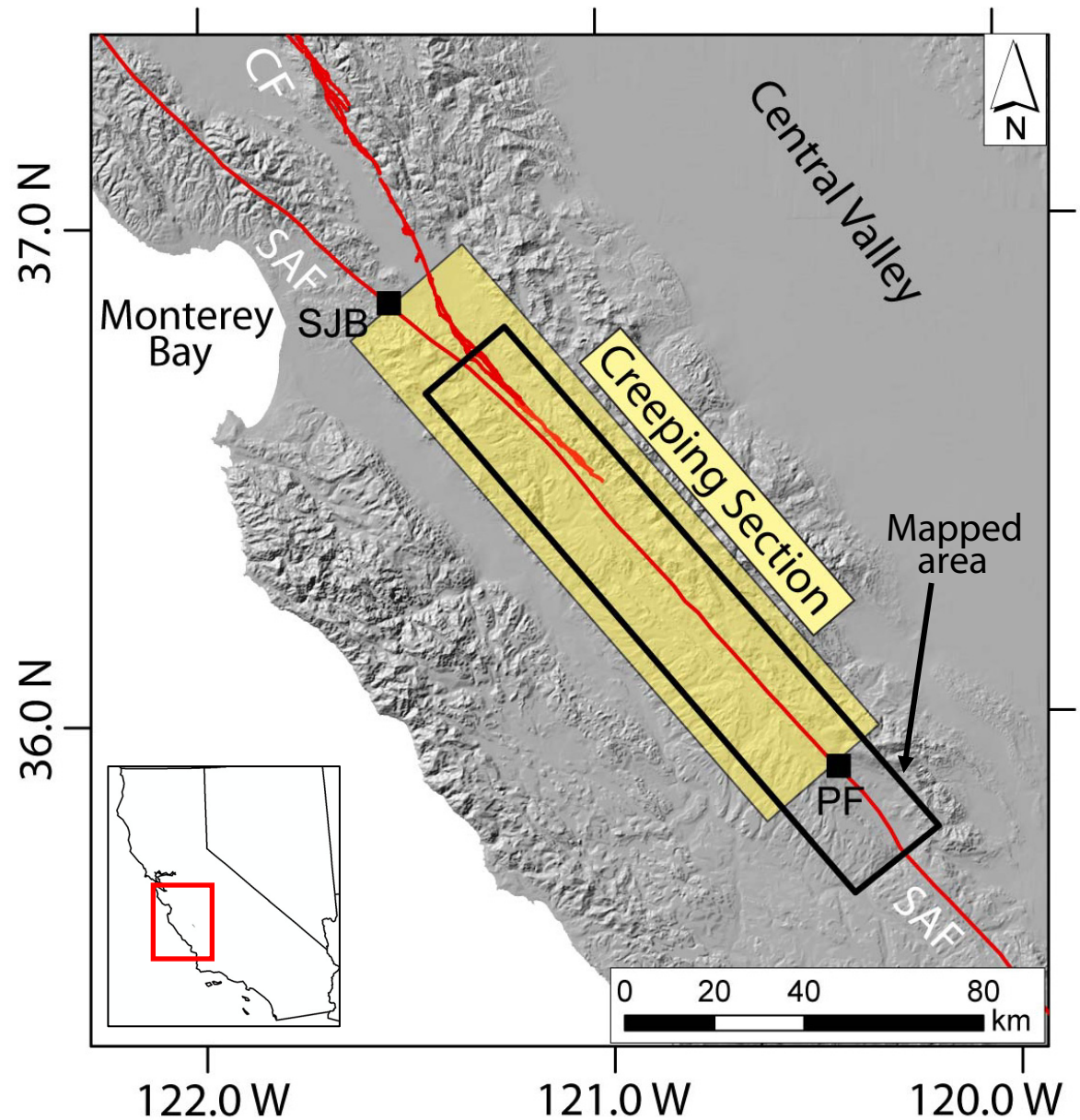
Outline

- **Study Site:** San Andreas Fault (SAF), California
- **Earthflow identification:** UAVSAR and aerial photos
- **Earthflow spatial distribution:** Central SAF
- **Central questions:**
 - Does reduced rock strength within fault damage zones promote earthflows?
 - Do large-magnitude earthquakes suppress earthflow development?

Central San Andreas Fault

Creeping section:

- Creep ~ 3 cm/yr
- Absence of large-magnitude earthquakes



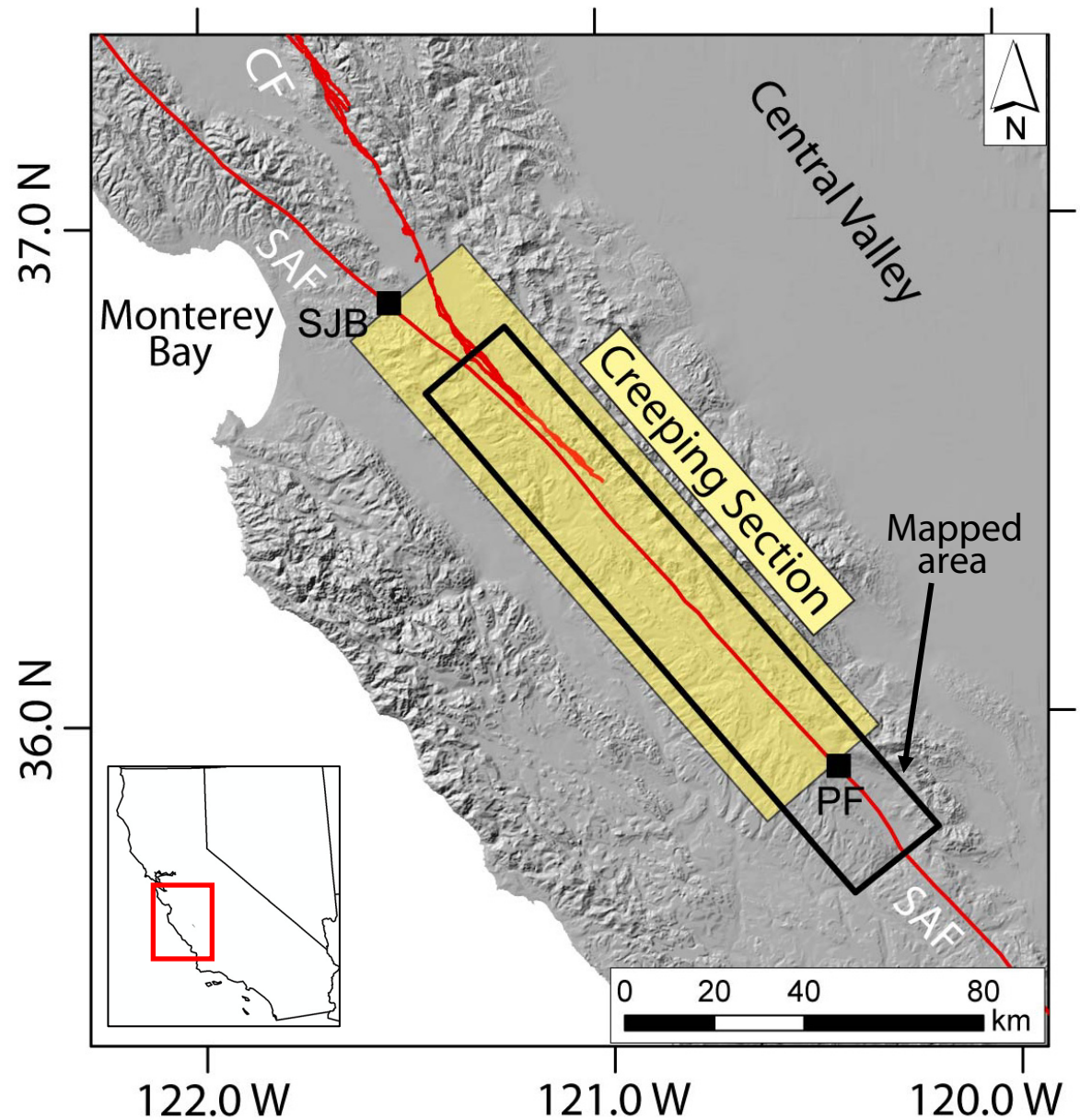
Central San Andreas Fault

Creeping section:

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Locked section:

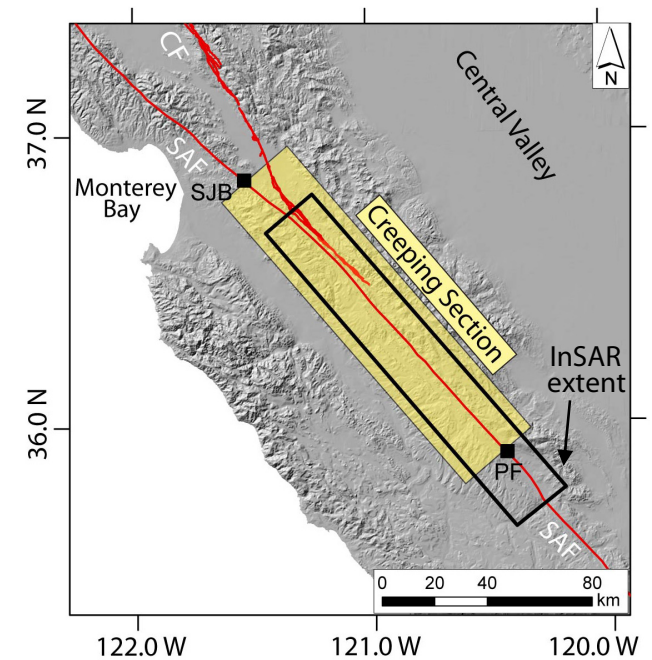
- Presence of large-magnitude earthquakes



Earthflow identification

1. Motion revealed by UAVSAR

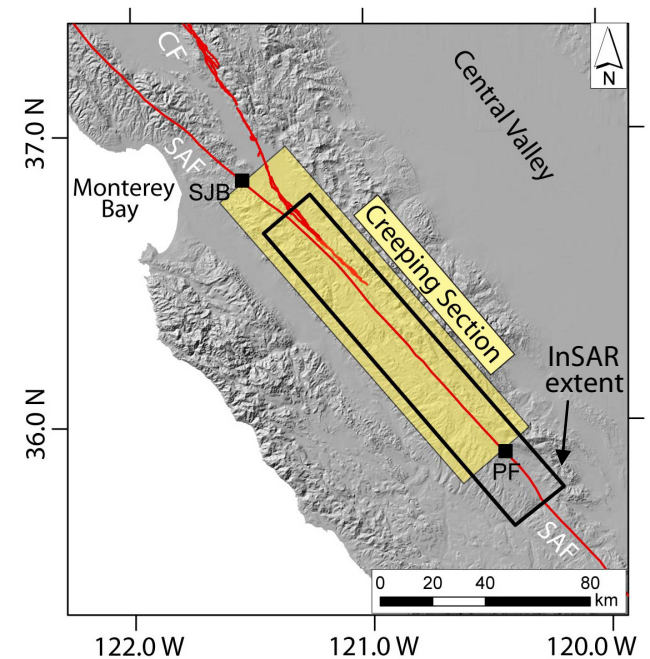
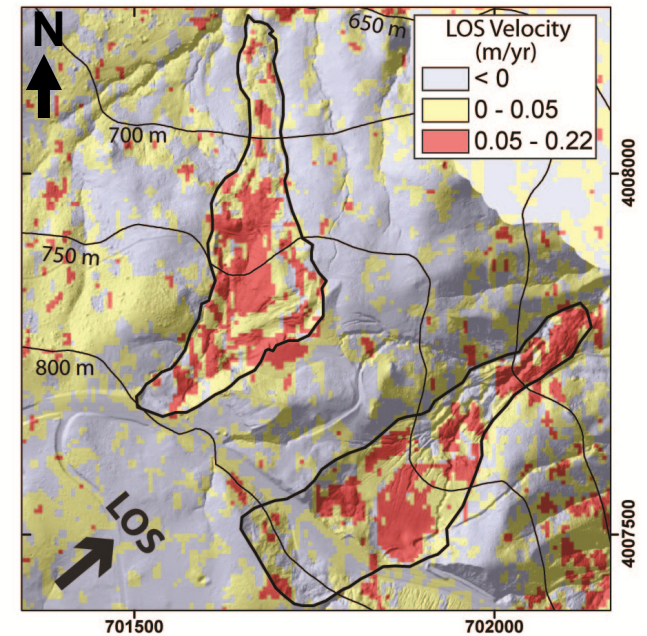
- 4 fault parallel interferograms
- 30 fault perpendicular interferograms
- Custom SNAPHU unwrapping



Earthflow identification

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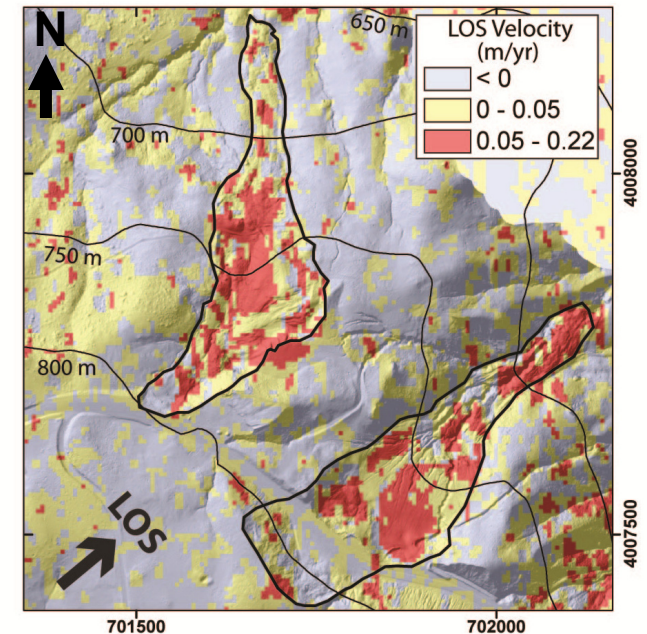
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Earthflow identification

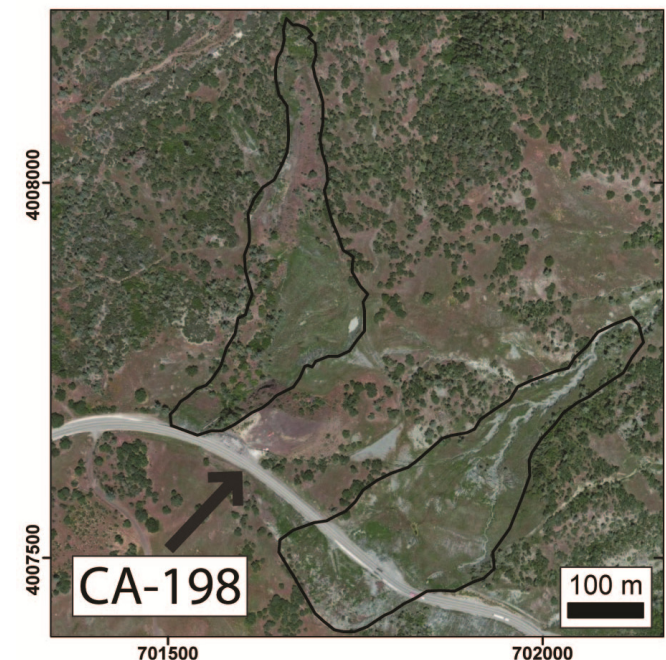
1. Motion revealed by UAVSAR

- 4 fault parallel interferograms
- 30 fault perpendicular interferograms
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2. Morphometric features in aerial photographs

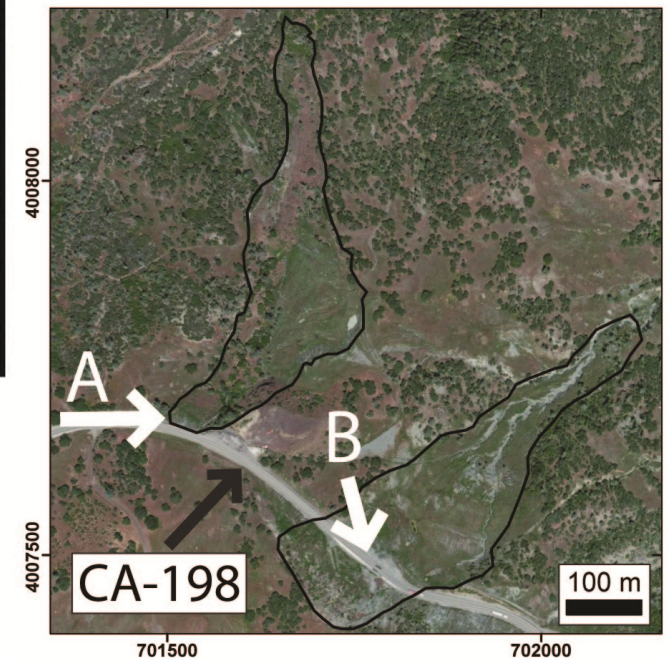
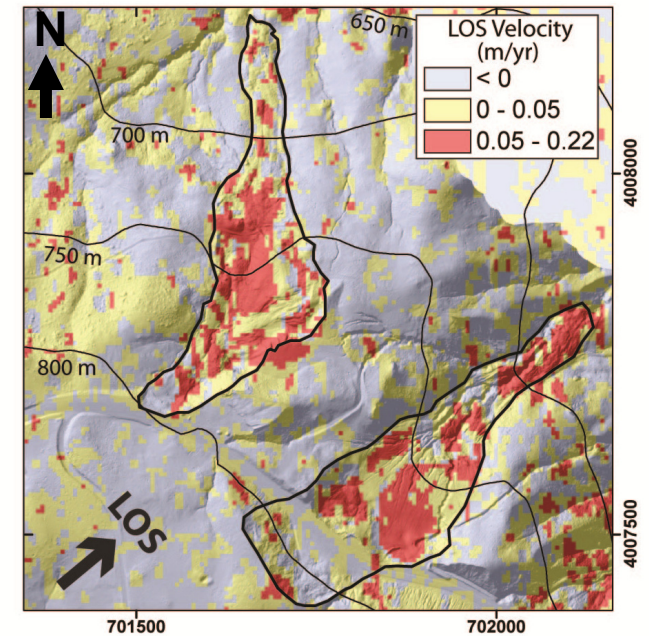
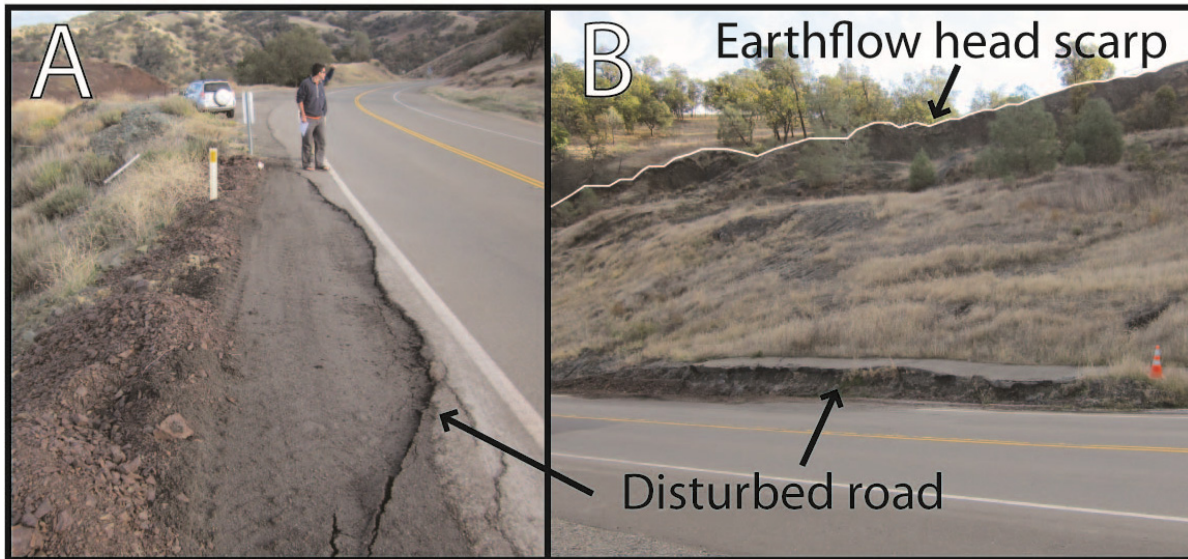
- Lateral margins, pressure ridges, and hummocky terrain (McKean and Roering, 2004)
- ~1 m² resolution aerial photographs (BING Maps)
- Accurately ortho-rectified within ESRI ArcMap



Earthflow identification

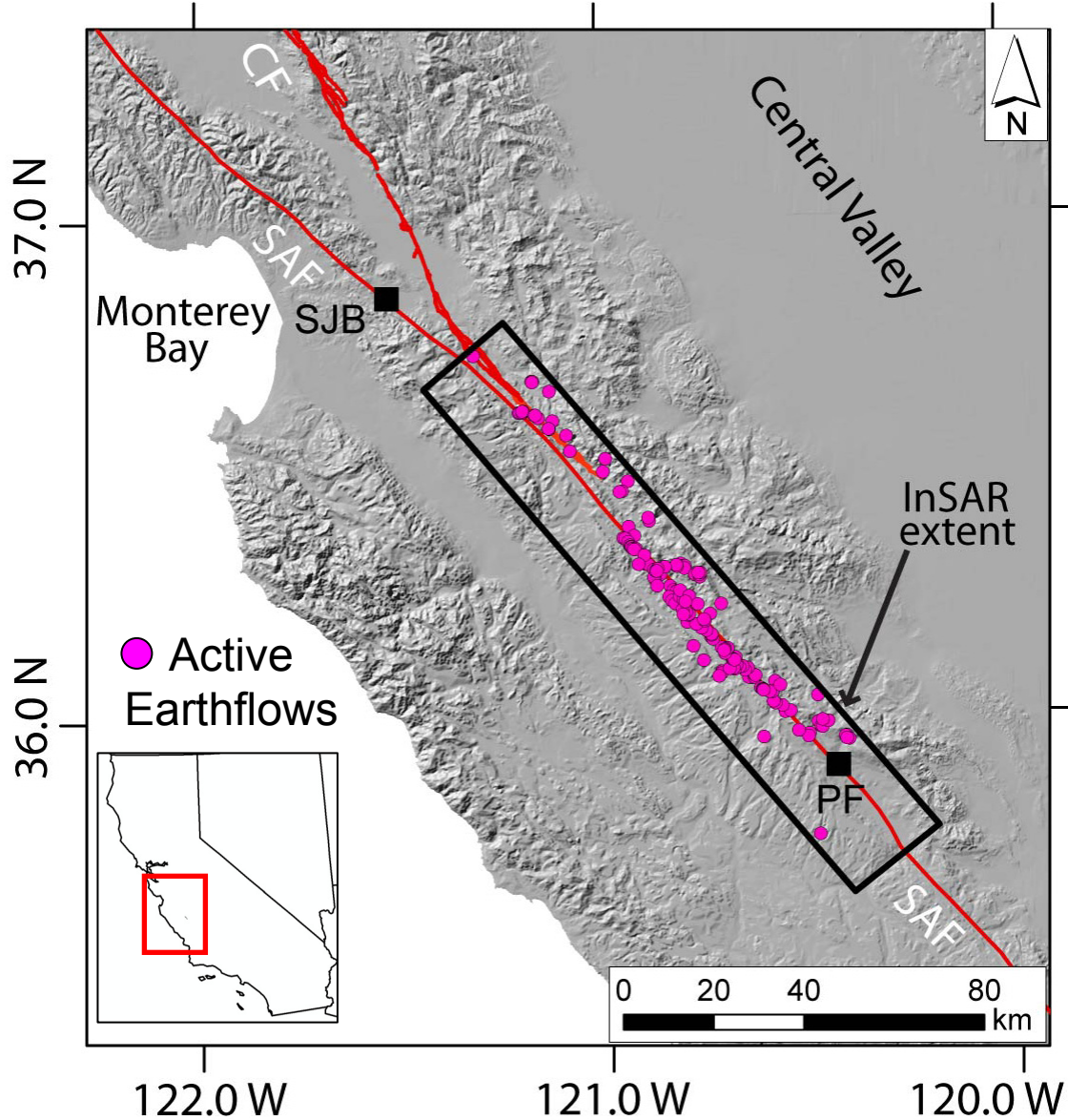
3. Field Visits

- Observations of deformed roads and active highway maintenance



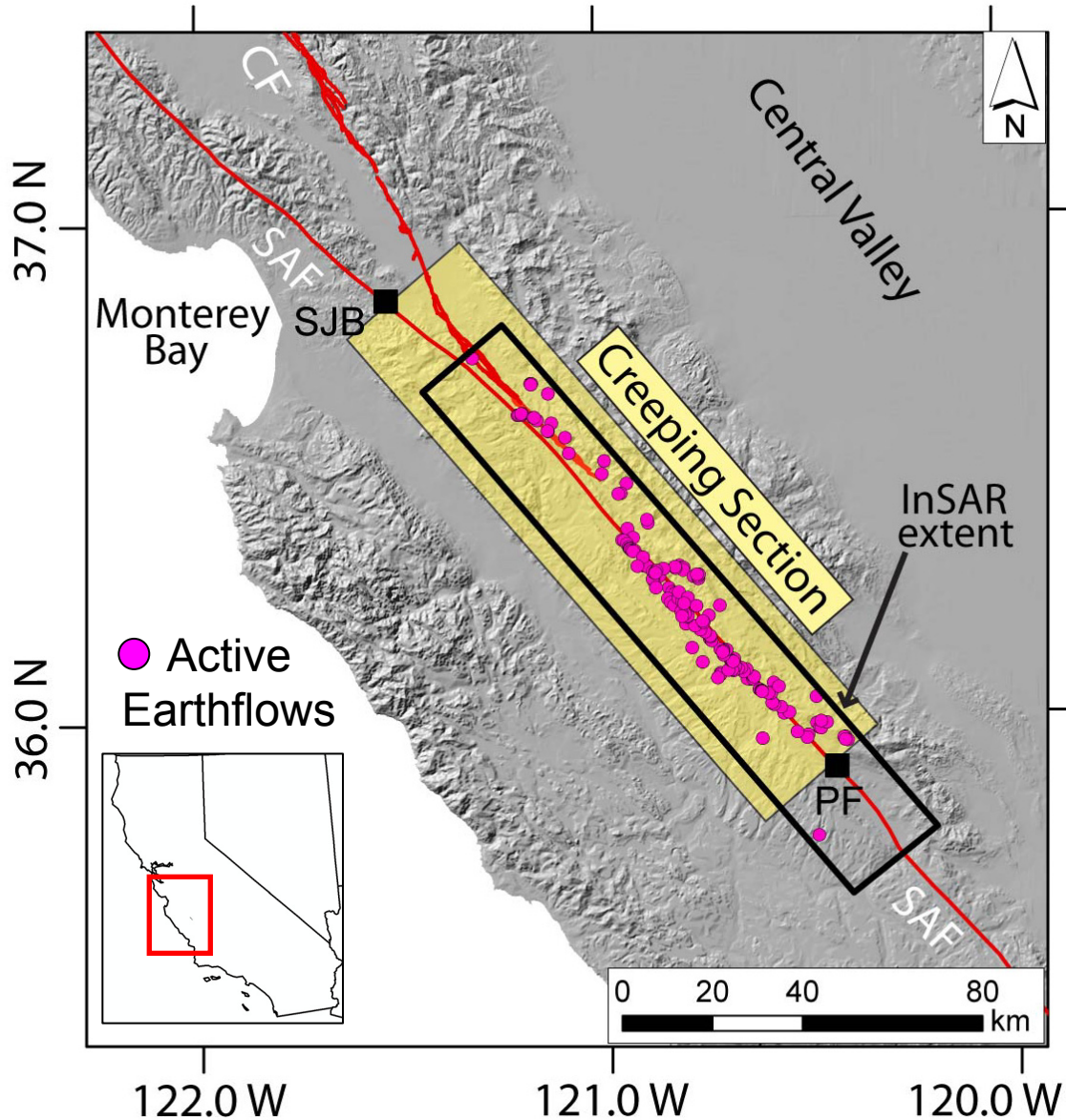
Earthflow spatial distribution

150 active earthflows identified May 2010 to July 2011



Earthflow spatial distribution

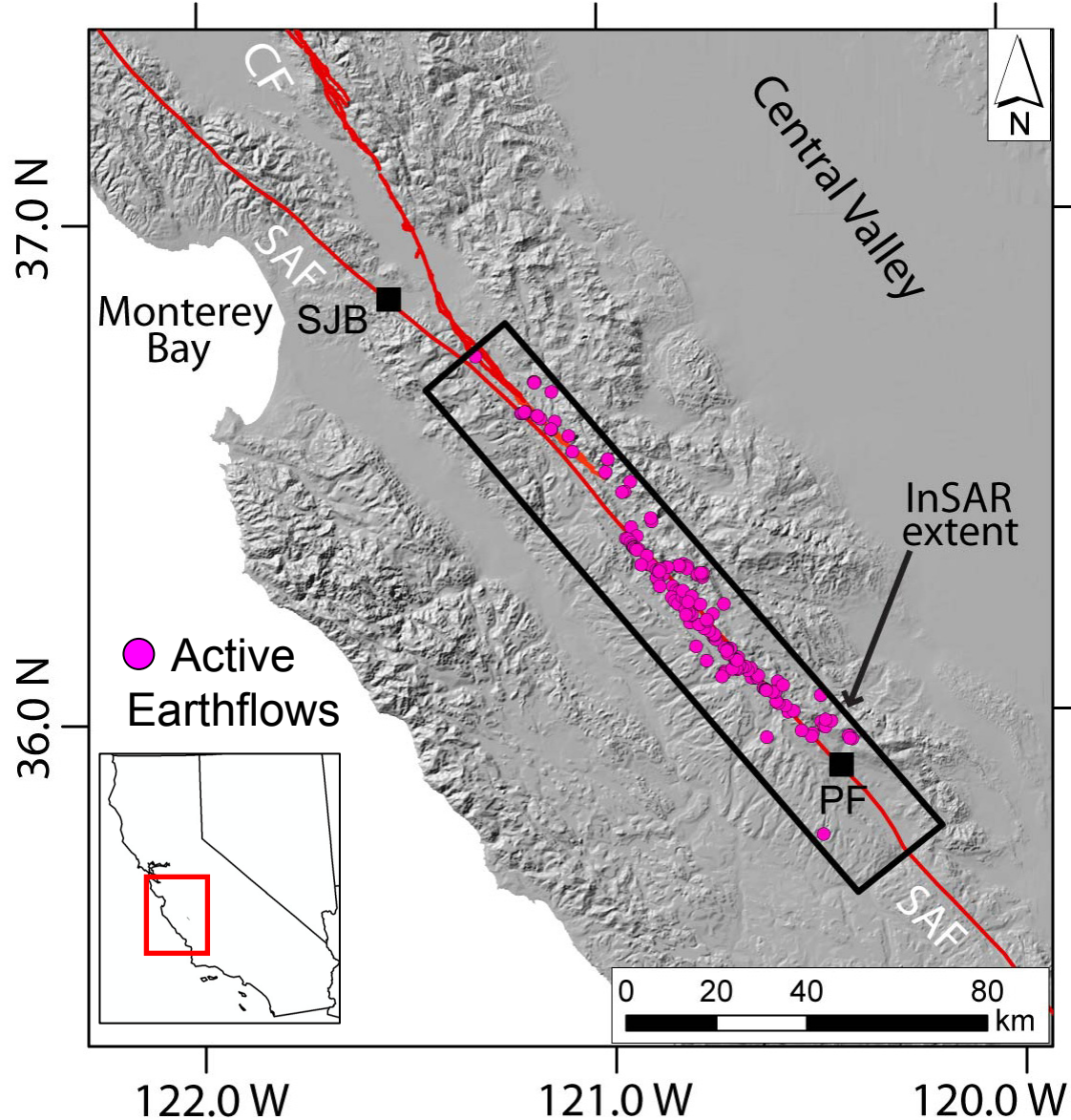
150 active earthflows identified May 2010 to July 2011



1. Almost all earthflows are within the creeping section of the SAF.

Earthflow spatial distribution

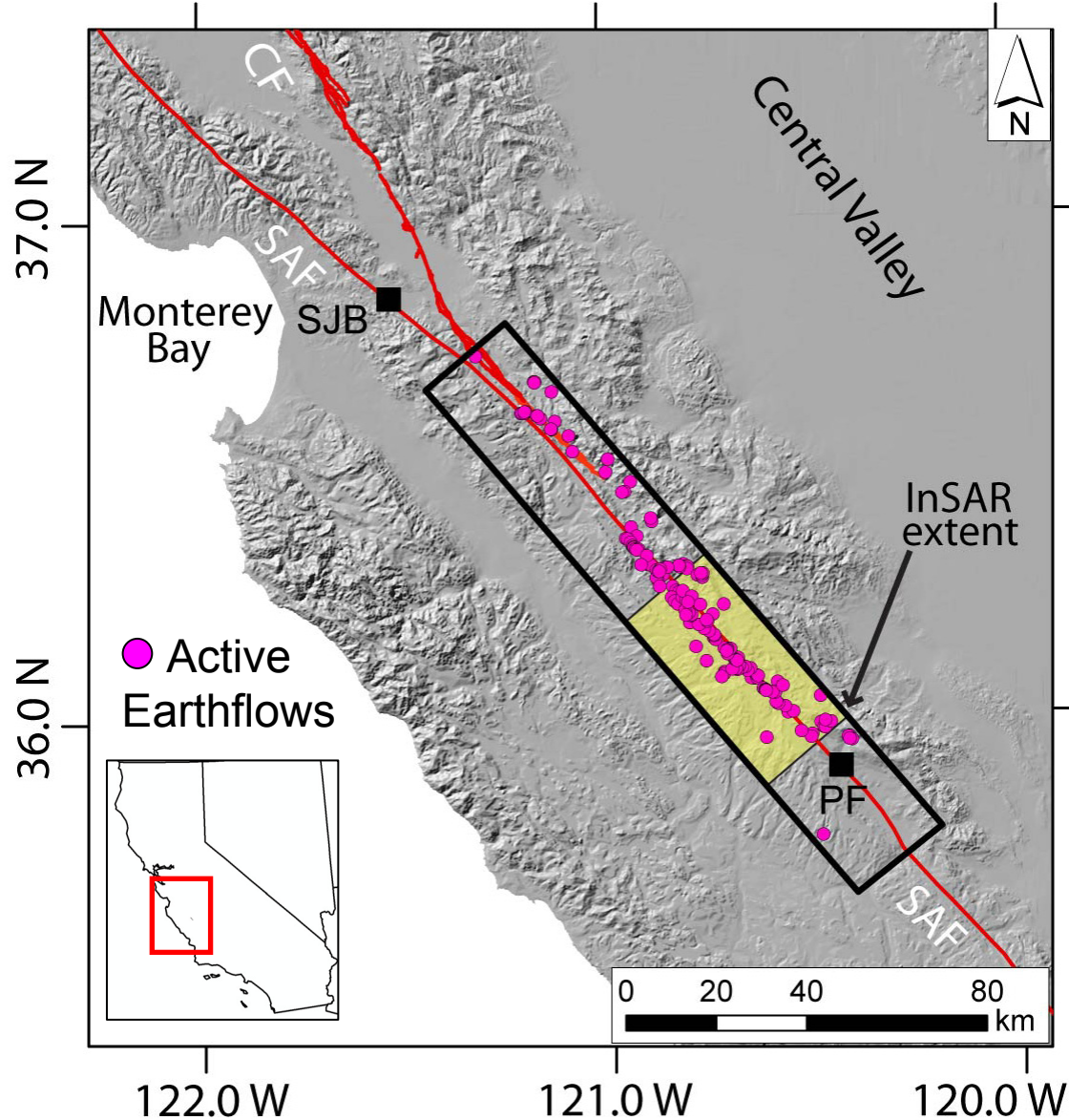
150 active earthflows identified May 2010 to July 2011



1. Almost all earthflows are within the creeping section of the SAF.
2. Earthflow cluster near the SAF.

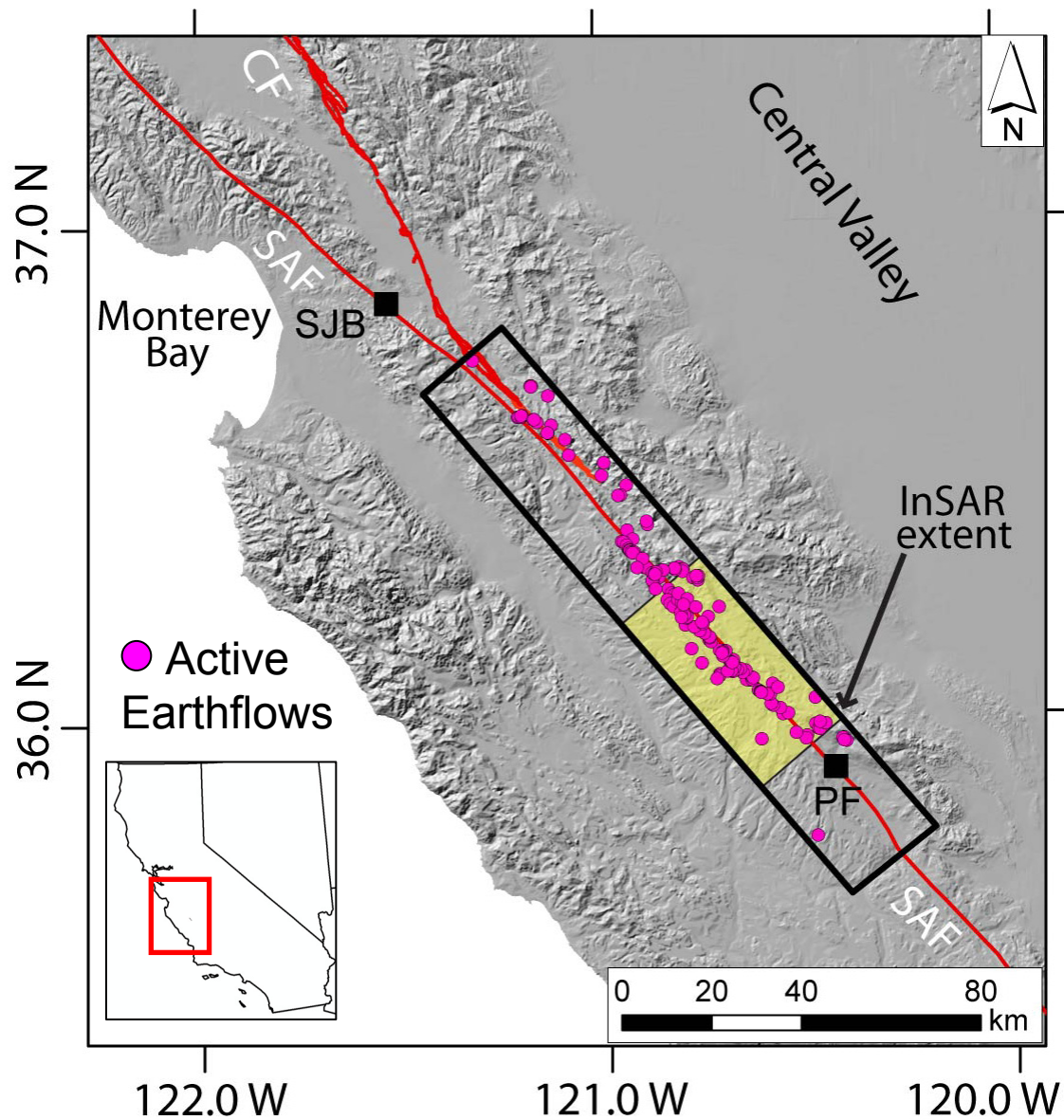
Earthflow spatial distribution

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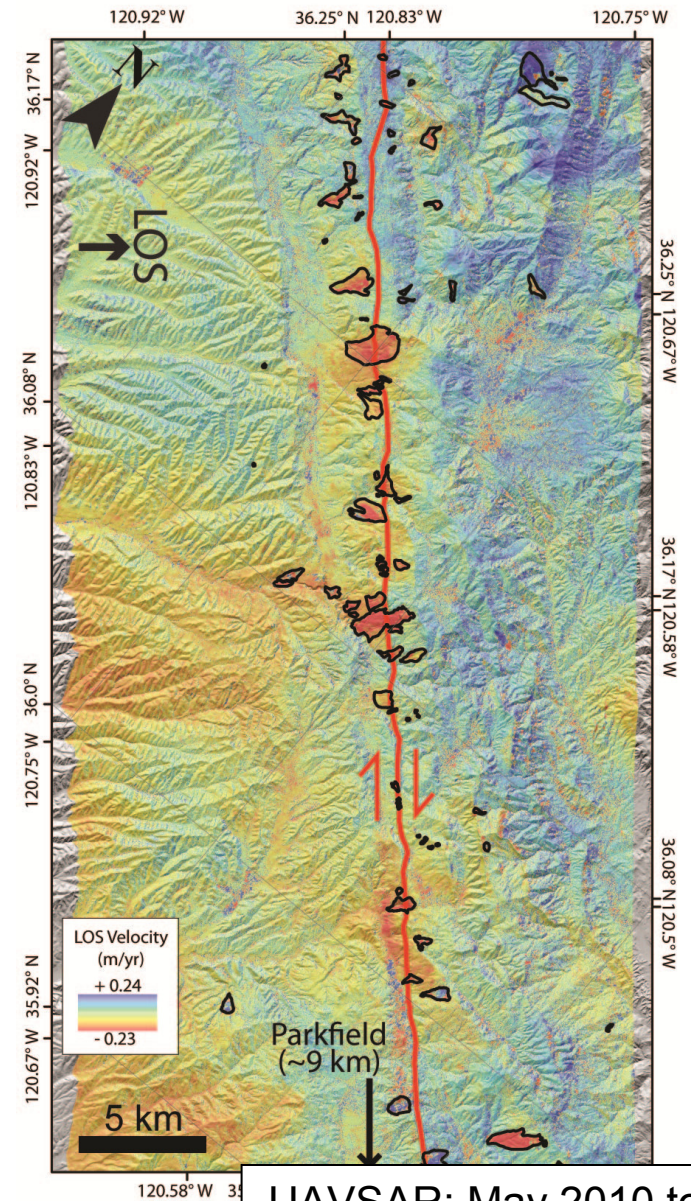


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Earthflow spatial distribution

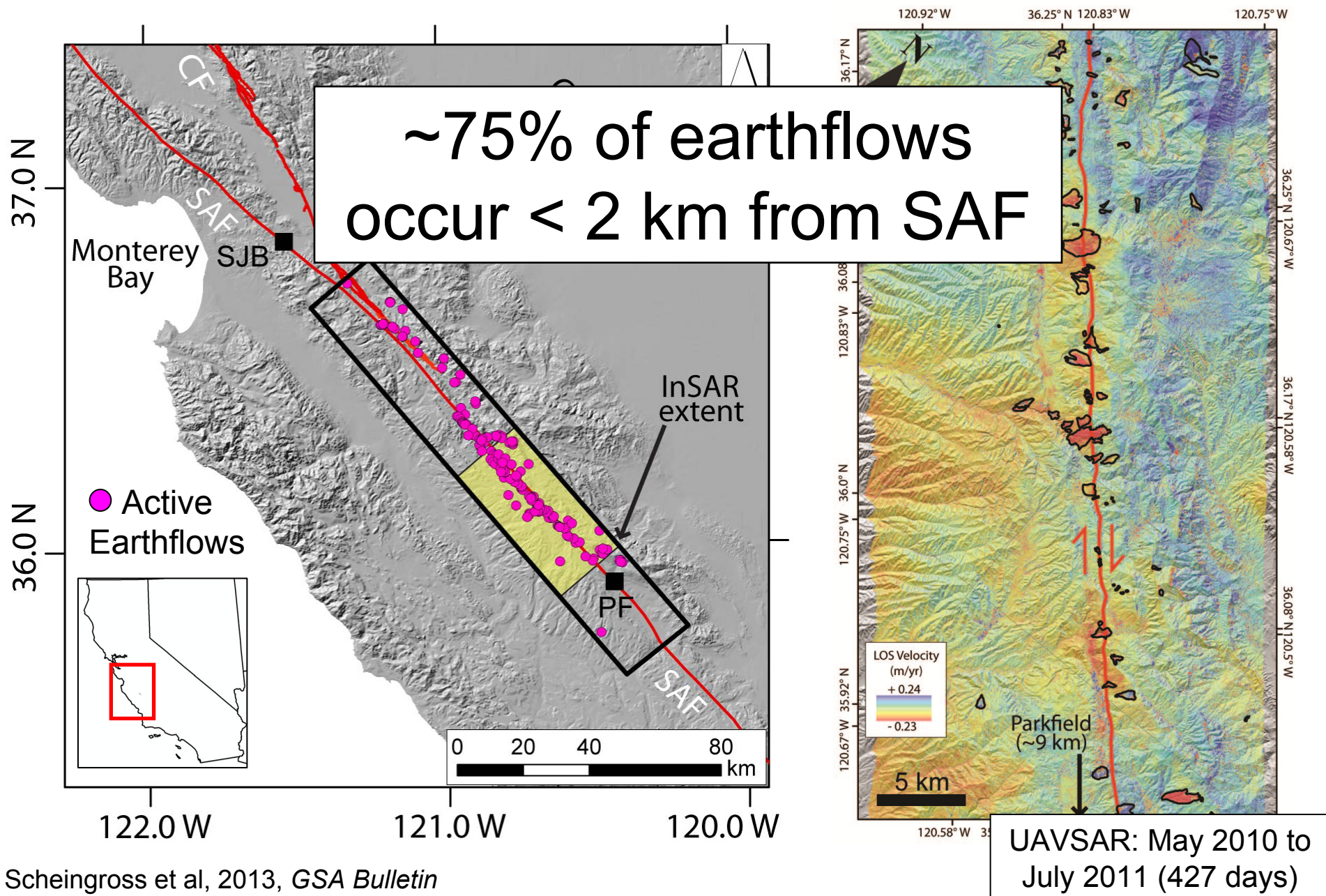


Scheingross et al, 2013, *GSA Bulletin*



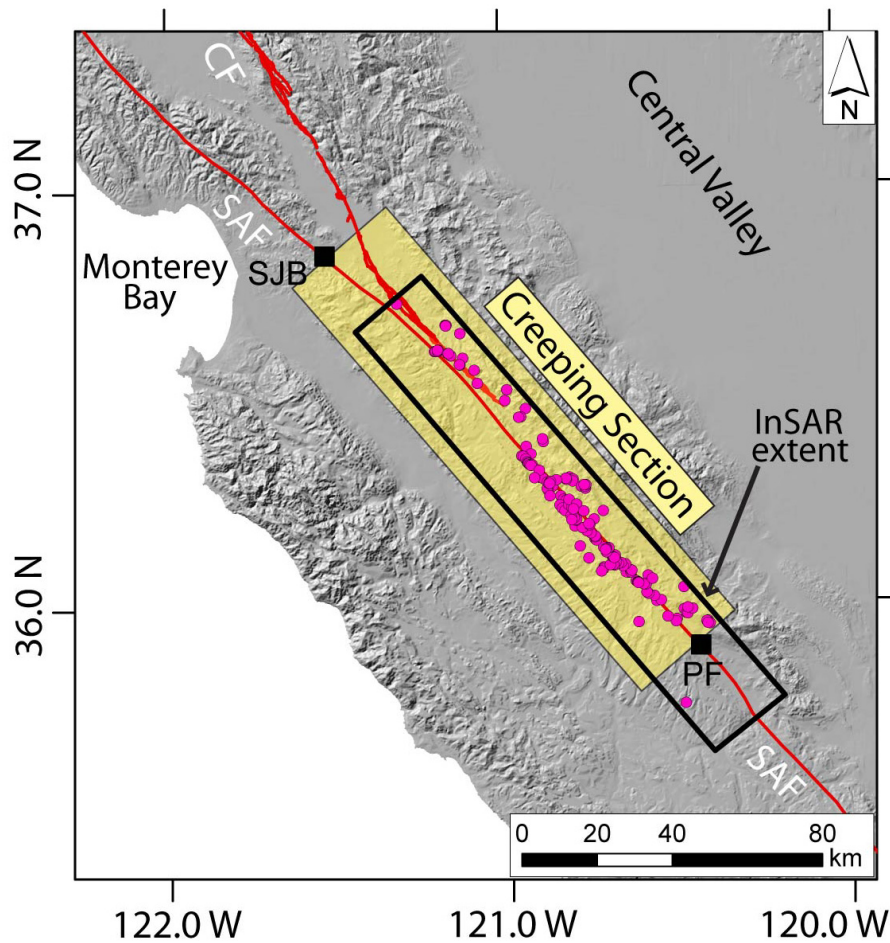
UAVSAR: May 2010 to
July 2011 (427 days)

Earthflow spatial distribution



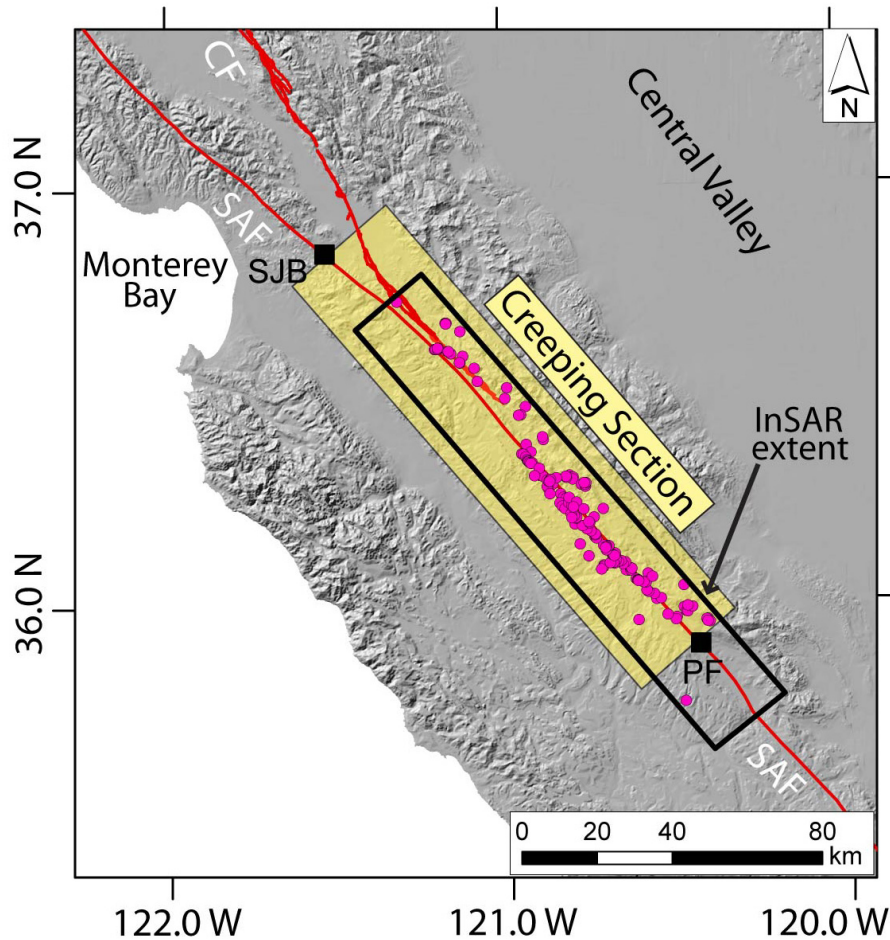
Central questions

- Does **reduced rock strength** within the fault damage zone influence earthflow spatial distribution?



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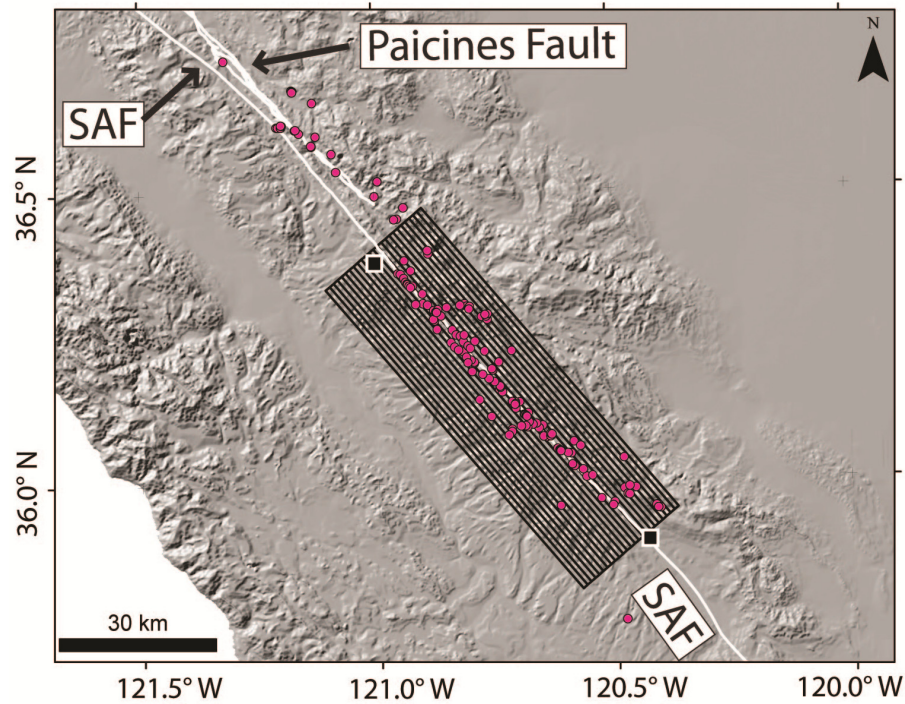


What about confounding variables?

- Topography
- Climate
- Lithology

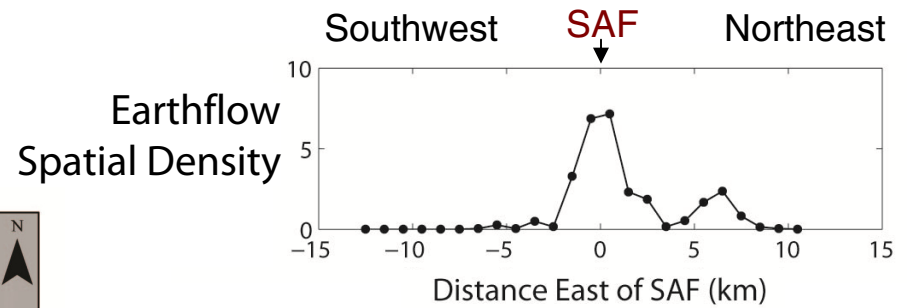
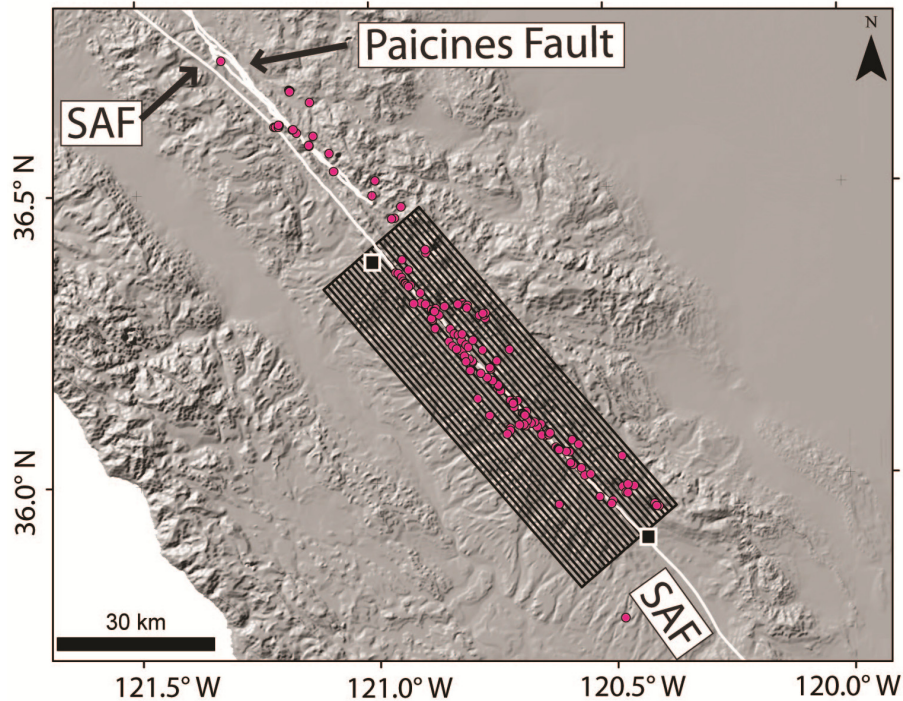
Swath Profile Analysis

SAF parallel swaths
(1 km x 75 km)



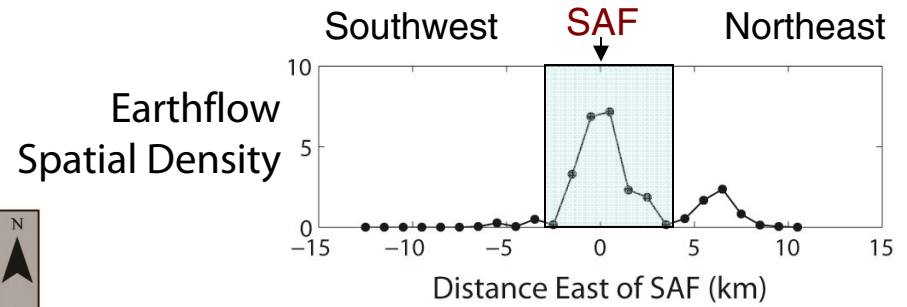
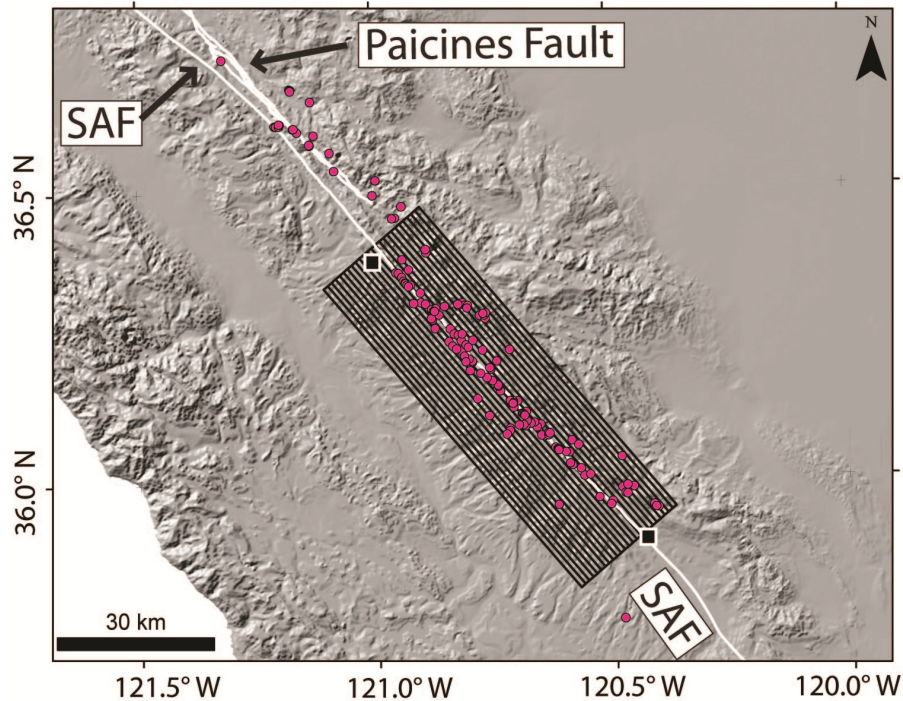
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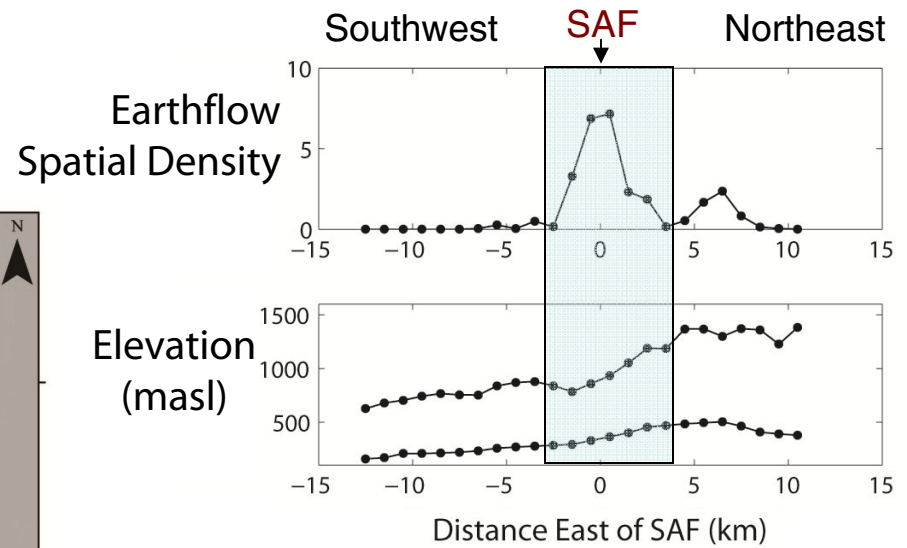
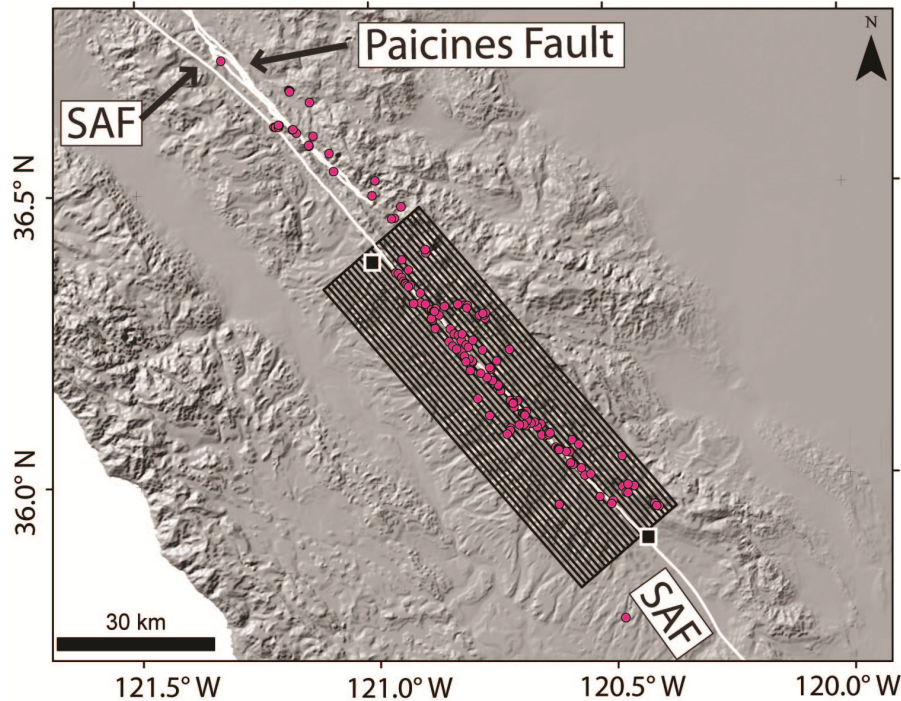
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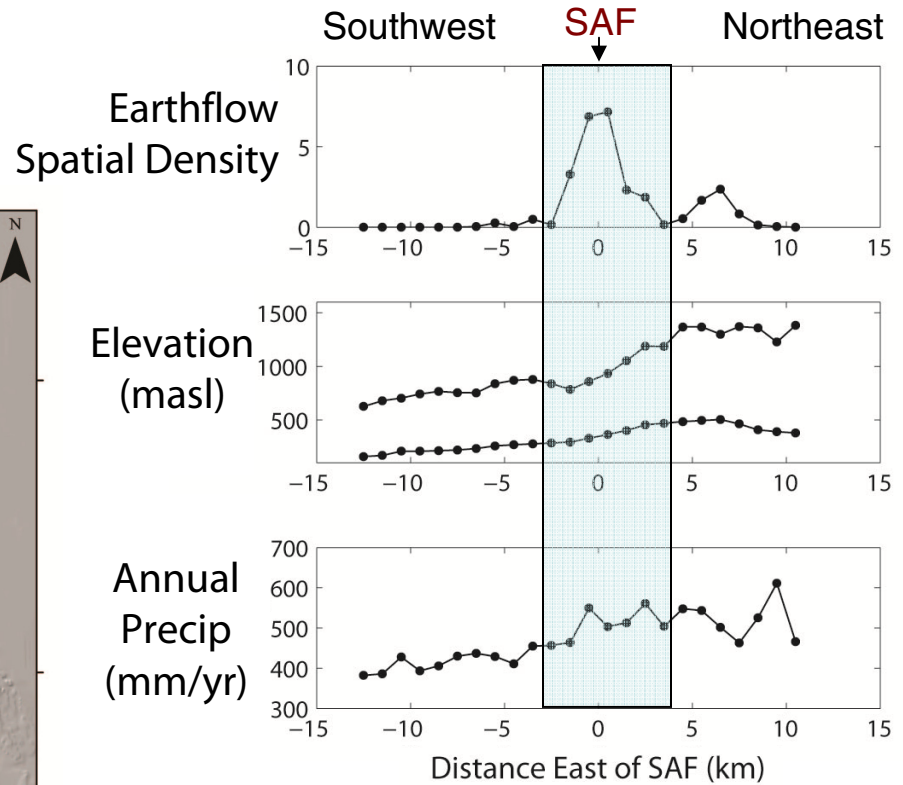
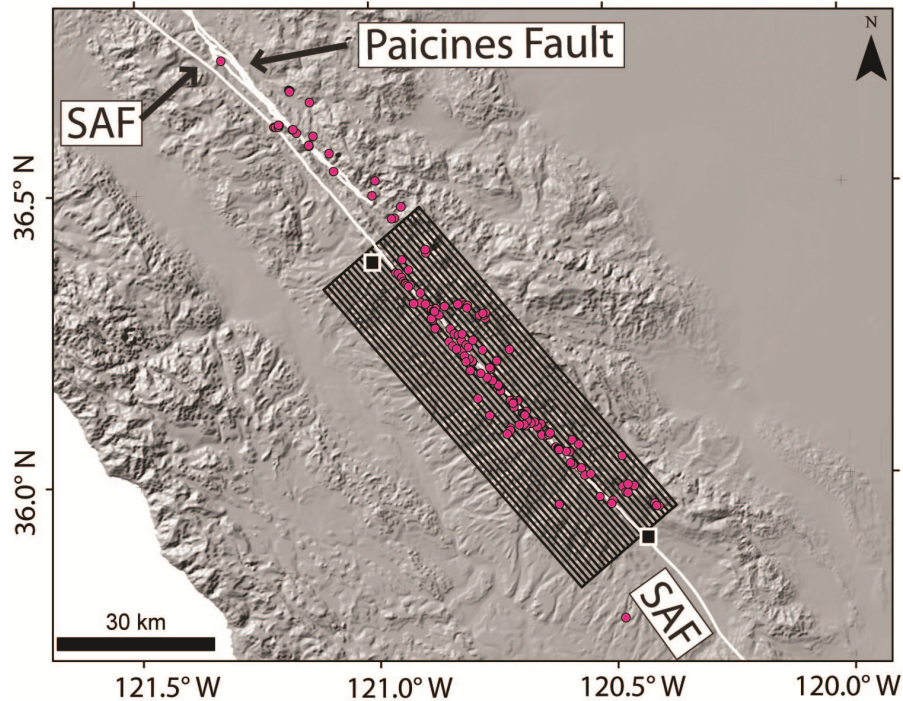
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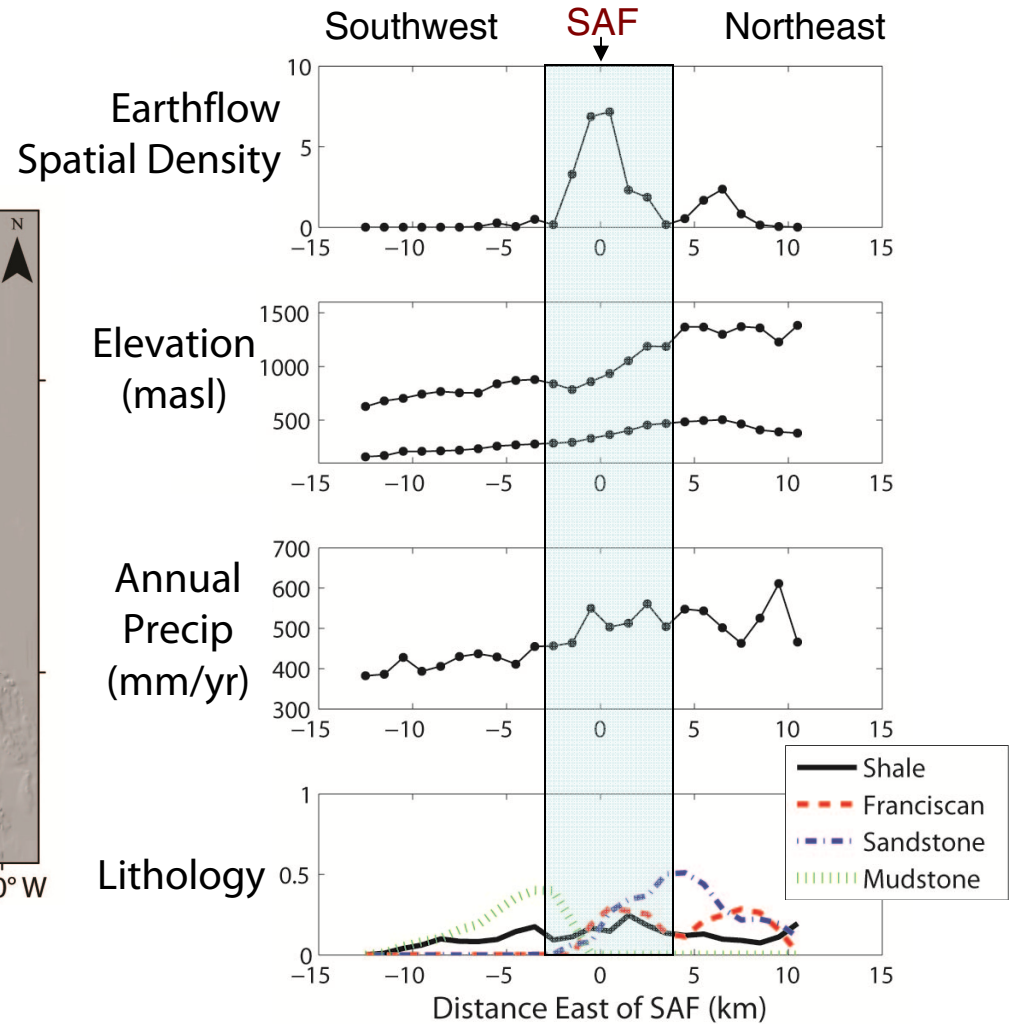
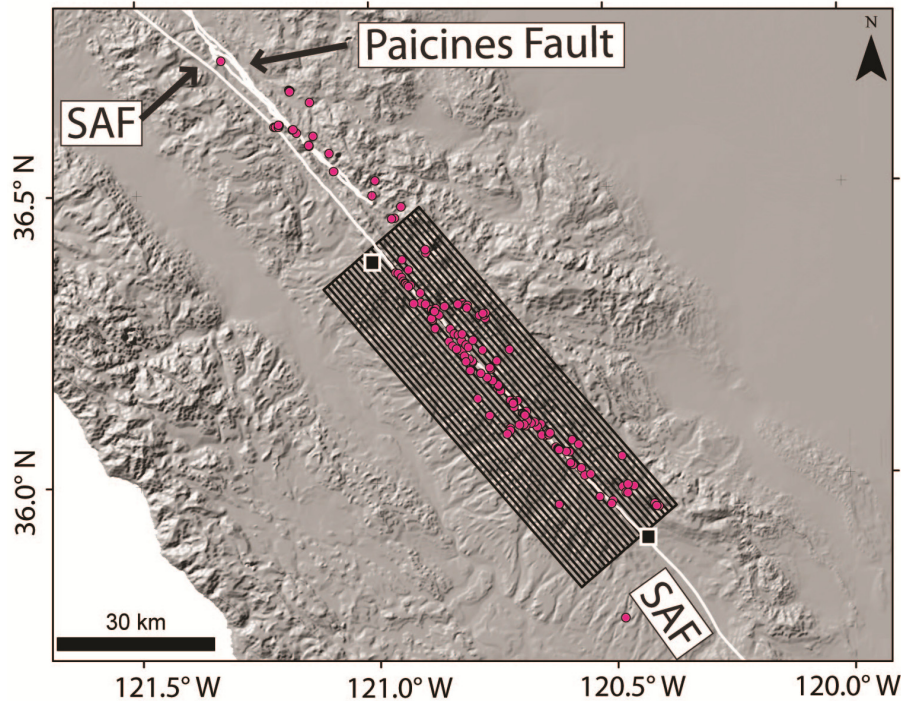
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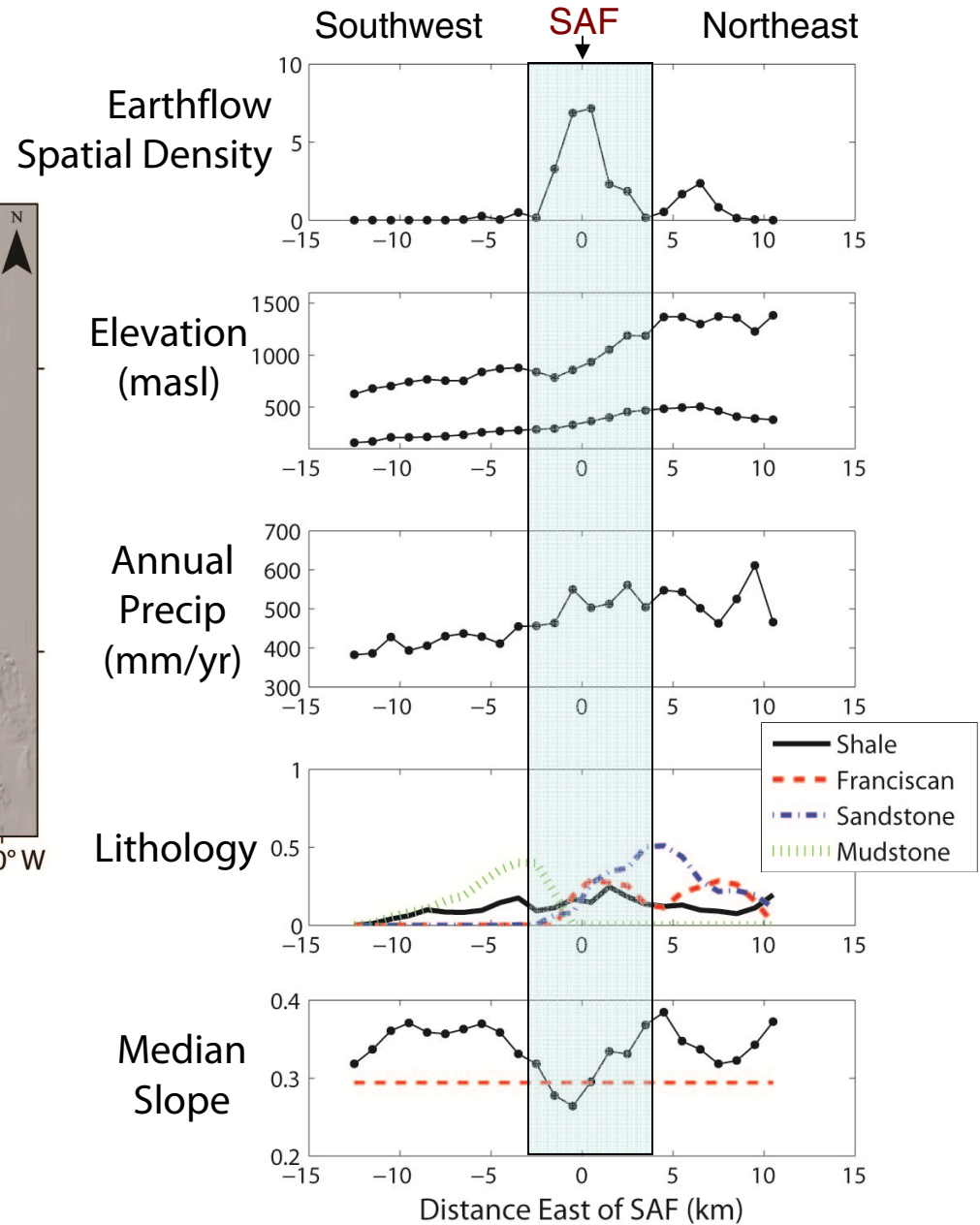
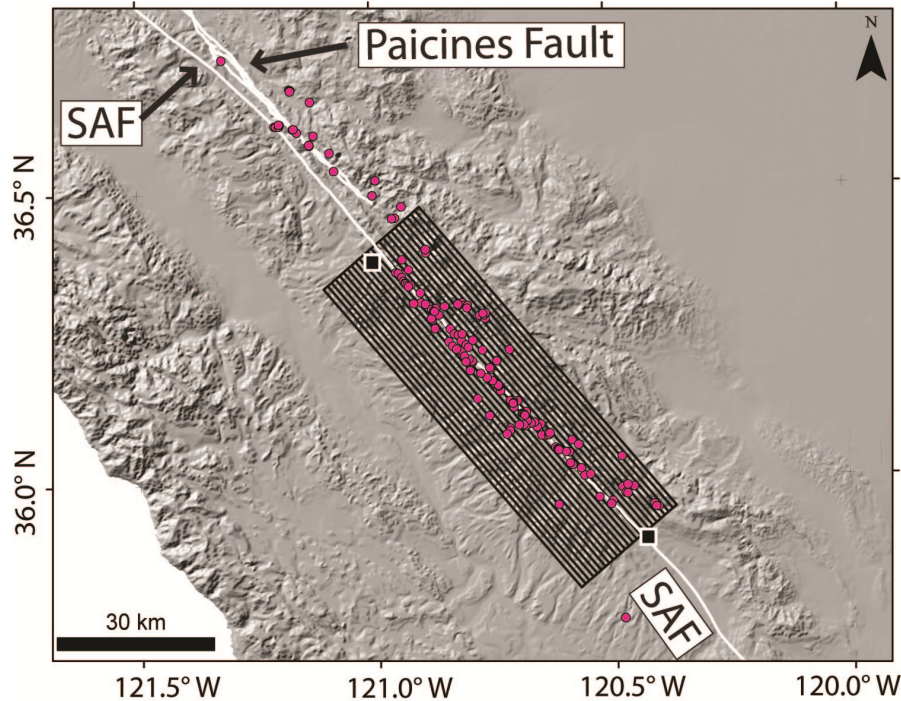
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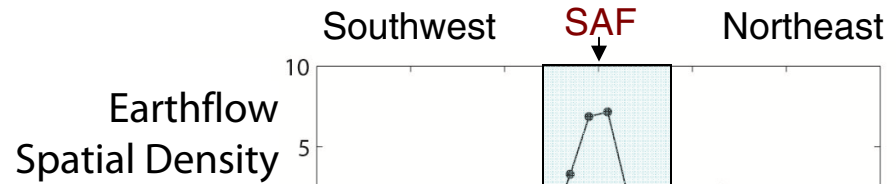
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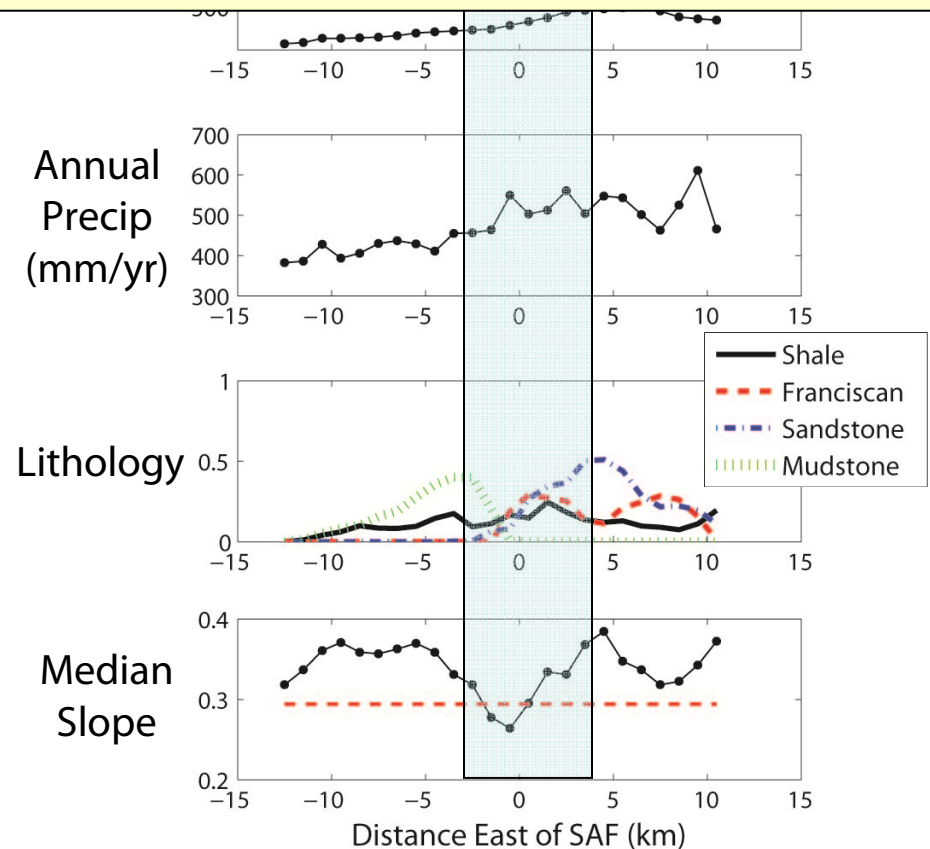
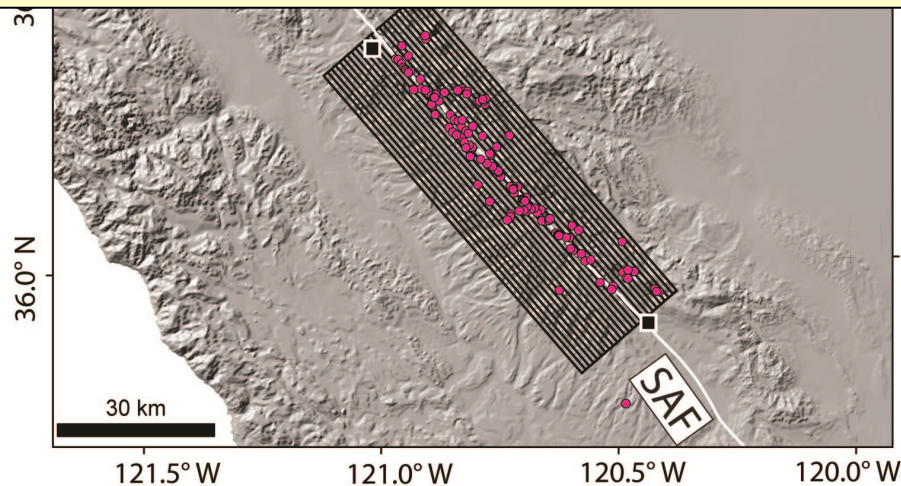


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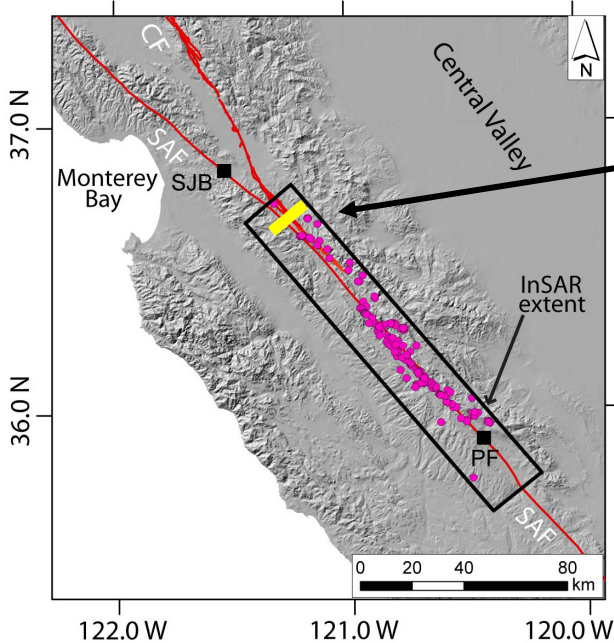


Topographic, precipitation, and rock type, metrics **alone** are not enough to explain the observed spatial distribution of earthflows.



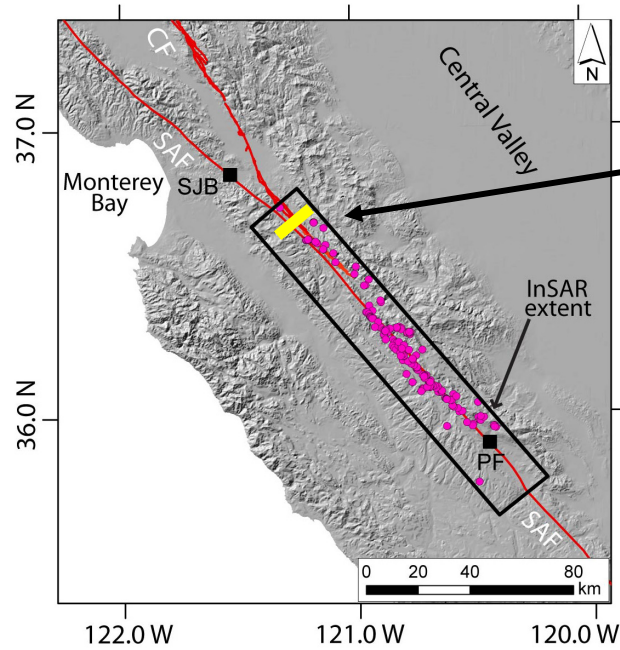
Fault damage zones

Fault damage zones



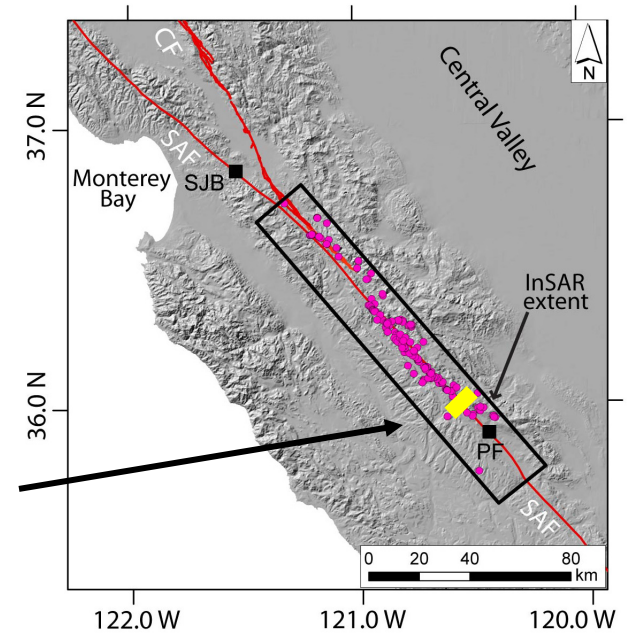
Seismic velocity profile
(Thurber et al, 1997)

Fault damage zones

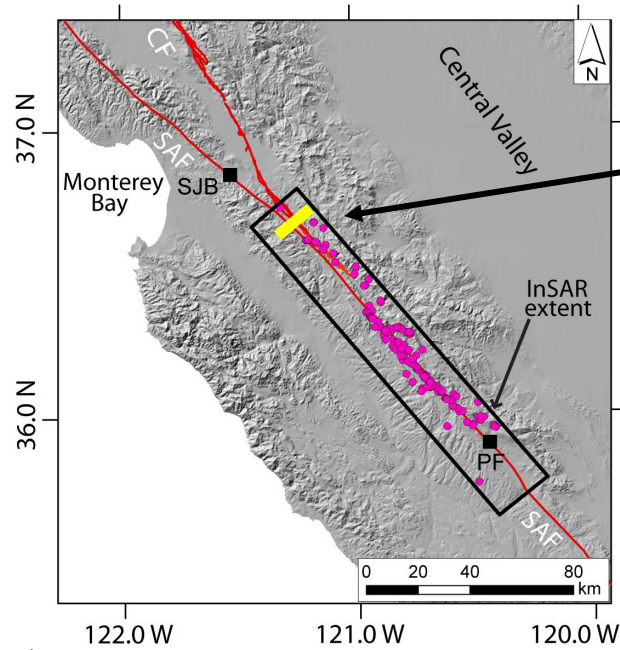


Seismic velocity profile
(Thurber et al, 1997)

Electrical resistivity profile
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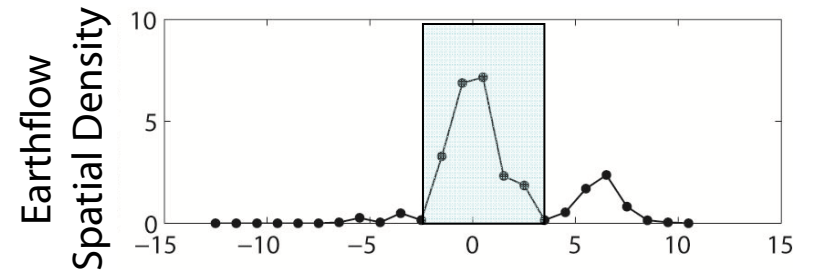
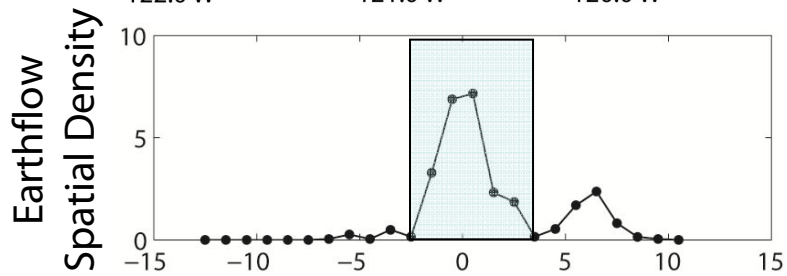
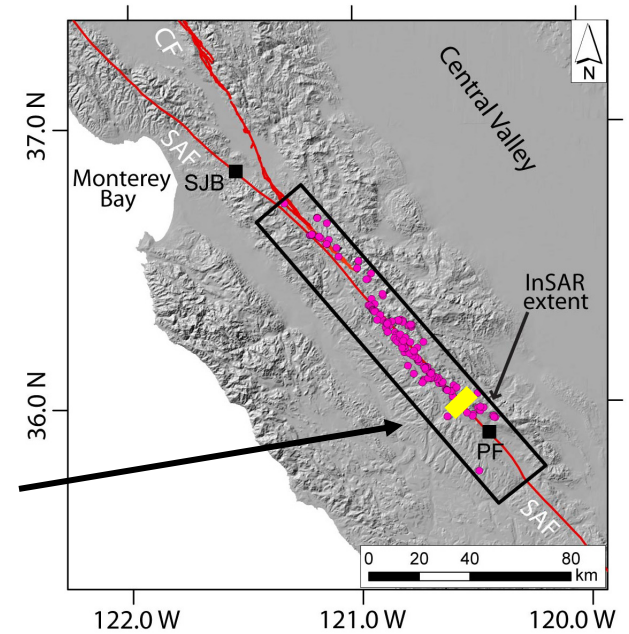


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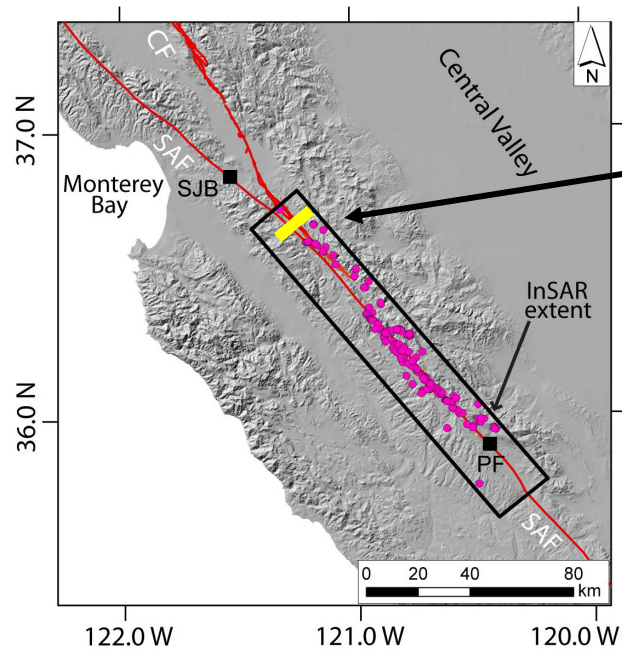


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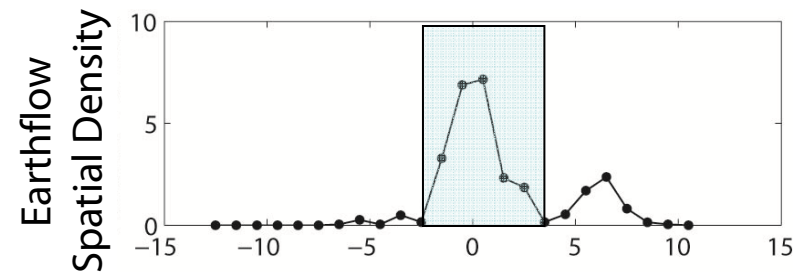
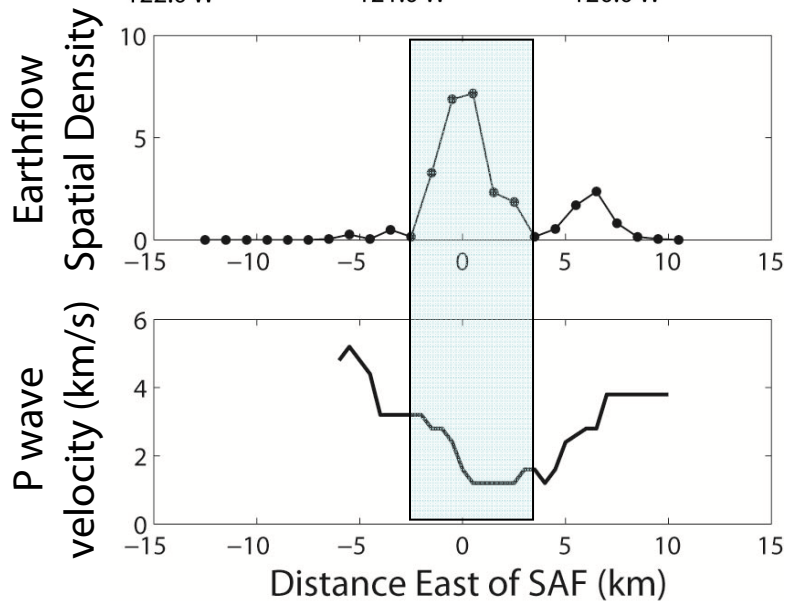
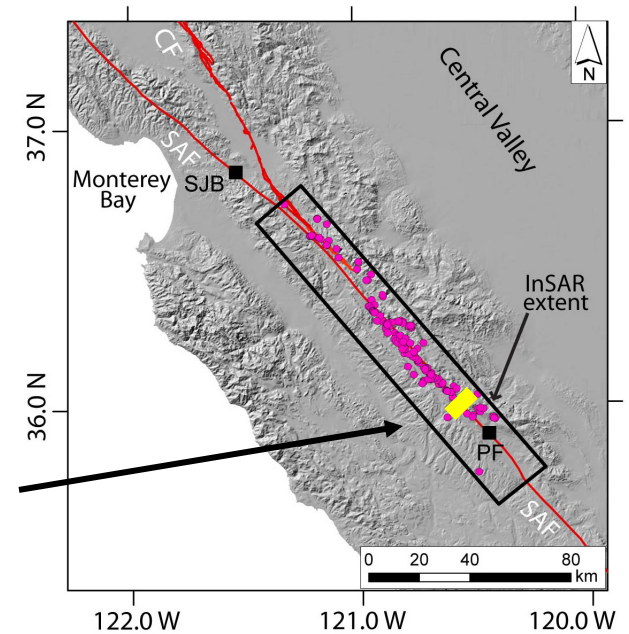


Fault damage zones



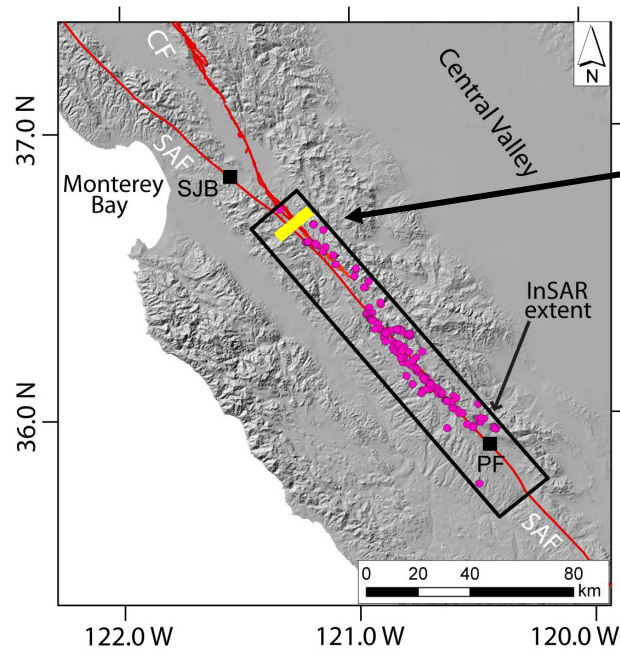
Seismic velocity profile
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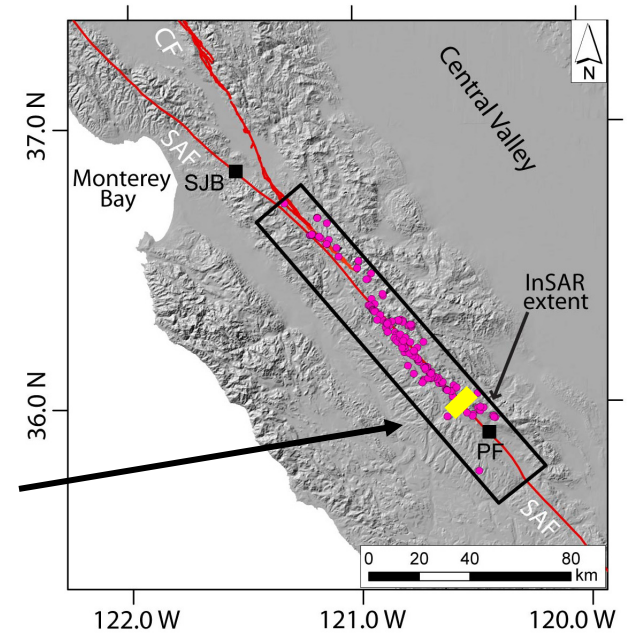


Decreasing
rock strength

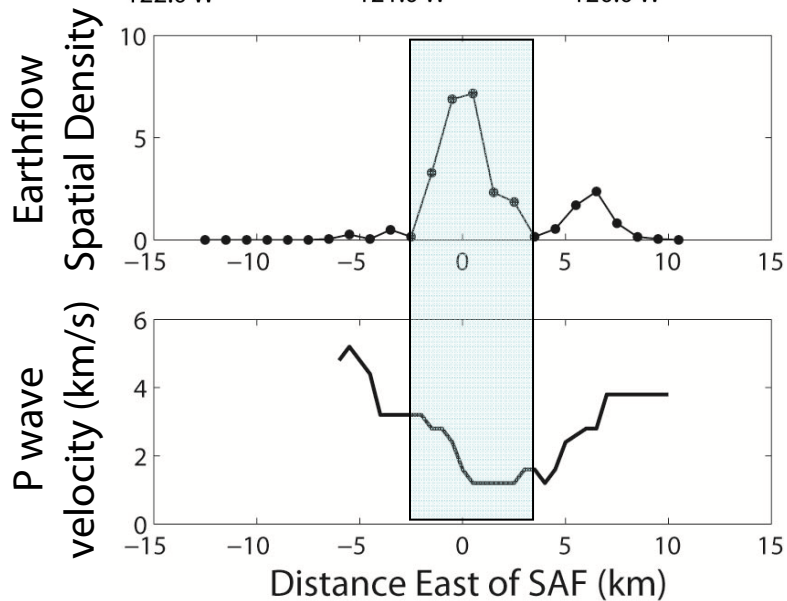
Fault damage zones



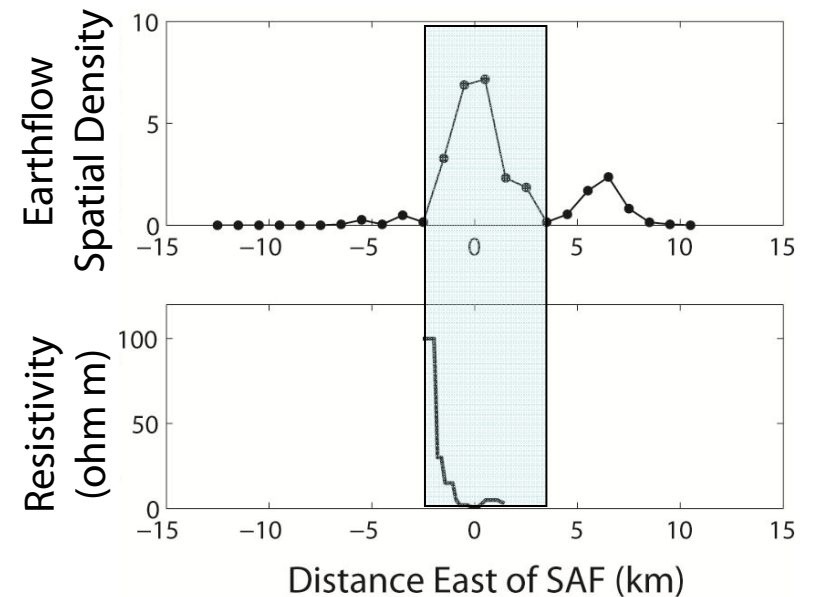
Seismic velocity profile
(Thurber et al, 1997)



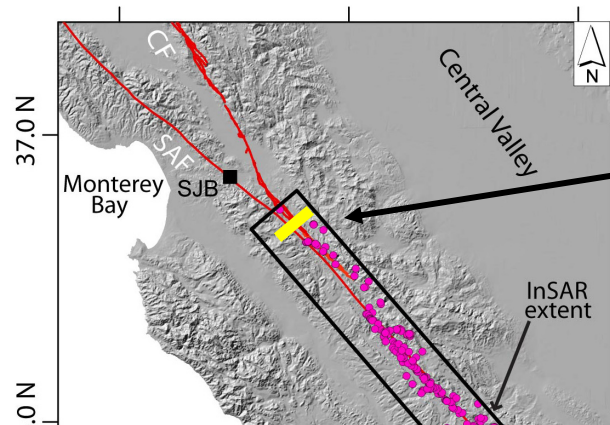
Electrical resistivity profile
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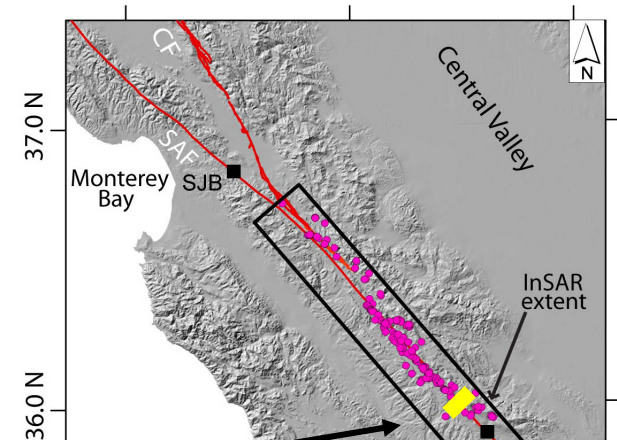
Decreasing
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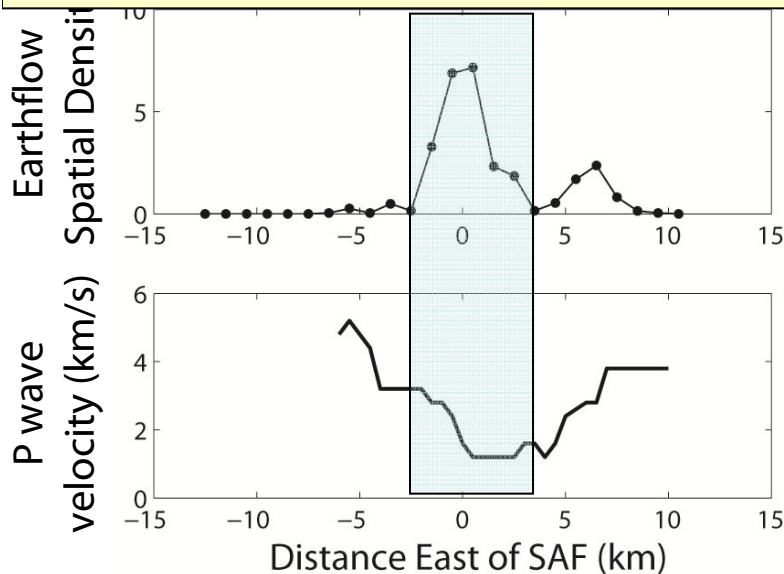
Fault damage zones



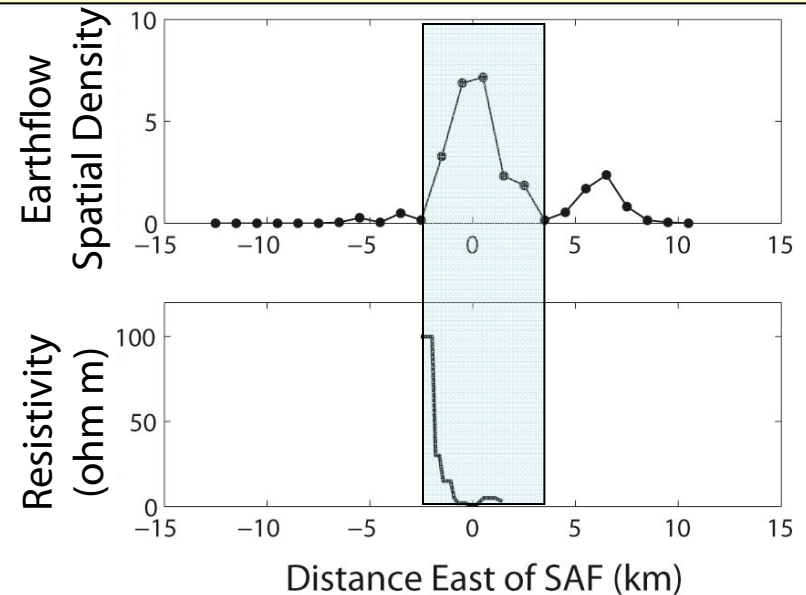
Seismic velocity profile
(Thurber et al, 1997)



Fault zone damage is observed to correlate with areas of high earthflow spatial density, and is **likely the first order control** on the cross-fault distribution of earthflows near the creeping SAF.

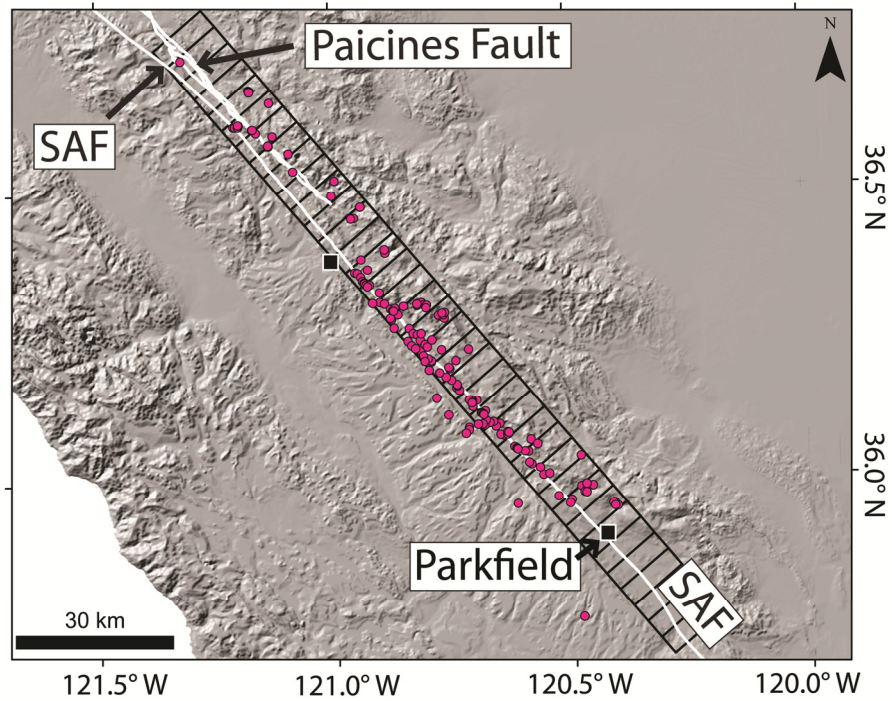


↓
Decreasing
rock strength



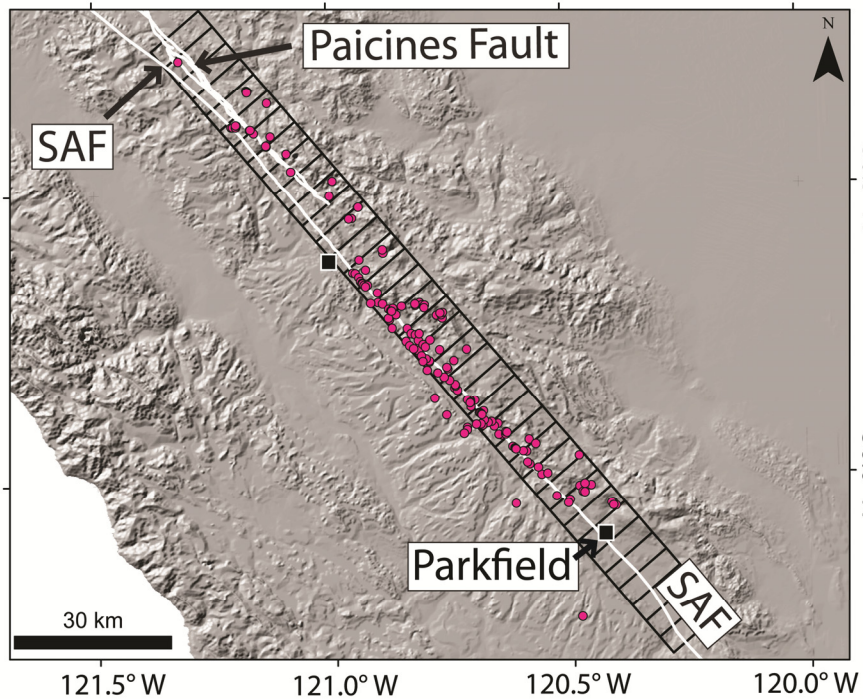
Swath Profile Analysis

SAF perpendicular swaths
(4 km x 12 km)

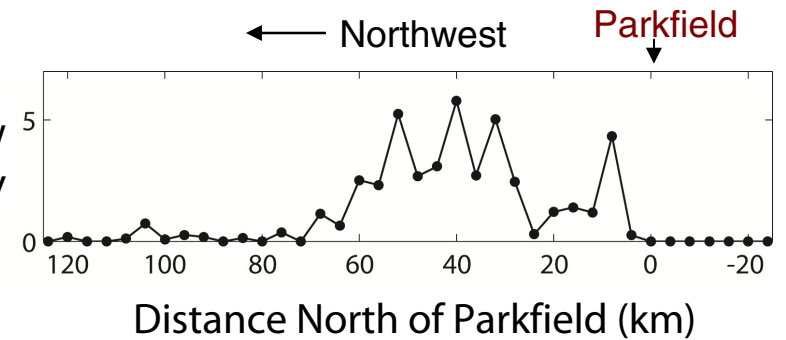


Swath Profile Analysis

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(4 km x 12 km)

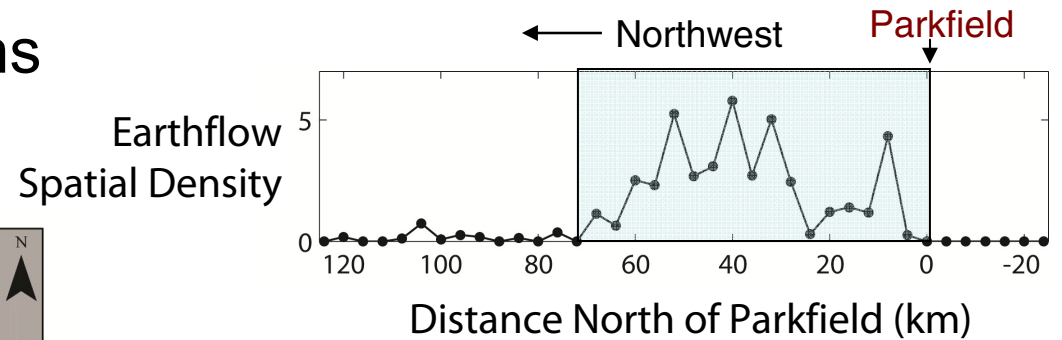
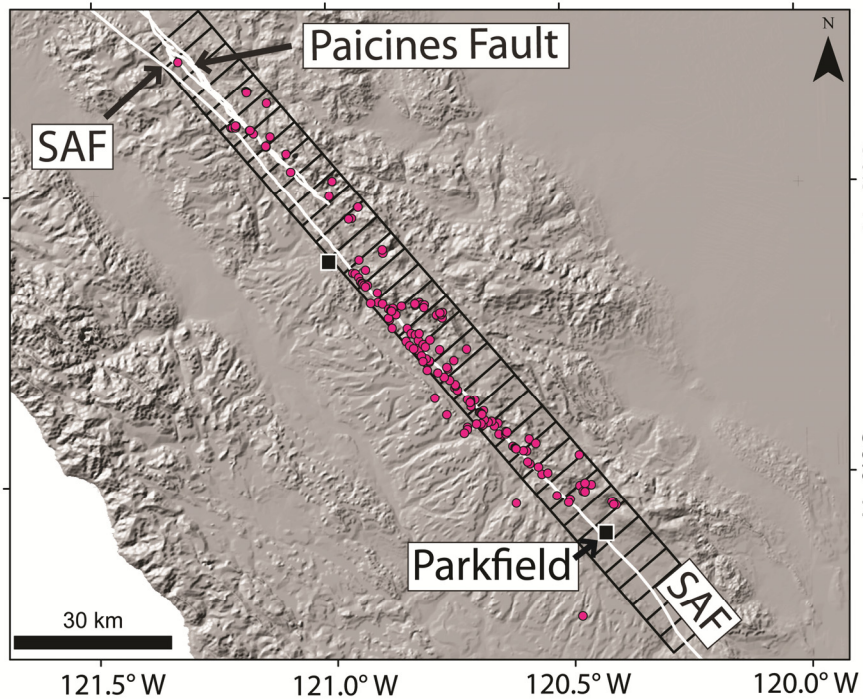


Earthflow
Spatial Density



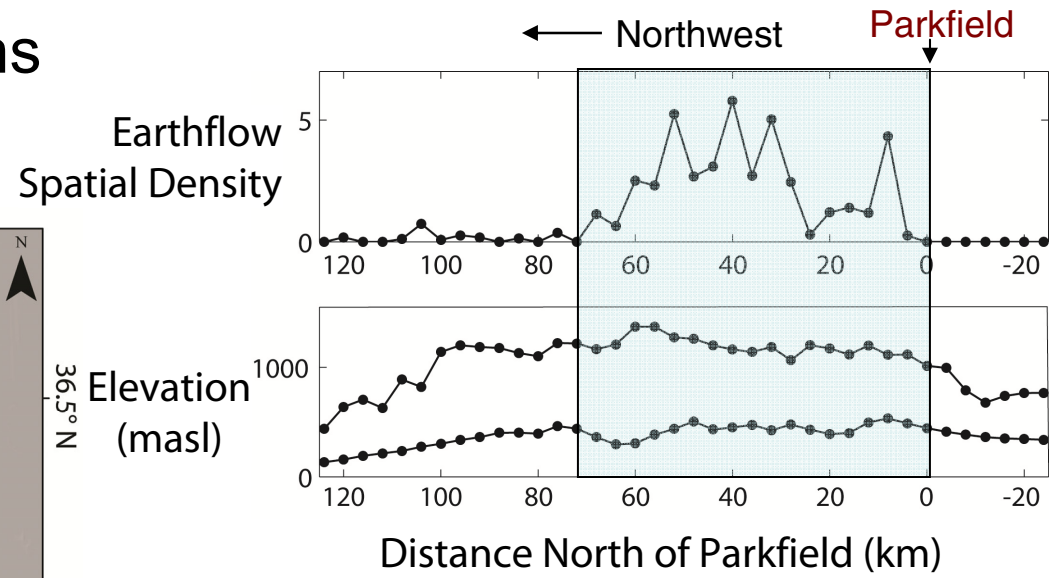
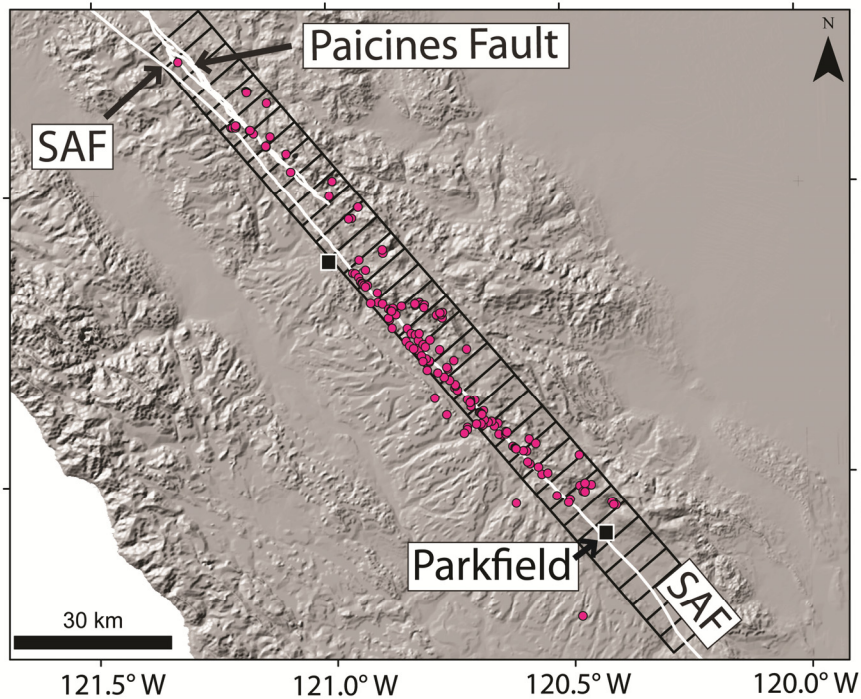
Swath Profile Analysis

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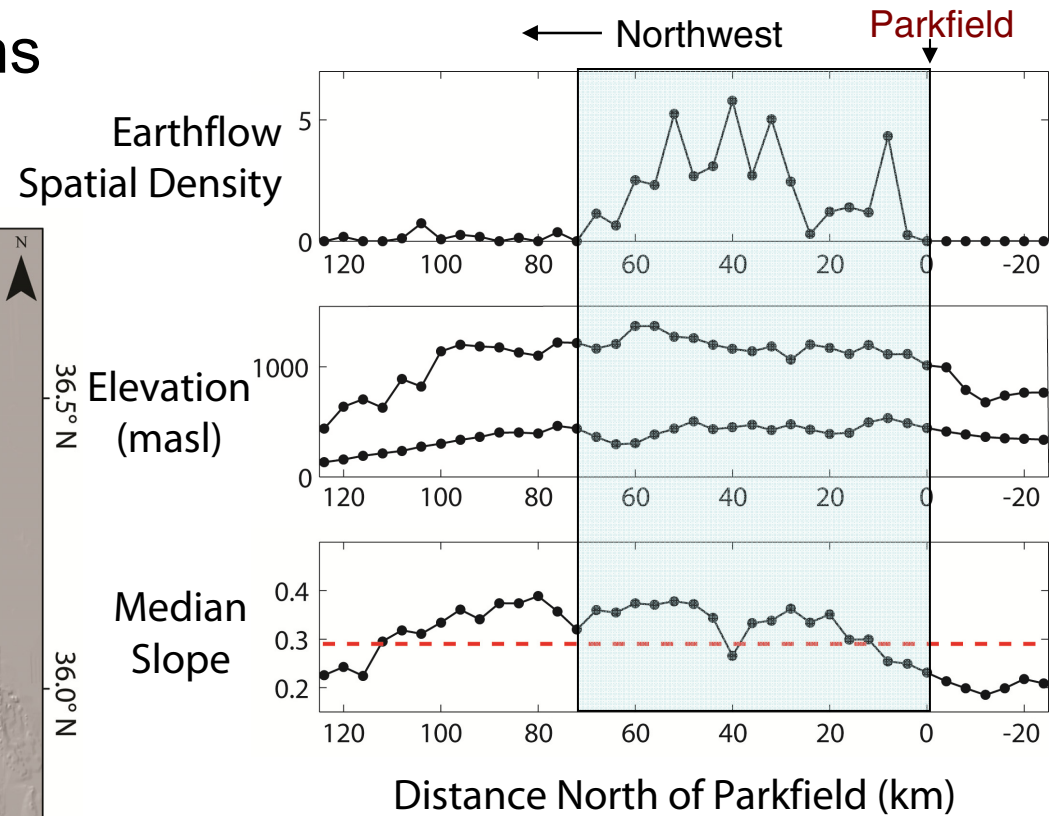
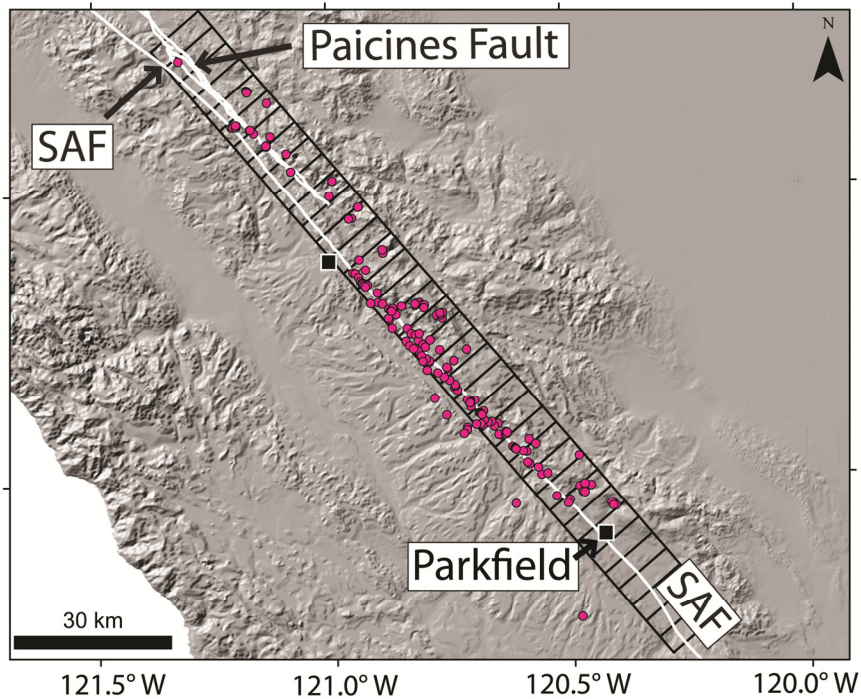
Swath Profile Analysis

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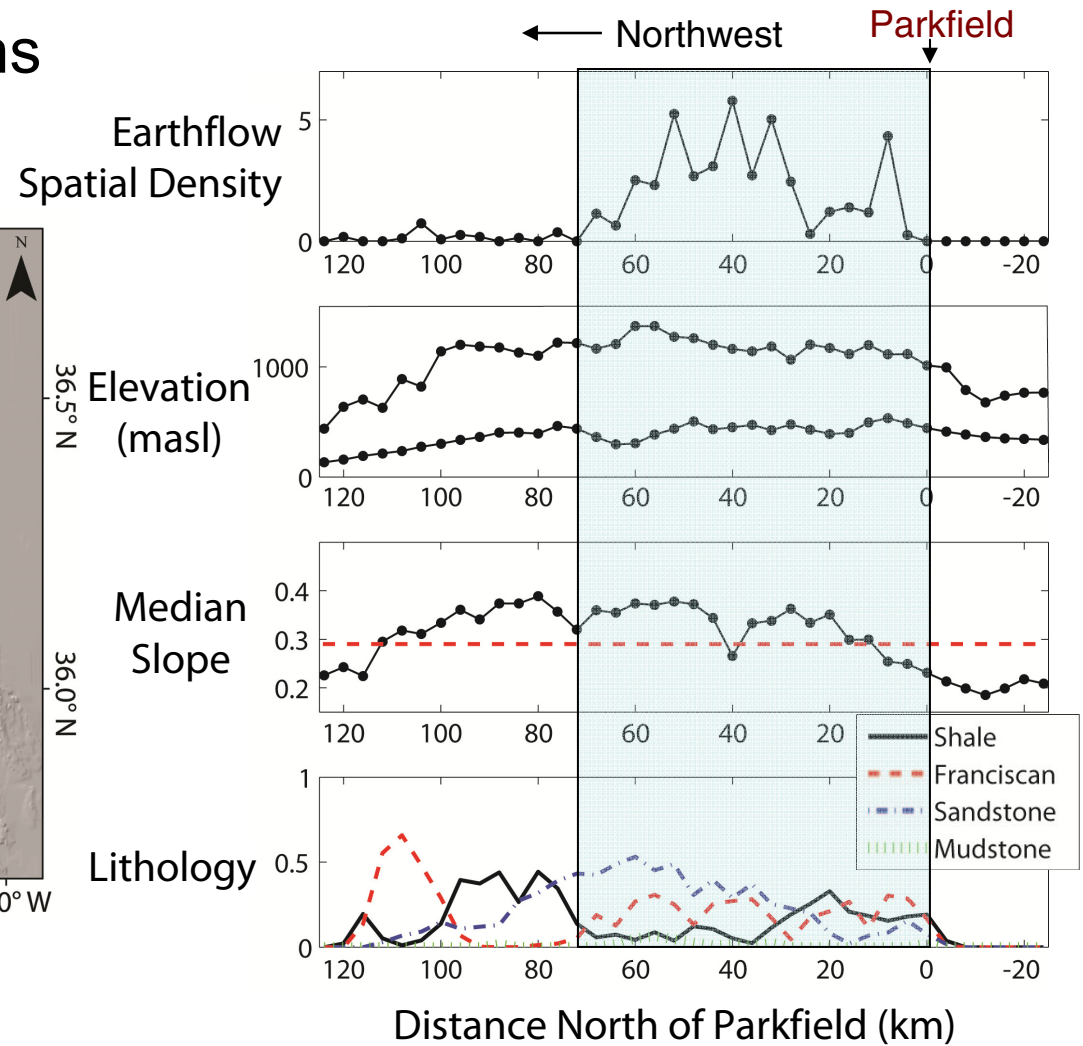
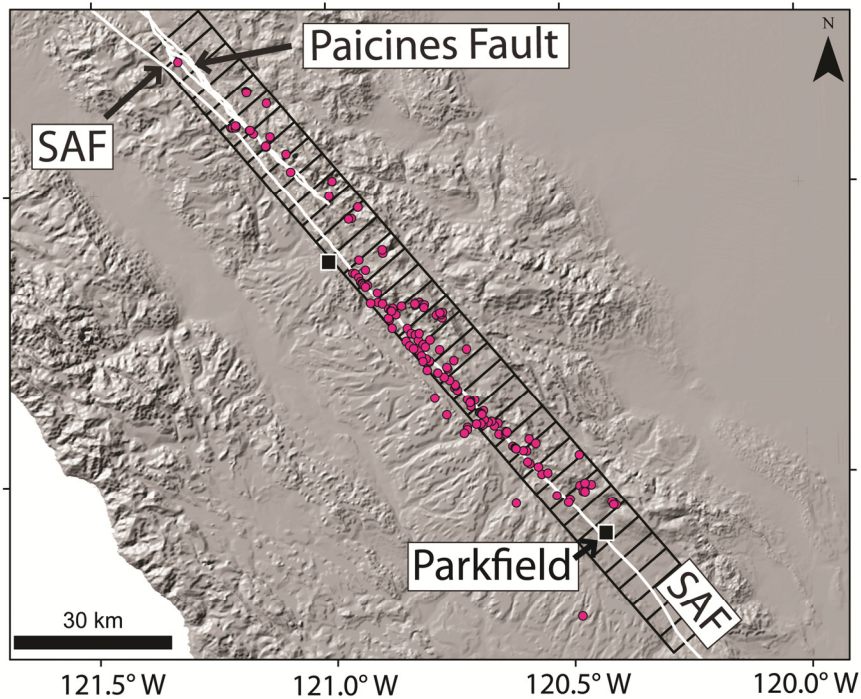
Swath Profile Analysis

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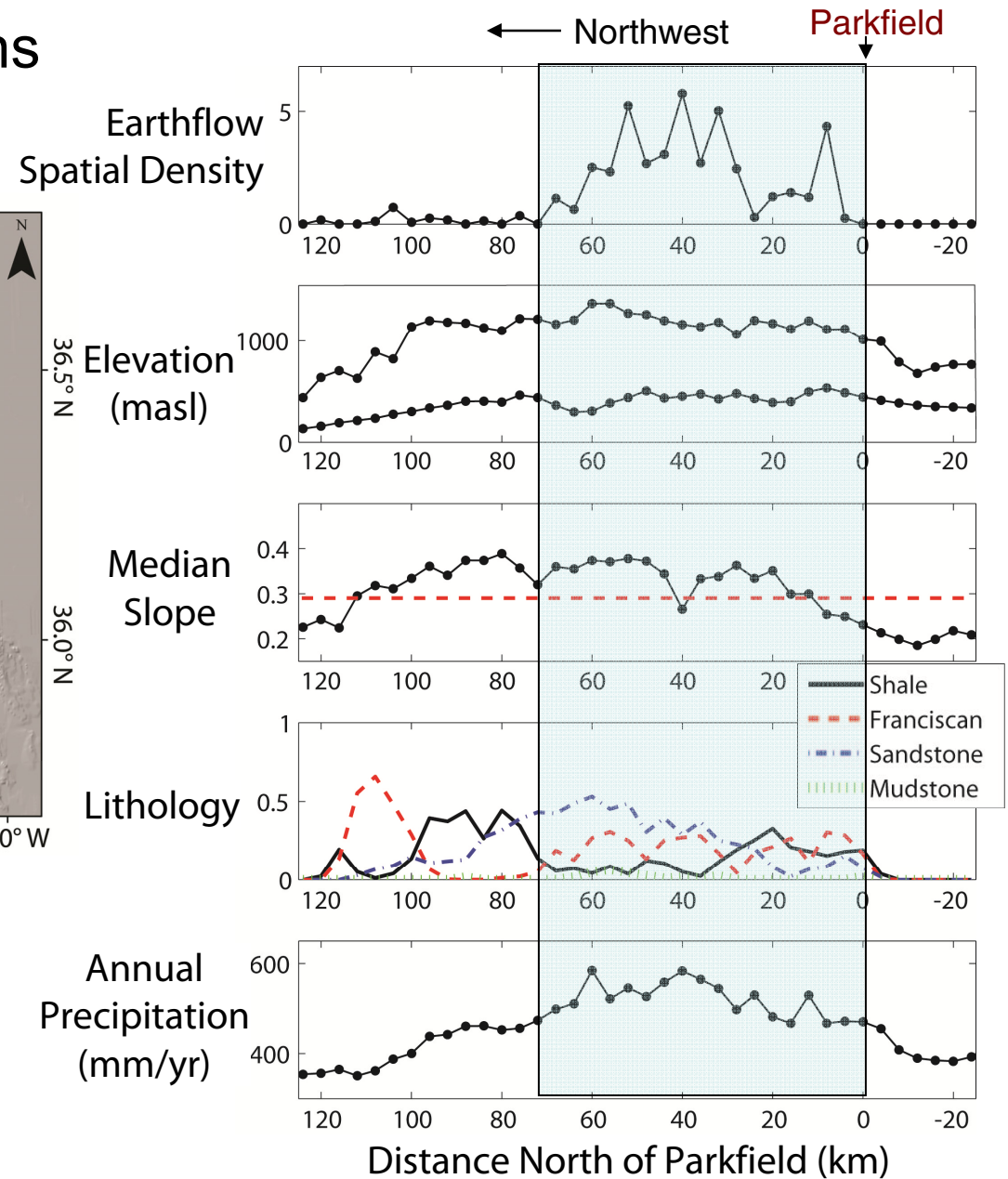
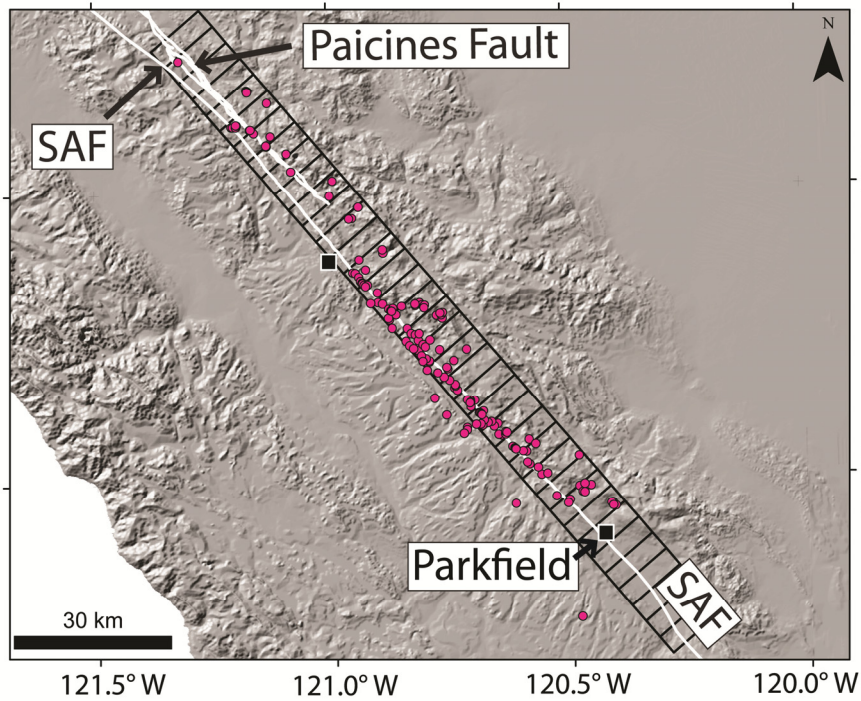
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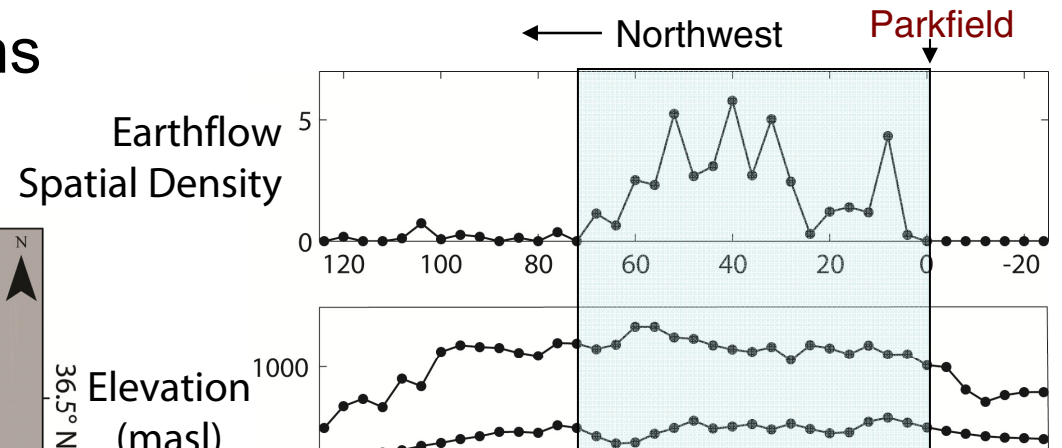
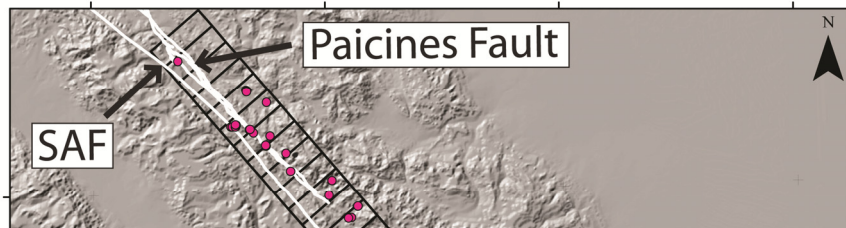
Swath Profile Analysis

SAF perpendicular swaths
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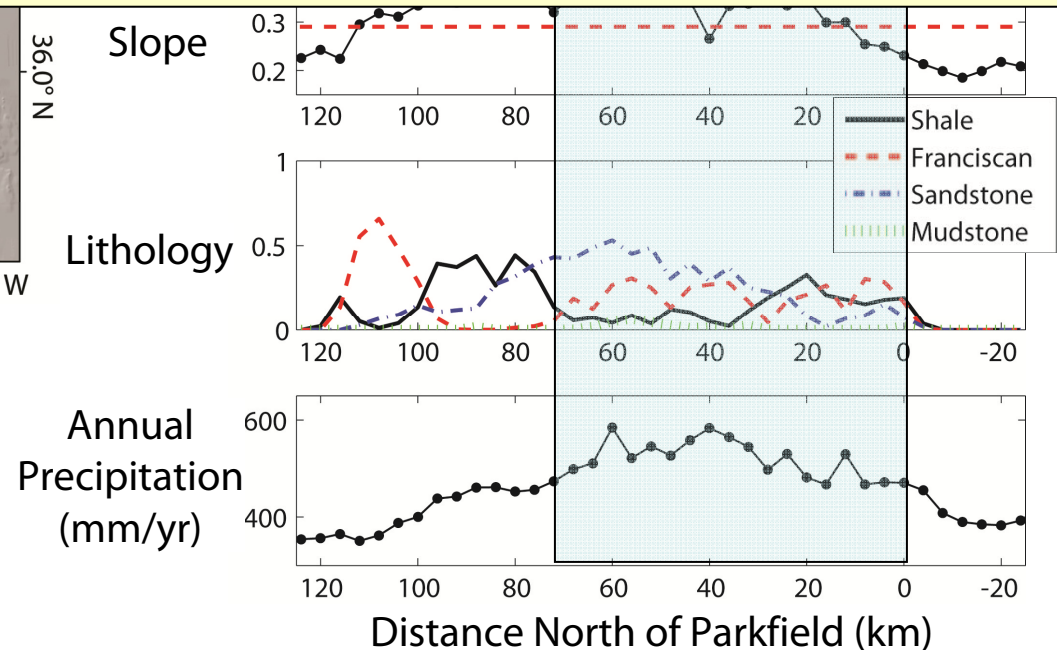
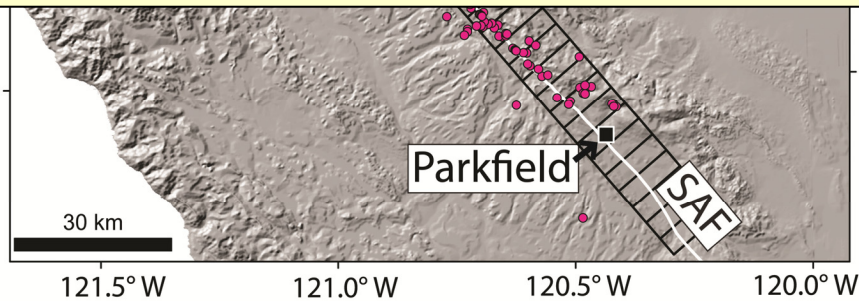


Swath Profile Analysis

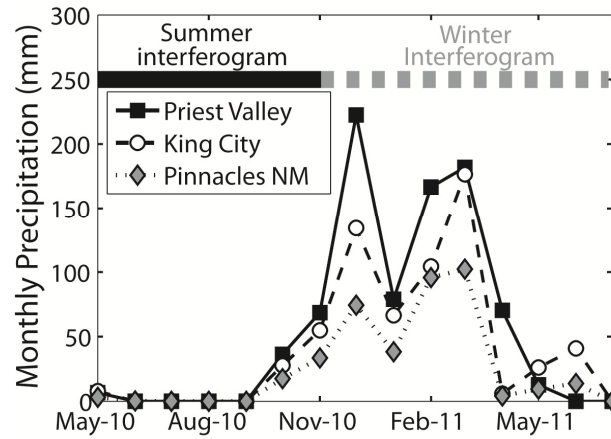
SAF perpendicular swaths
(4 km x 12 km)



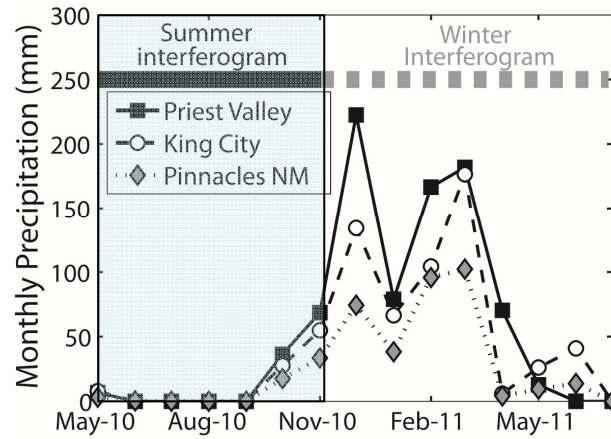
Annual precipitation may be an important **second order control** on earthflow spatial density after fault zone damage is accounted for.



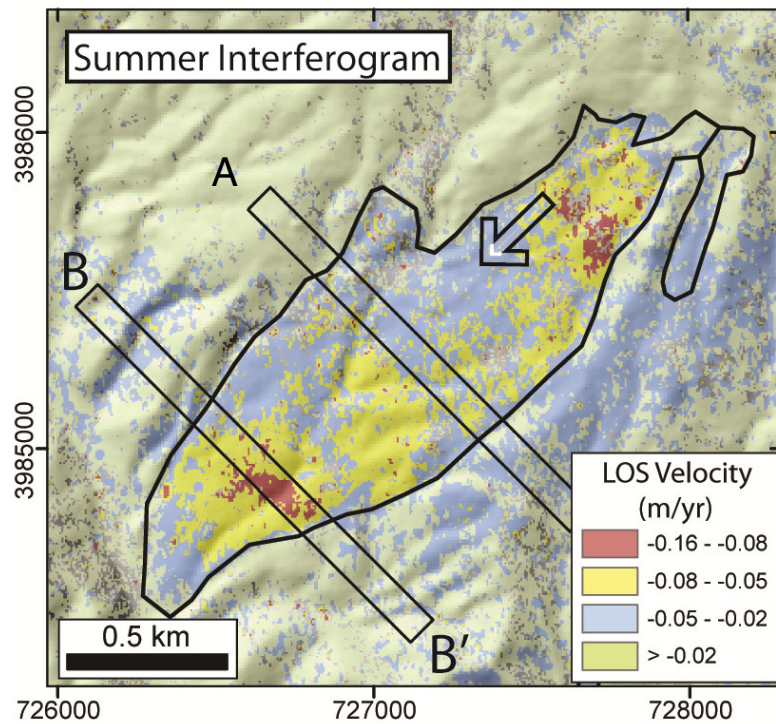
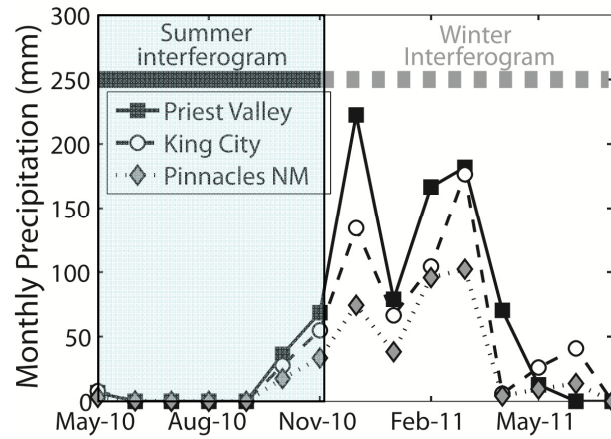
Movement in response to rainfall



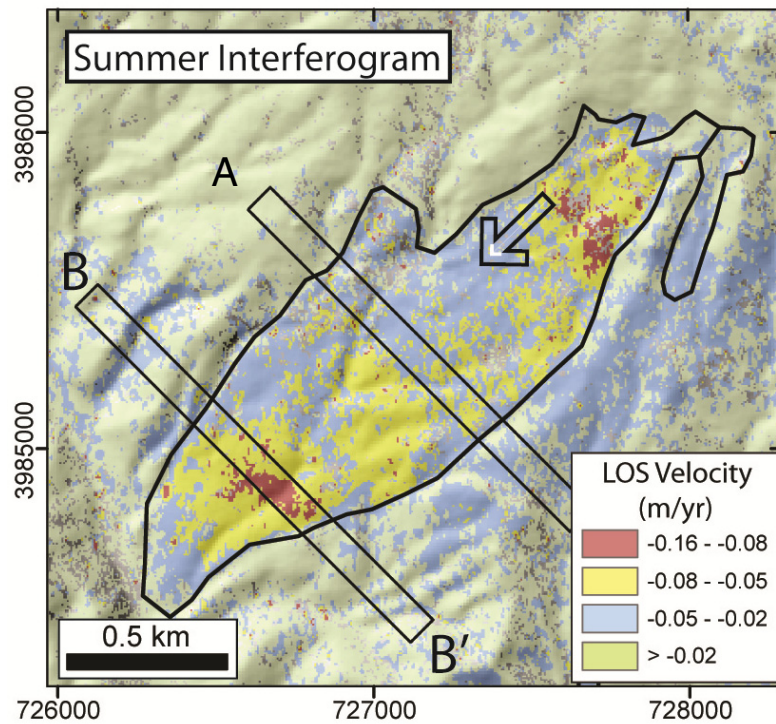
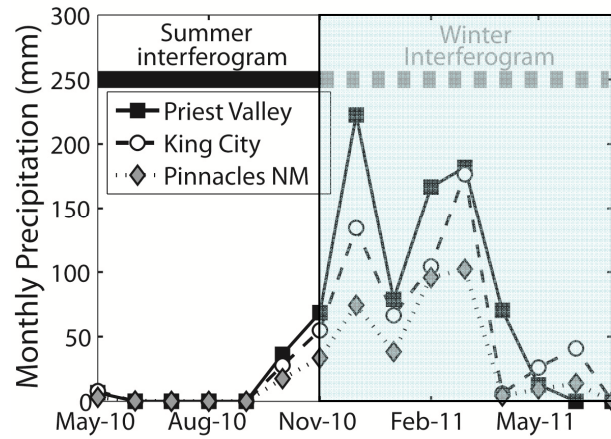
Movement in response to rainfall



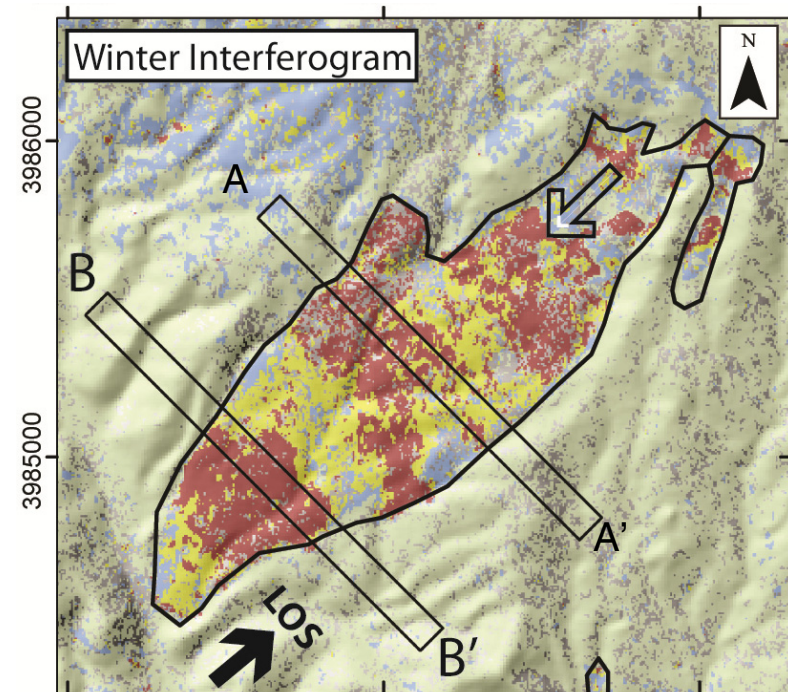
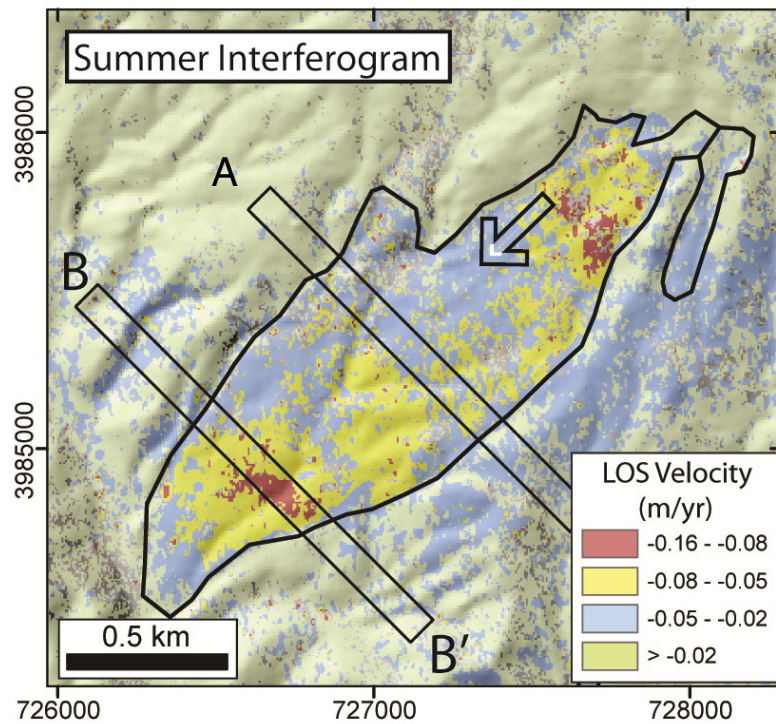
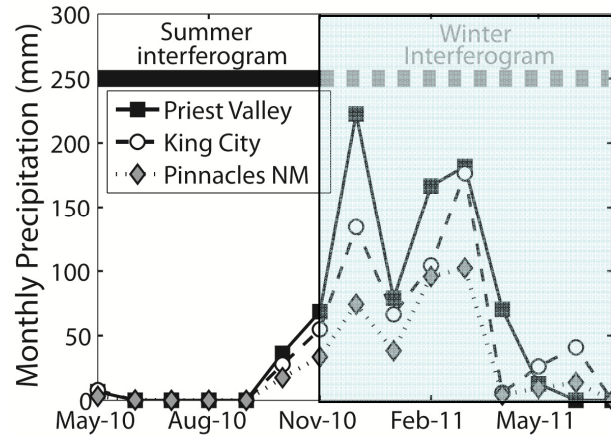
Movement in response to rainfall



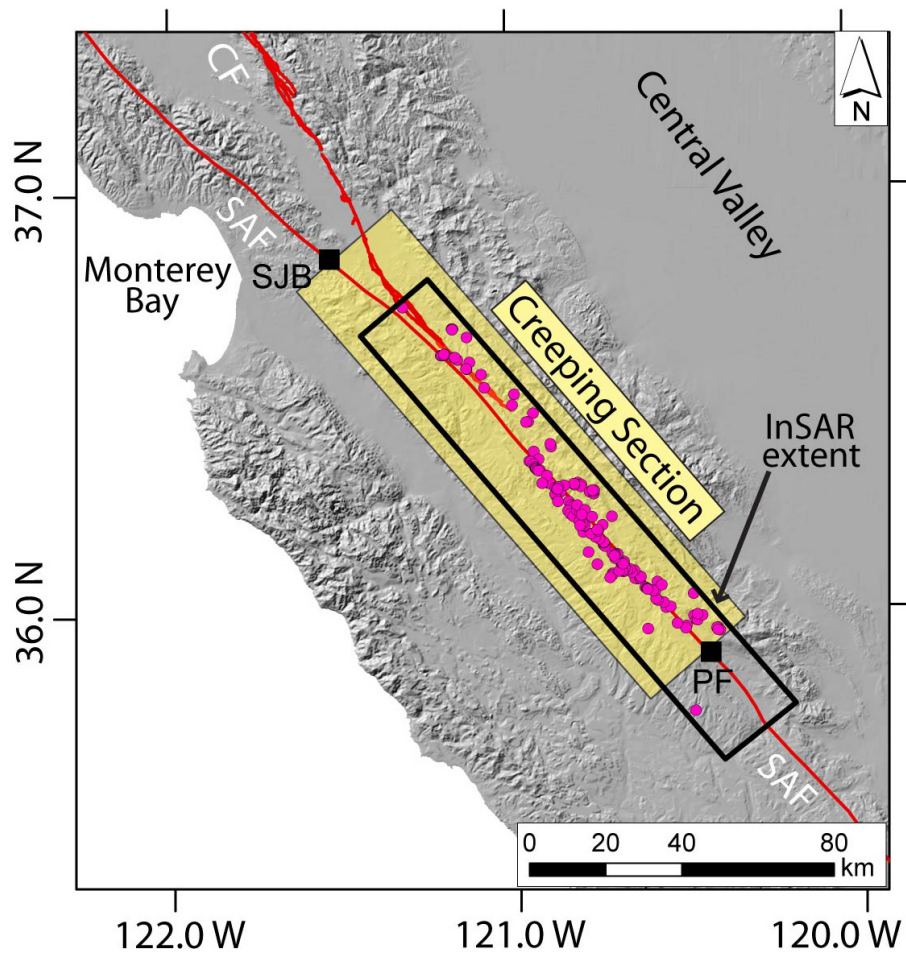
Movement in response to rainfall



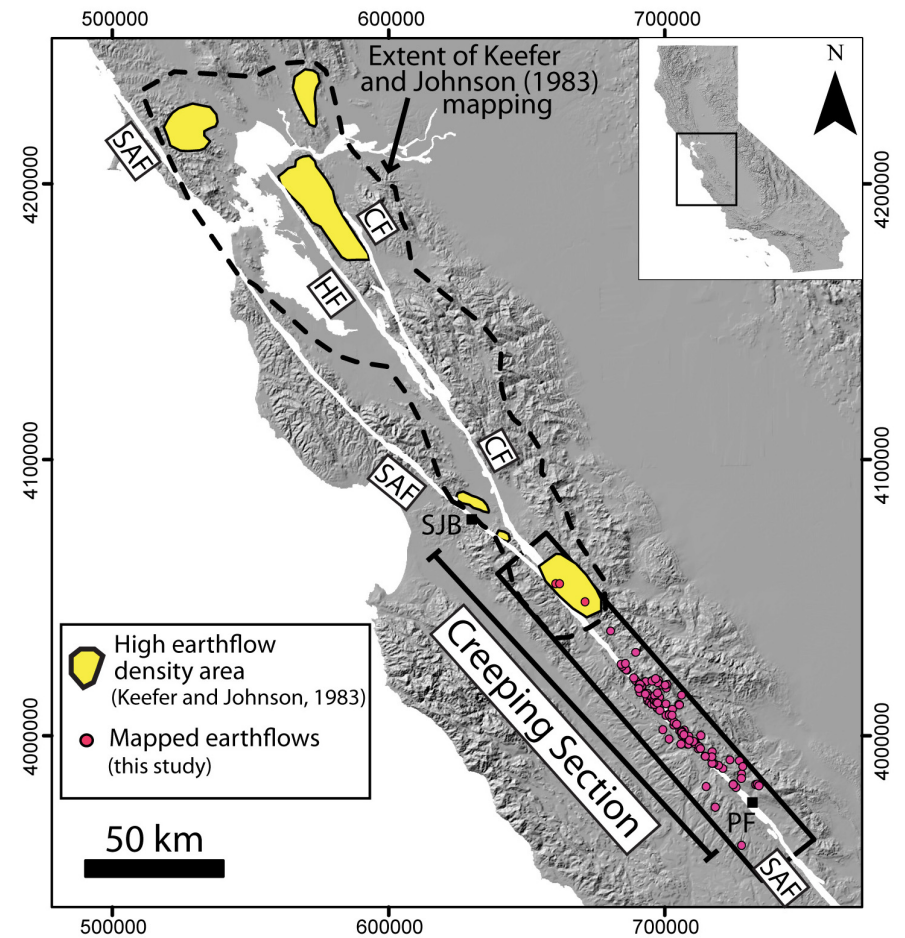
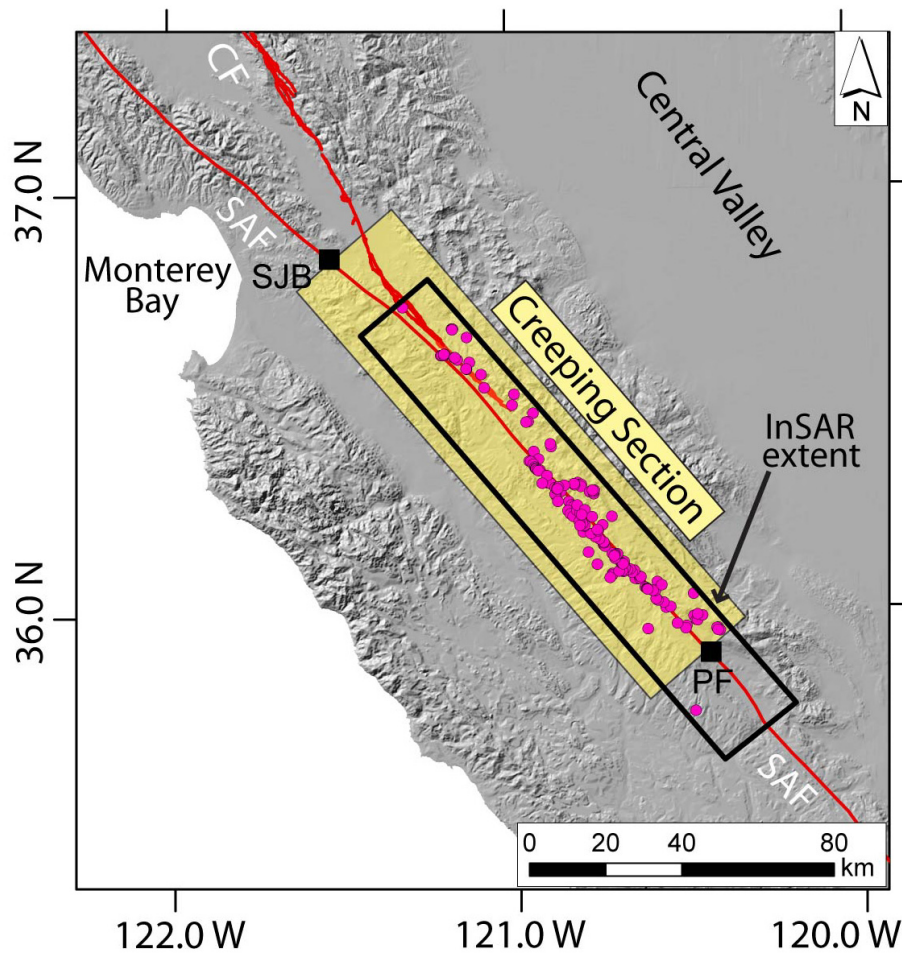
Movement in response to rainfall



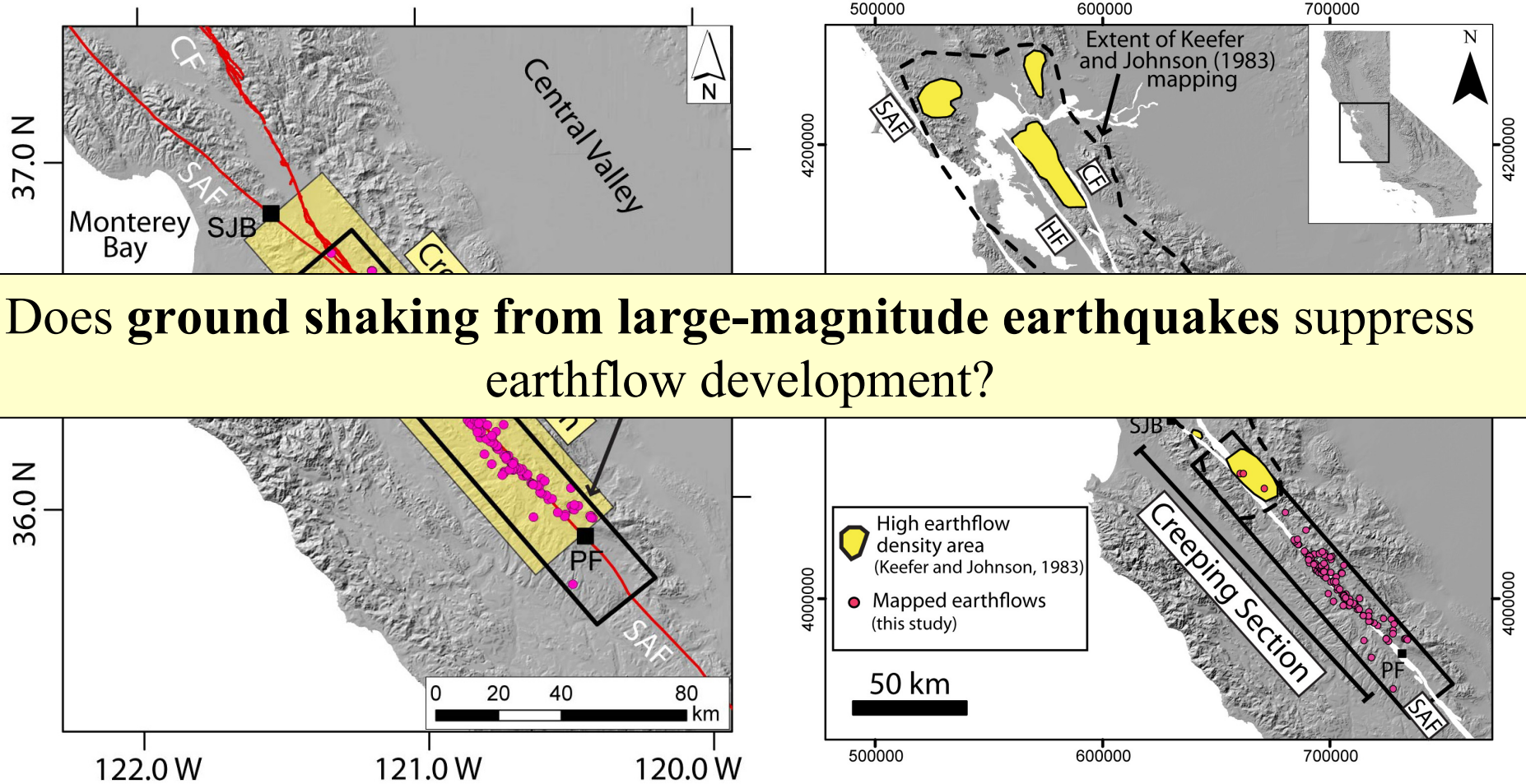
Do earthflows occur along the locked SAF?



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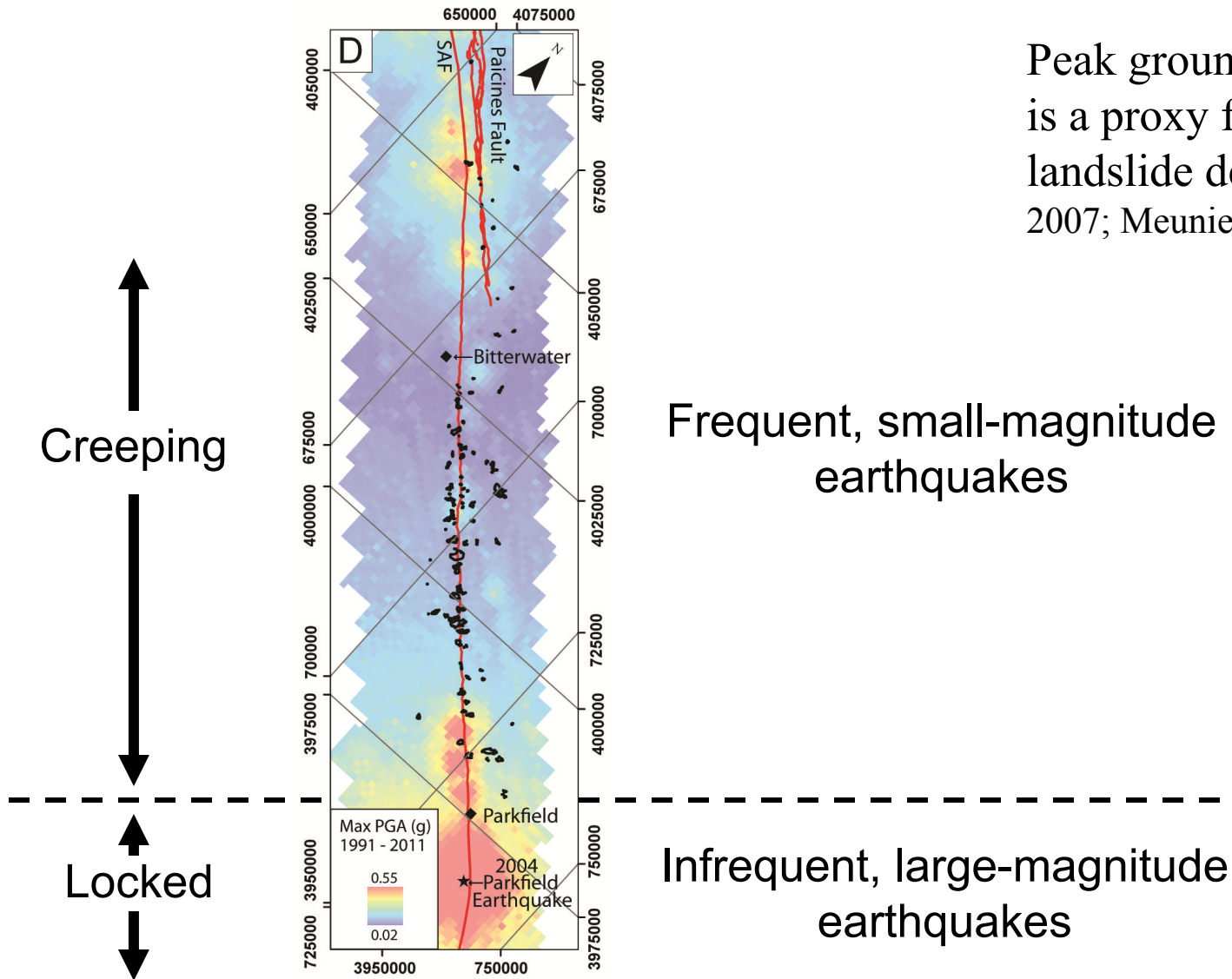


Does ground shaking from large-magnitude earthquakes suppress earthflow development?

Do large-magnitude earthquakes suppress earthflows?

Peak ground acceleration (PGA)
is a proxy for co-seismic
landslide density (Meunier et al.,
2007; Meunier et al., 2008).

Do large-magnitude earthquakes suppress earthflows?



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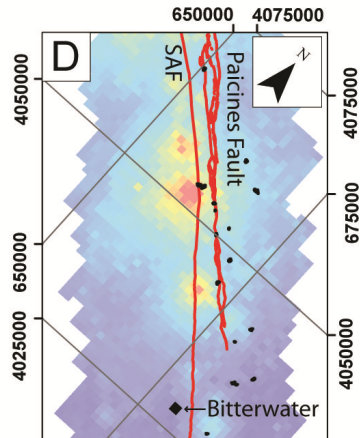
Frequent, small-magnitude earthquakes

Infrequent, large-magnitude earthquakes

Predicted maximum **peak ground acceleration**
42,731 earthquakes between 1991-2011 (NCEDC)
(Boore and Atkinson, 2007)

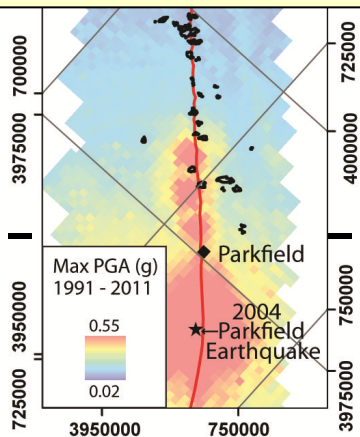
Scheingross et al, 2013, *GSA Bulletin*

Do large-magnitude earthquakes suppress earthflows?



Peak ground acceleration (PGA) is a proxy for co-seismic landslide density (Meunier et al., 2007; Meunier et al., 2008).

Anti-correlation of earthflows and peak ground acceleration is consistent with our hypothesis of co-seismic landslides limiting earthflow extent.

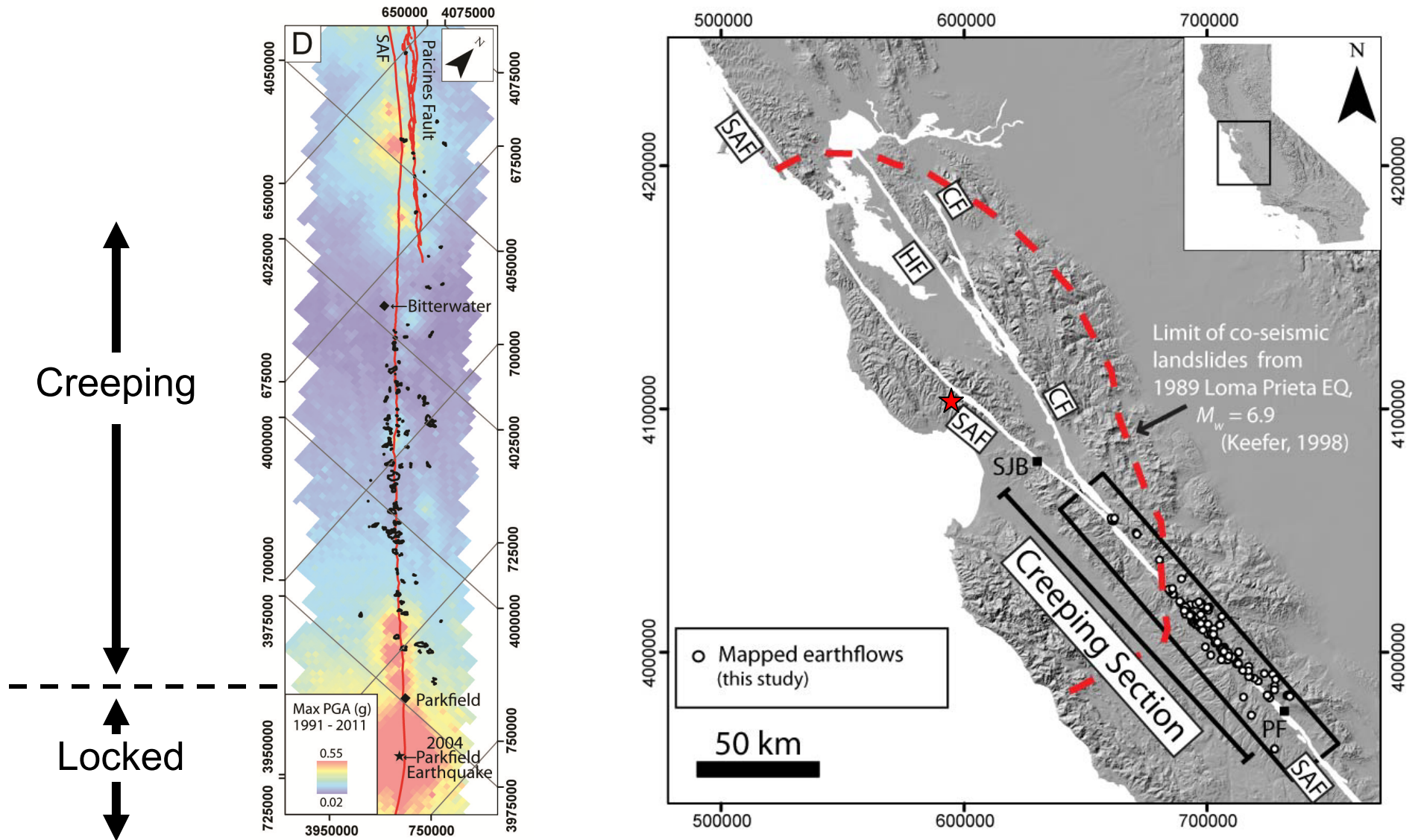


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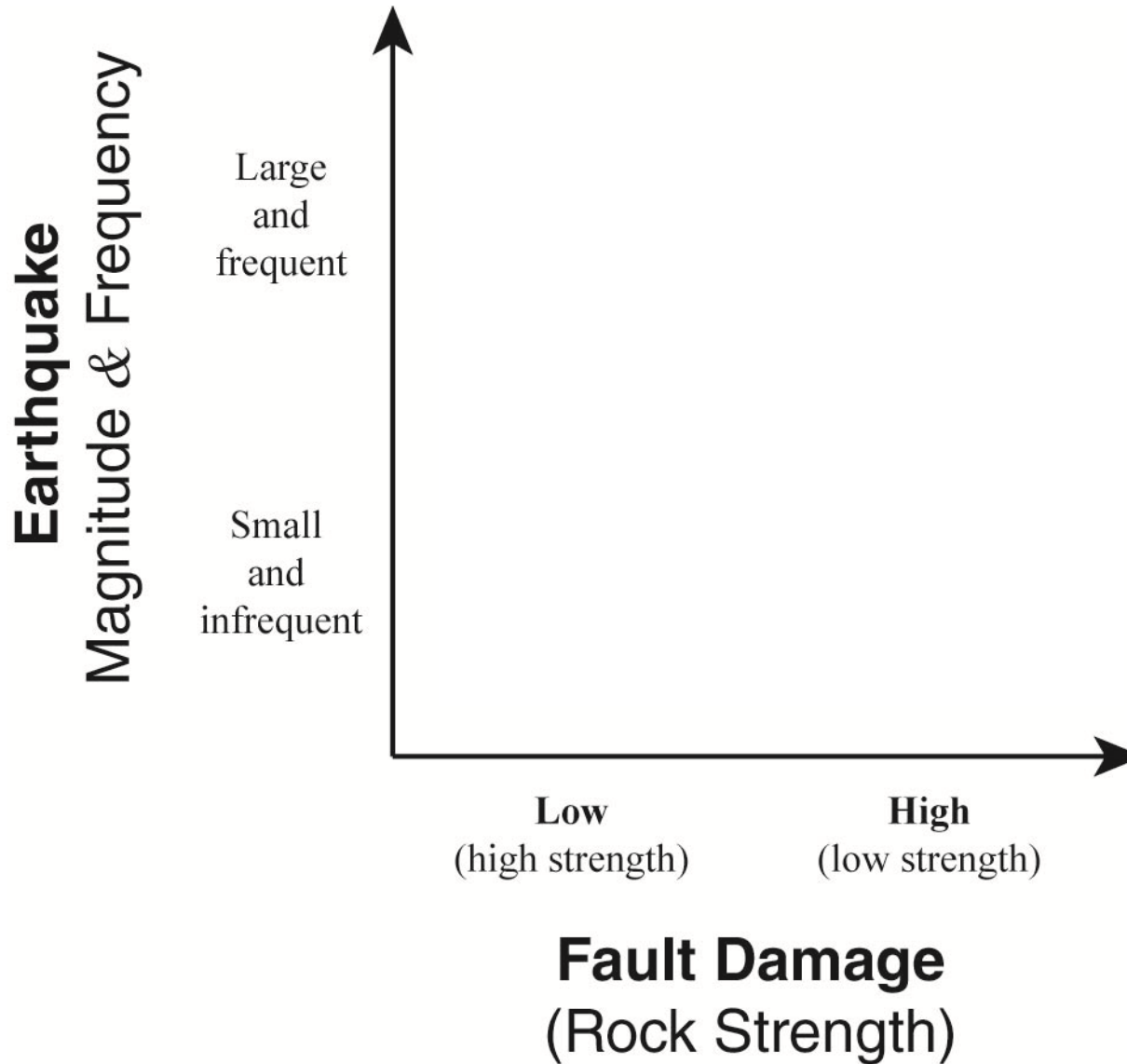
Do large-magnitude earthquakes suppress earthflows?



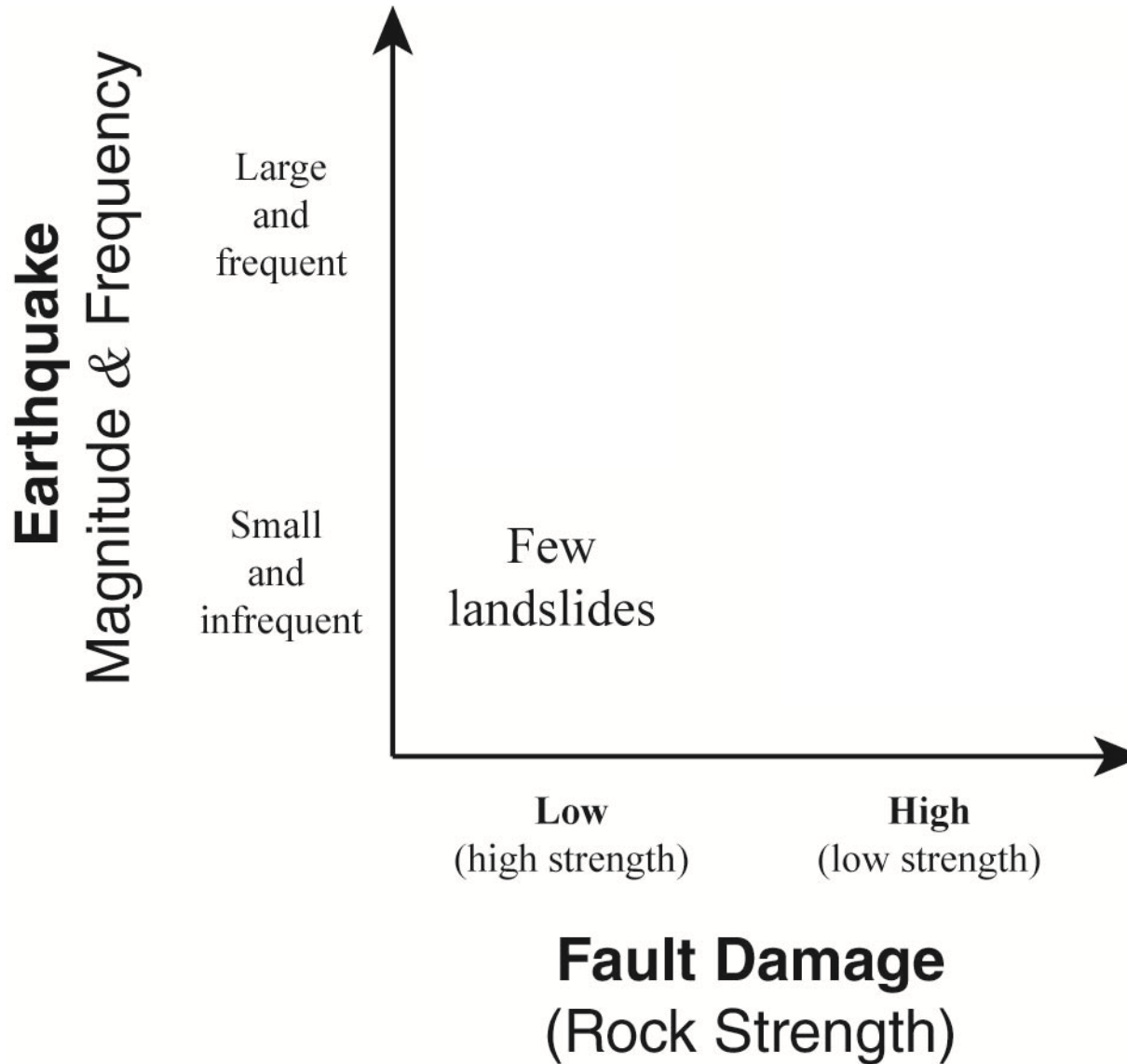
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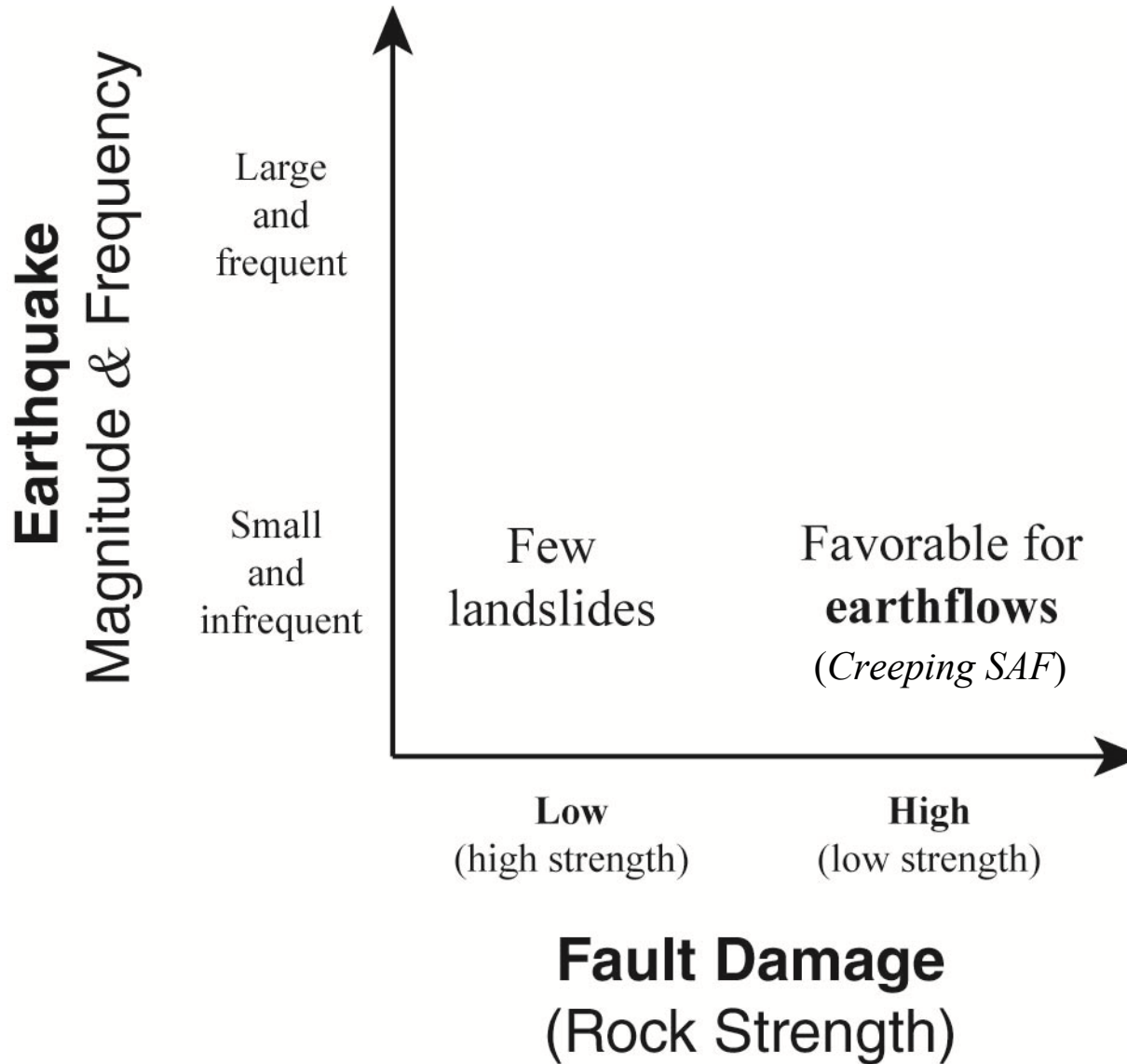
Fault damage versus earthquakes



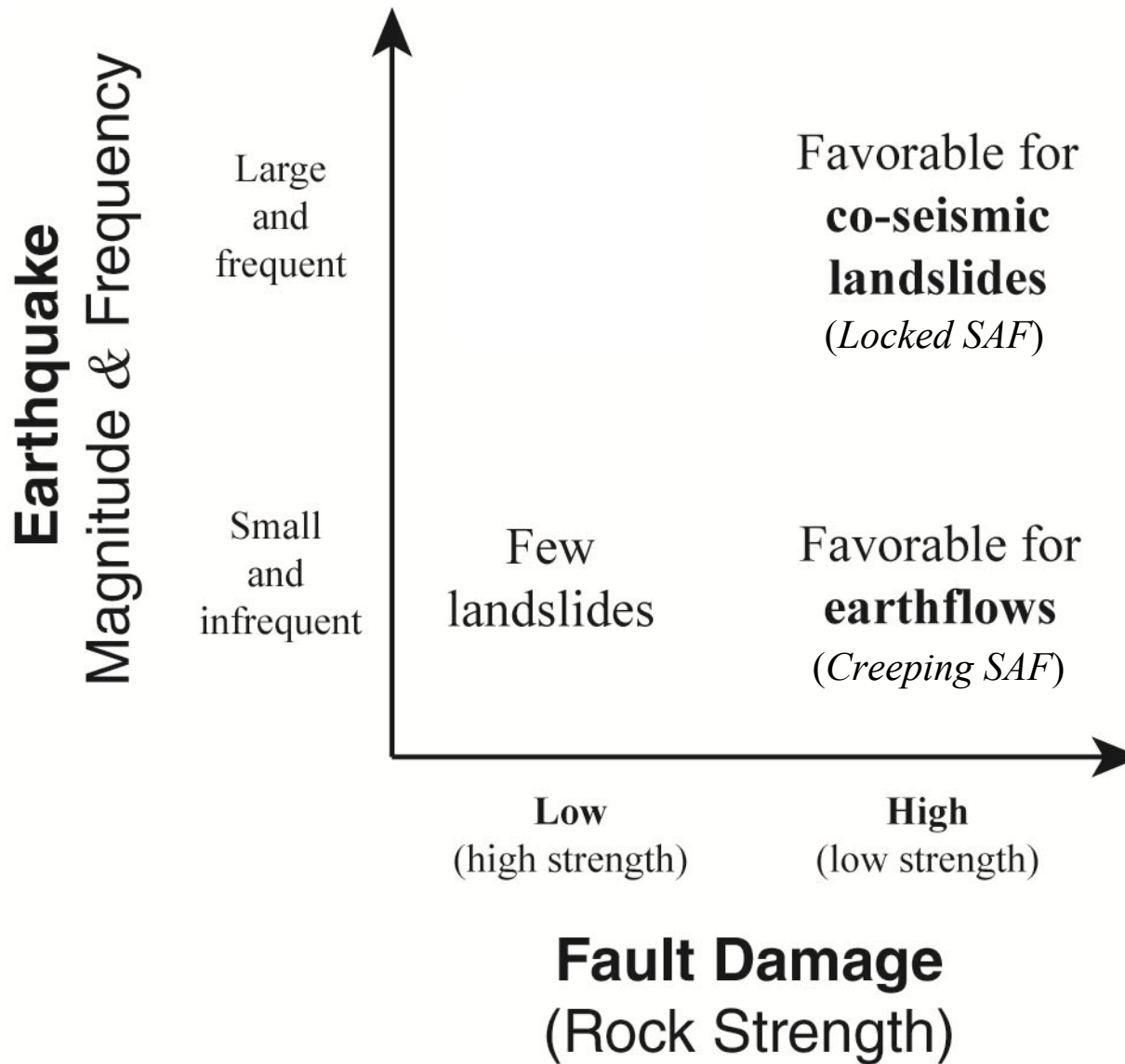
Fault damage versus earthquakes



Fault damage versus earthquakes



Fault damage versus earthquakes



Conclusions

- Faulting introduces competing influences that can both **PROMOTE (via fault damage zones)** and **SUPPRESS (via large-magnitude earthquakes)** the occurrence of slow-moving landslides.
- Along the central SAF, fault zone damage and large-magnitude earthquakes appear to be the primary controls on the occurrence of slow-moving landslides.
 - Predictions of earthflow spatial distribution in other tectonically active landscapes should account for these variables.
 - Total annual precipitation and seasonal variation in precipitation may be important secondary controls on earthflow movement.